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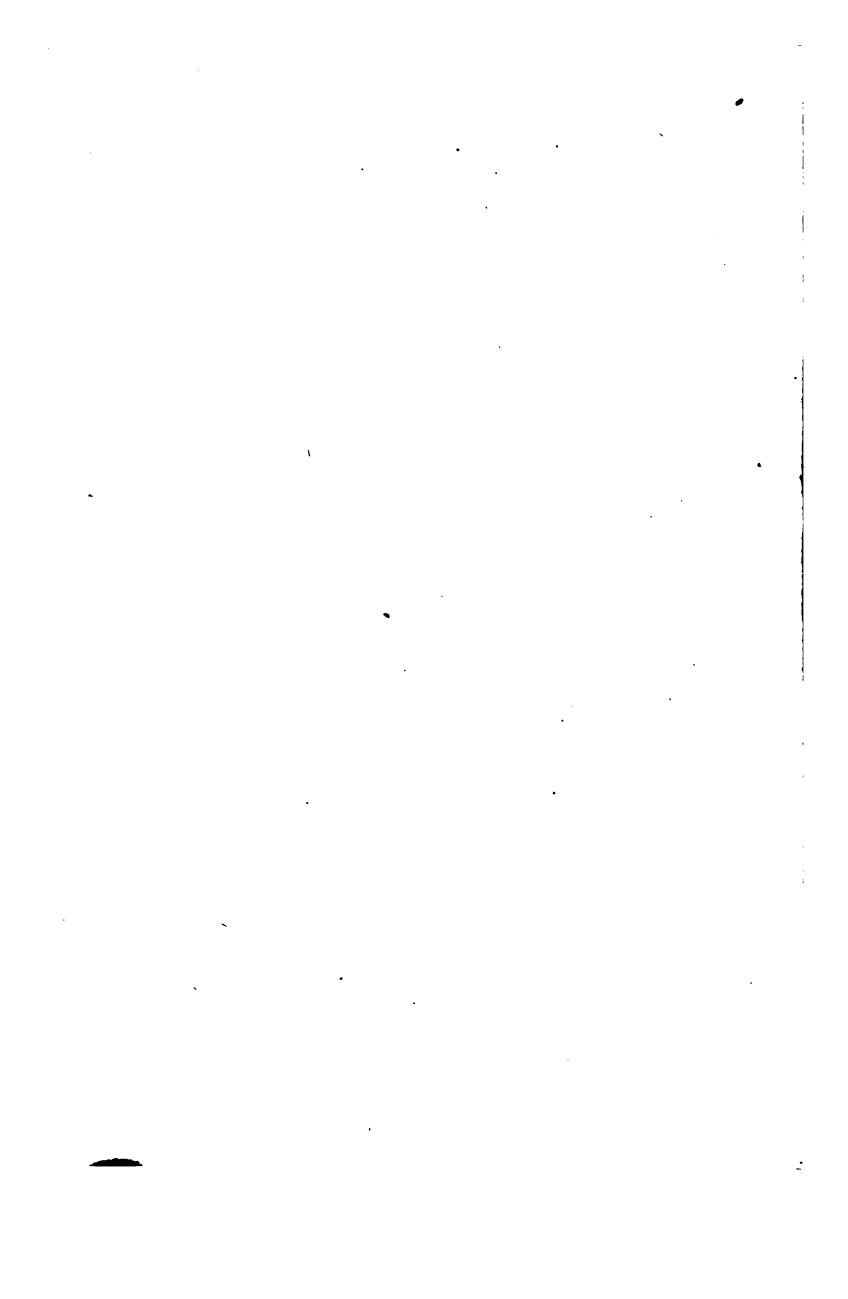


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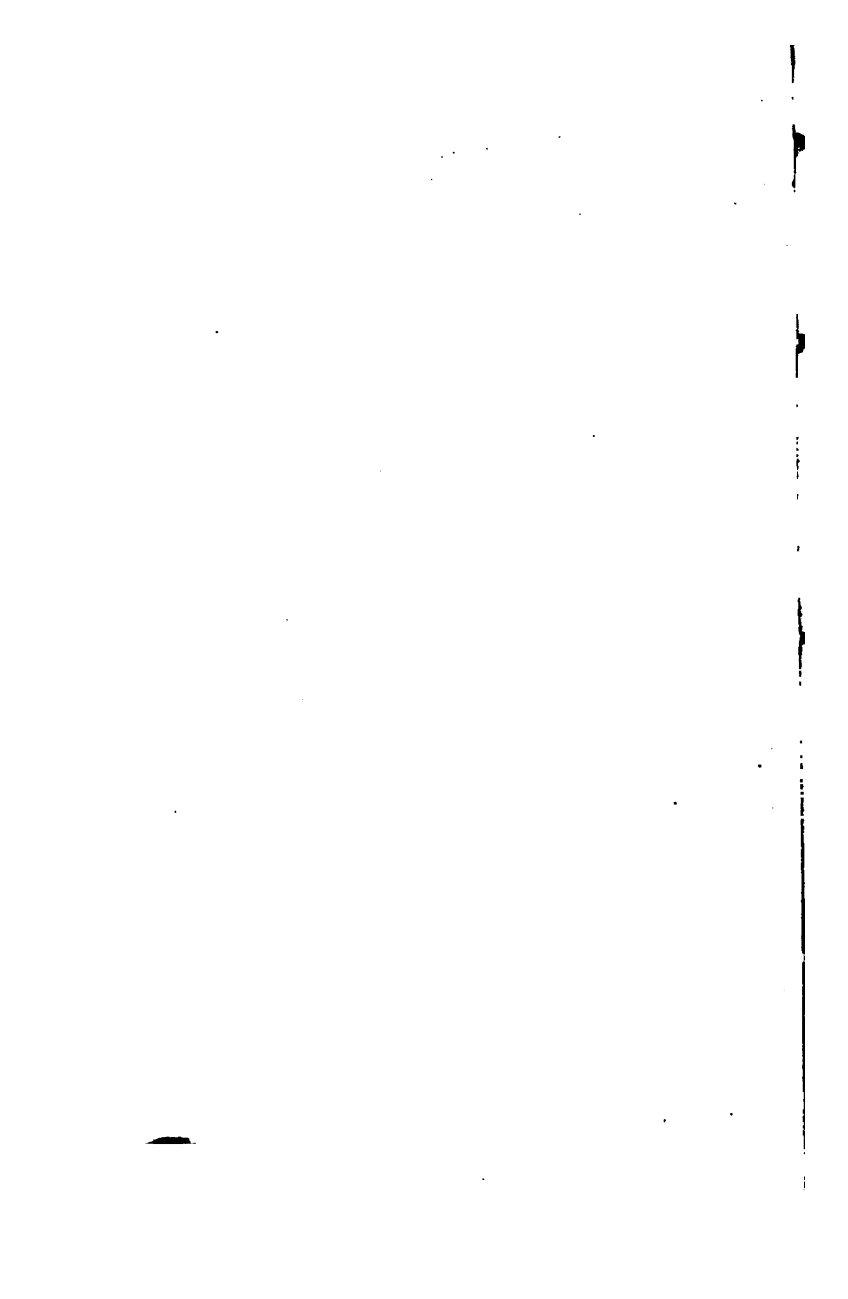
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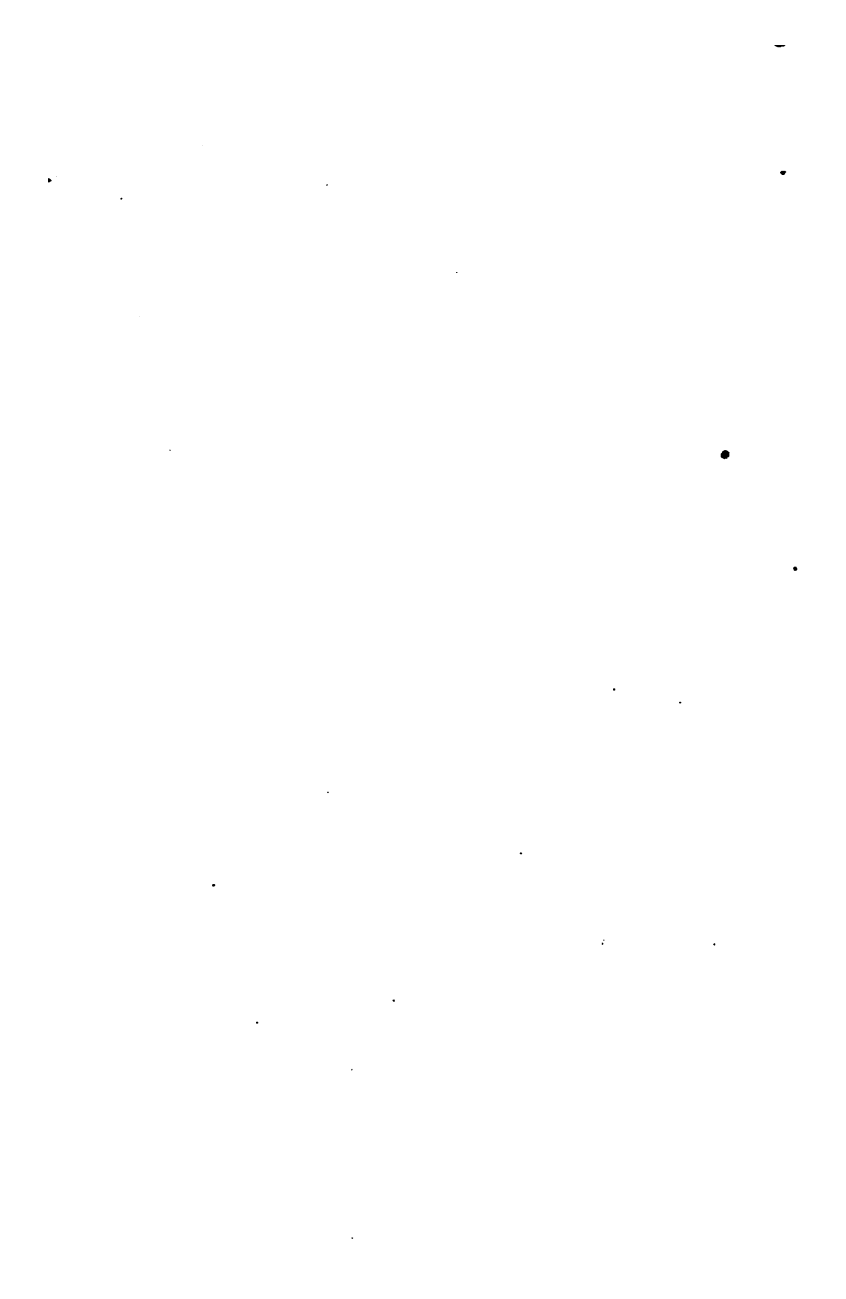






ELEMENTARY LESSONS
IN
PHYSICAL GEOGRAPHY.







OCEAN-WAVES. (After J. M. W. Turner.)

◦
ELEMENTARY LESSONS
IN
PHYSICAL GEOGRAPHY.

BY
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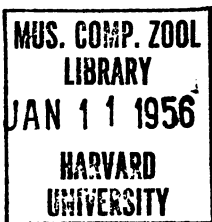
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P R E F A C E.

A SIMPLE but systematic and yet interesting description of the familiar features of the Earth's surface may be regarded as a natural introduction to the teaching of science. Treated from this point of view, Physical Geography may be made a valuable instrument of education. To give it such importance the most advantageous method is to make use of the common knowledge and experience of the pupils, and, starting from this groundwork, to train them in habits of observation and in scientific modes of thought and inquiry among every-day phenomena. From the very outset the instruction should be as far as possible practical. A shower of rain, the flow of a brook, the muddy water of a river, the shape of a cliff, the outlines of a mountain, the undulations of a plain—these and the thousand other common features of landscape should be eagerly seized by the teacher and used as vivid illustrations of the broad fundamental

principles which it will be his main object that his pupils should thoroughly master. Thus employed, Physical Geography is not learnt as an ordinary task, but rather becomes a delightful recreation, in which, however, the observing faculty is exercised, the power of induction cultivated, and the imagination kept constantly active.

Having been long convinced that such a method of instruction would place this branch of science upon a firmer and broader footing in our educational system, and would moreover prove of great service in fostering a spirit of observation and reflection even among children, I projected many years ago the *Primer of Physical Geography*, which has been recently published in the series of *Science Primers*. The rapid sale of large impressions of that little work encourages the hope that the method advocated has been found successful in practice.

The present volume may be regarded as a further development of the same plan of instruction. As its title implies, it still deals mainly with the broad elementary questions of Physical Geography. It would have been impossible to find a place within its Lessons for the treatment of every branch of the wide subject, and as impossible, had it even been desirable, to

bestow equal fulness upon every branch for which room has been made. I have devoted most space to those aspects of the science which, in my own experience, have been found best suited for practical instruction. While as much general information as may be feasible should be communicated to them, young pupils cannot of course be expected to find the same interest in all divisions of the subject. It is of far more consequence to awaken in them a taste for such pursuits, and lead them to carry on the study of their own accord, than to try to charge their memories with dry facts and figures which, in the absence of intelligent and suggestive teaching, are too commonly meaningless and repulsive.

While, therefore, adhering to a systematic treatment, I have been led to dwell, for example, on the phenomena of the atmosphere at much greater length than is usual in elementary class-books. These phenomena are among the most familiar and universal features of the globe; examples of them can be constantly adduced, and they may thus be used with singular advantage to illustrate how the facts of science are observed and its laws are deduced.

I acknowledge with pleasure my obligations to my friend Mr. Buchan, who not only kindly allowed me

to make use of his Charts of Atmospheric Pressure and Temperature (which I have sought to render more effective by the use of shades of colour), but who also read over the proof-sheets of the first two chapters and gave me valuable suggestions on subjects regarding which he is so high an authority.

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ELEMENTARY LESSONS

IN

PHYSICAL GEOGRAPHY.

INTRODUCTION.

1. At night when the sky is clear, the largest stars seem to stand out in front, with others less in size and brightness crowding behind them. As we gaze into these depths still remoter and feebler twinkling points appear, until at last our eyes can no longer shape out any distinct specks of light. Such a sight impresses our minds, as nothing else can do so vividly, with the vastness of the Universe. We feel how comparatively small must after all be the distance which we can see into that "star-dust" which has been sown through the regions of space. And even when with the aid of a good telescope we return to these same skies, it is to find more cause than ever to acknowledge how immeasurably vast is that part of the Universe which man can thus explore, but at the same time to meet again with a vague limit, beyond which we cannot see, not because we have reached the utmost verge of creation, but because our instruments

can carry our vision no further. Far beyond that limit it may be that the regions of space contain other stars and systems, though too remote ever to be brought into view even by the most powerful telescopes which human skill could construct.

Astronomers have calculated the distance of some of the largest and nearest stars. But their figures, expressing sums of many millions of miles, are too vast to carry any definite ideas to our minds. When we reflect that each of those stars, from the brightest to the faintest twinkling point, is really a sun, many of them, no doubt, far vaster in size than our sun, but dwindled into such seeming feebleness by reason of their inconceivable distance, we cannot but feel what a little speck of dust, in comparison, must this dwelling-place of ours really be which we call the Earth.

2. It is useful to get this comparative insignificance of the Earth firmly realized by our minds. And in no way can this be done so well as by watching the starry sky, and learning what has been discovered regarding the motions, sizes, and distances of the heavenly bodies. What then is the Earth in relation to these bodies? Has it always been in the same condition as now, or has it perhaps passed through long ages of change and progress? Mankind has had a long and varied history. May not the Earth itself have had one also? If so, can we learn anything about the story of the Earth?

3. Again, when, on the other hand, we look upon the face of the Earth by day, how boundless and varied it seems! From the district in which we may chance to live, we can pass in thought to the country at large, then to other countries, and then to the idea of the whole wide globe, with its continents and oceans, its mountains, valleys, and plains, and all that wonderful diversity of form and colour which makes its surface so unceasingly beautiful.

4. This variety is everywhere associated with life and movement. Consider, for instance, the unvarying succession of day and night ; the orderly march of the seasons ; the constant blowing of the winds ; the regular circling of the ocean tides ; the ceaseless flow of rivers ; the manifold growth and activity of plant and animal life ! Surely it was no strange thought when men in old times pictured this world as a living being. And even though we cannot look on the earth as a living thing in the sense in which a plant or animal is so called, yet in view of all that multitudinous movement which is ever in progress upon its surface, and on which, indeed, we know that our own existence depends, there is evidently another sense in which we may speak of the life of the Earth.

5. Now this Life of the Earth is the central thought which runs through all that branch of science termed Physical Geography.¹ The word geography, as ordinarily used, means a description of the surface of the earth, including its natural sub-divisions such as continents, and oceans, together with its artificial or political sub-divisions, such as countries and kingdoms. But Physical Geography is not a mere description of the parts of the earth. It takes little heed of the political boundaries, except in so far as they mark the limits of different races of men. Nor does it confine itself to a mere enumeration of the different features of the surface. It tries to gather together what is known regarding the Earth as a heavenly body, its constitution, and probable history. In describing the parts of the earth—air, land, and sea—it ever seeks so to place them before our minds as to make us realize, not only what they are in themselves, but how they affect each other, and what part each plays in the general system of our globe. Thus Physical Geography endeavours to

¹ This term as here used is synonymous with Physiography, which has been proposed in its stead

present a vivid picture of the mechanism of that wonderfully complex and harmonious world in which we live.

6. Many of the facts with which this branch of science deals are familiar to everyone. With nothing, for example, have we better acquaintance than with the air which surrounds us. We have breathed it every moment of our lives ; we know it at rest as well as in storms ; we have watched the growth in it of mists and clouds, as well as the fall of dew and rain. At first there might seem hardly anything more perfectly known to us, and therefore regarding which we should have so little new to learn. But Physical Geography, taking up the subject as one great whole, strives to show how all these different features and changes of the air are connected together, to explain their causes, and to point out the essential part which the air takes in the great movements which affect both sea and land. Thus one of the great advantages of this study lies in the very commonplace character of the phenomena of which it treats. We cannot go anywhere without meeting with illustrations of some of the great lessons which it teaches. When we have once learnt these lessons, therefore, and have learnt too how to apply them as we go, we carry an additional source of pleasure in every walk and journey. We get into the habit of using our eyes, and noting an endless variety of things which would otherwise have been passed by unseen, and this habit of observation, besides the pleasure its exercise brings with it, will be found of no slight service in the ordinary affairs of life.

7. In completing its picture of the Life of the Earth, Physical Geography must necessarily draw its illustrations from all parts of the world. It thus brings before us phenomena of which we may have no experience in our own country, and which very probably we may never have any opportunity of seeing for ourselves. In the

study of it, we journey, so to speak, all over the world, and learn in a short time far more about the world than we should be able to do from reading a few ordinary books of travel. Indeed, a good treatise on Physical Geography may be regarded as a condensed and well-arranged book of travels in all countries, with this distinction, that although it has no personal adventures to describe, it enables us to understand how one region of the earth differs from another, and explains these differences by connecting them all together with the general principles or laws on which they depend. So that, although a man may never have been in India or Africa, or the Arctic regions, he may from the study of Physical Geography have a far better notion of the general features of these countries, and why they differ from each other so much in climate, than many other men who have travelled to, or even lived for years in them. It is a matter of no little encouragement for all of us to know that the more we watch what takes place around us in our own country, and the more thoroughly we understand it, the more easily do we realize what goes on in other and distant parts of the world.

3. By taking up the study of Physical Geography, not merely as a subject to be learnt from books, but as a practical pursuit to be followed out by our own observations, as opportunity offers in the course of our daily occupations, we make most progress in it and get the largest amount of pleasure out of it. This is the spirit in which the following chapters are written. They are not meant merely to describe the different parts of the earth, in such a way as to enable you to learn these most easily by heart, but rather to incite you to use your own eyes, and to examine, compare, and contrast what you see to take place from day to day. They are so arranged as to begin, where practicable in each sub-division, with our common

knowledge. Then they point out what can be ascertained on the subject by our own simple observation and experiment. Lastly, they present such further information as may be acquired from the observations and travels of men who have given much time and thought to the collection and investigation of the facts. In the case of the Air, for example, starting from what each one of us knows by everyday experience, we shall proceed to consider what we can ourselves easily find out about the air, and from this basis of knowledge we shall follow what has been still further discovered by prolonged investigation in all parts of the world.

9. At the outset it may be well to group together in due order the different subjects which, coming within the scope of Physical Geography, will have to be attended to in these Lessons.

10. First of all we shall consider what the Earth is, as a heavenly body, how it is related to other heavenly bodies, and specially to the Sun, as the source of light and heat.

11. Secondly, looking at the Earth in itself, we find it wrapped round with an outer envelope of Air, which will next deserve our attention. What this envelope consists of, and the part it takes in the phenomena of the Earth's surface, furnish materials for much interesting inquiry.

12. Thirdly, beneath the surrounding shell of air, but covering the greater part of the surface of the solid globe, lies that vast expanse of water known as the Ocean. We shall follow its tides and currents, and trace the vapour which, ascending from its surface, is carried through the air until it falls as rain and snow upon the land, whence it is borne by rivers back again into the ocean. The wonderful beauty and high importance of this circulation will claim our attention.

13. Fourthly, the solid Land will be considered, with

its continents and islands, its mountains and valleys, its earthquakes and volcanos. The evidences of continual change in the surface of the land will lead us back to the action of the air and of water, and we shall find how largely the forms even of "the everlasting hills" have been determined by that action.

14. Fifthly, we shall inquire what Climate is, how different kinds of climate are distributed over the globe, and whether any causes can be assigned to account for such differences; this will lead us to note that as there is a geographical distribution of climates, so likewise is there one of plants and animals. Each great region of the earth's surface which has a peculiar climate of its own, has also its own distinguishing assemblage of peculiar plants and animals. Even in the way in which the races of man are grouped over the earth we shall find evidence of the same connection between the distribution of climates and of life. Thus the geographical distribution of Life over the earth's surface will form the concluding part of these Lessons.

CHAPTER I.

THE EARTH AS A PLANET.

LESSON I.—THE EARTH'S FORM.

1. What then is this Earth on which we live? The answer to this question must be chiefly given by astronomy, but there are some familiar features in the earth itself which, if rightly considered, help to make the answer more vividly realized. Evidently there is much advantage in obtaining a clear idea of what the earth is as a whole, before we begin to trace the events which take place upon its surface.

2. In the first place, then, the earth, like the sun and moon, is a globe. This may be proved in various ways.

(1) If we set sail from England and steer westward without turning back, we shall eventually find ourselves in England again. This is called **Circumnavigating the world**. It would not be practicable unless the earth were really a globe.

(2) Standing by the margin of the sea, and looking seawards, we observe that a ship which is moving away from us, by and by seems to sink into or under the sea. (Fig 1) The hull first disappears, and then by degrees the sails. So, on the other hand, when a vessel approaches us, we can first, with a telescope, make out the tops of the sails and masts just peeping above the distant surface of the ocean. Little by little the sails rise as it were out of

the water, until at last the hull and the whole vessel come into sight. This could not happen unless the surface of the sea instead of being flat were really curved, that is, part of the curved surface of a globe.

(3) If the surface of the earth were flat, as men once supposed it to be, the sun would rise at the same time at all places on the earth. But this is not the case; sunrise takes place later or earlier as we travel west or east. Again, on ascending a mountain we should see the whole of the earth's surface if that surface were really a plain.



FIG. 1.—Curvature of the Earth's Surface, as shown by ships at sea.

But the extent to which we see depends on the height to which we climb. This shows that the earth must be a globe.

(4) When the earth comes directly between the sun and moon so as to cut off the light of the former from the latter, the moon is said to be eclipsed. Now, if we watch the earth's shadow as it creeps over the moon's surface, we see it to be circular, and hence we know that the real form of our planet must be that of a globe.

3. Were it possible for us to quit the earth and look back upon it from a distance of a few millions of miles it would appear as a large bright moon, hanging in space

just as our moon does, with the stars twinkling on all sides of it. Its surface would shine like that of the moon, with the light falling on it from the sun, and would be seen to be marked by irregular bright and dark strips and patches. There would be two specially bright parts on opposite sides of the earth's body, marking the great



FIG. 2.—The Earth and Moon as seen from space.

regions of snow and ice at the north and south poles. Between these two areas, certain bright irregular bands would indicate the position of the dry land, while the intervening darker portions would show the extent of the seas (Fig. 2.). If we could transport ourselves still

further away ; if, for instance, we could get as far off as the sun, that is to say more than ninety-one millions of miles from where we are living at this moment, our earth would be seen shining merely as a bright star. And were it possible for us to reach one of the nearest fixed stars, our globe would be no longer visible, while the sun itself if seen at all would appear merely as a twinkling star.

4. But the earth is not a perfectly regular sphere. In particular it is somewhat flattened on two opposite sides, so as to resemble an orange in shape. How this is known and measured will be seen from Lesson IV. A line from the middle of one of these flattened sides passing through the centre of the planet to the middle of the opposite flattened side, is called the *Axis*, and the points at which it reaches the surface are known respectively as the *North Pole*, and *South Pole*. A line drawn midway between the poles round the bulging part of the earth is termed the *Equator*.

LESSON II.—THE EARTH'S MOTIONS.

1. So long as men in early times looked upon the earth as a great plain placed in the centre of the universe, with sun, moon, and stars moving over and under it every day and night, they could not for a moment imagine that the Earth itself is moving. But the truth in this matter has now been so long known, that we have no difficulty in seeing how the apparent movements of the sun and stars are explained by the real movements of the earth. Of these movements the two chief are called *Rotation* and *Revolution*.

2. *Rotation*.—The most obvious movement of our earth is its rotation on its axis, to which the succession of day and night is due. A complete rotation is performed

in about twenty-four hours. During that time each part of the earth's surface is alternately presented to, and turned away from, the sun. As the sun rises in the east to give us day, and appears to travel westward across the sky, the real motion of the earth must be in the reverse direction, or from west to east.

3. It is evident that as the earth rotates, the rate of motion and the distance travelled will vary for different parts of the surface according to their nearness to or distance from the axis. When a cart-wheel, for example, is made to spin round on its axle, a mark made on its rim must move faster and describe a much larger circle than one made near the axle. Places situated on the earth's equator must have the maximum velocity and travel over the greatest distance, while at either pole the velocity must be *nil*. The circle described by a place on the equator is indeed the whole circumference of the globe. So that when the number of miles in the circumference is known and then divided by the number of hours required for a complete rotation, we obtain in figures the rate of motion. At the equator this rate is found to be upwards of 500 yards in a second; it gradually diminishes to the poles.

4. Now if the earth is rotating with such speed, why are we not jerked off its surface? When a stone is thrown into the air, why does it not immediately rush away into space instead of falling back again to the ground? Because what is called the earth's attraction is far stronger than this tendency to fly off. Everything in and on the surface is pulled down towards the centre of the earth by this attracting force, which is called gravity. But it will be seen that the influence of rotation is nevertheless unmistakably marked upon the great atmospheric currents, since these have their direction modified by it (Lesson XI. art. 13).



Fig. 2

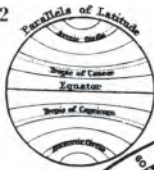


Fig. 5

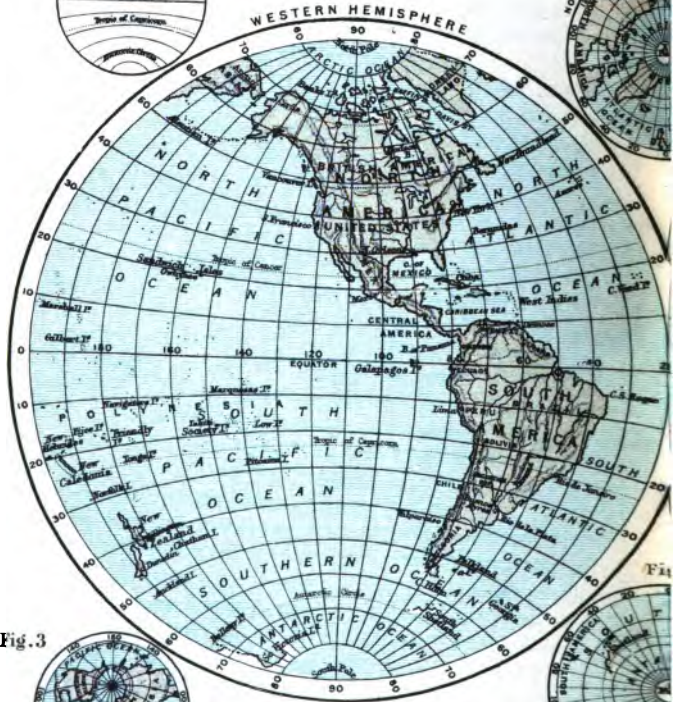


Fig. 3



Fig.1

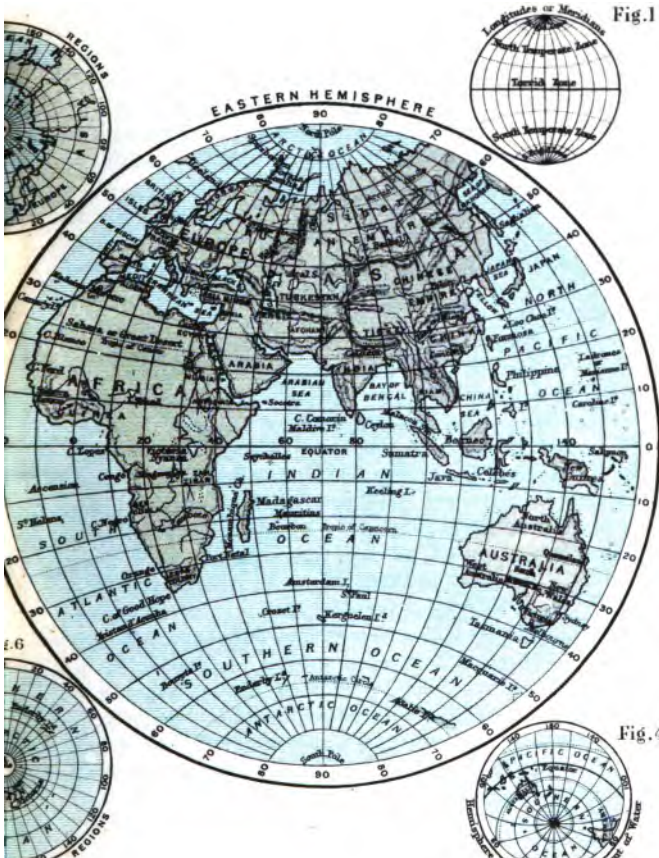
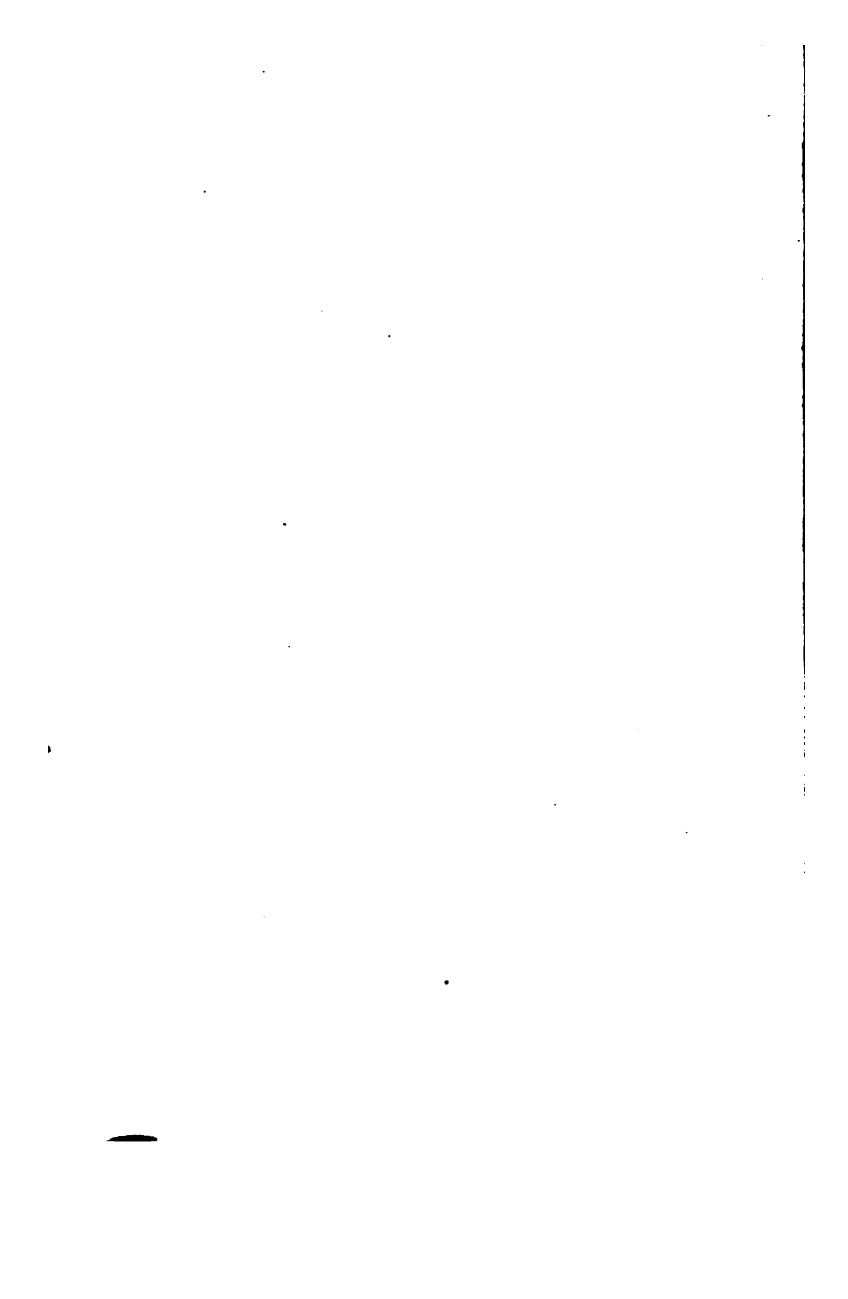


Fig.4

Inburgh and London



5. Revolution.—The earth revolves round the sun and takes about 365 days to perform the circuit. What we call a year is simply the length of time taken by the earth to make one complete revolution round the sun. The path which it follows, called its orbit, has been ascertained not to be a perfect circle but an ellipse, so that our globe is at one part of its course nearer to the sun than at another, its mean distance being computed at 91,328,600 miles. From this number and from the time taken for a single revolution it is easy to find the average rate at which our globe must rush along its orbit. That rate is about 68,000 miles in an hour.

6. There are certain circumstances connected with the movement of revolution which serve to explain why the days and nights throughout the year are not all of the same length, and why instead of perpetual summer or winter there is the regular succession of the Seasons. If the axis of the earth were perpendicular to the plane of the orbit, that is, if it moved along in a perfectly upright position, there would be equal day and night all the year round all over the globe. But it is really inclined to the path of the planet round the sun at an angle of about $23\frac{1}{2}^{\circ}$ and remains always parallel to itself, that is, always points to the same star. In summer the North Pole is turned towards the sun, in winter it is turned away from it. Hence while midway between the two poles the day and night remain each of twelve hours' duration, they depart more and more from this uniformity towards the poles, until there we have a day lasting for one half of the year, and a night lasting for the other half.

7. About the 22nd of March, and again about the 22nd of September, the earth is so placed in its orbit that the sun is exactly vertical over the equator. The line between the sun-lit half and the dark half passes through the poles. At these times, therefore, day and night are

each of twelve hours' duration over the whole globe ; and these points in the year are accordingly called Equinoxes, that is, "equal nights." After the March equinox, the northern parts of the globe, as the earth moves round in its orbit, come more and more into the sunlight ;

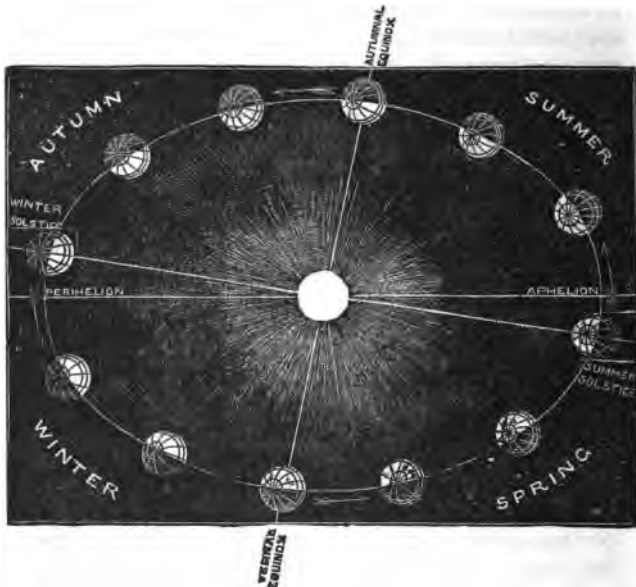


FIG. 3.—The Earth's path round the Sun.

in other words, the days get longer and longer until, at the North Pole and for a certain space all round it within what is called the Arctic Circle (see Plate I.), the sun does not set at midsummer at all, and there is continuous daylight. The further north we go the longer are the days in June, so that a summer-day in England differs very

much in length from one in India, and one in the north of Scandinavia from one in England. On the other hand, when the earth has got round to the September equinox, the North Pole is beginning to point away from the sun, and the days in the northern part of the earth begin to get shorter, till in mid-winter, at the pole and over the ground within the Arctic Circle the sun does not rise, and there is continued night. The further north we go in December, therefore, the longer are the nights; so that while a summer-day in London is not so long as a summer-day in the north of Scotland, a winter-day is considerably longer.

8. It is evident that when the North Pole is enjoying perpetual daylight, the South Pole must be in continuous night. It is only at the two equinoxes that their day and night are equal. As the days about the North Pole shorten, those at the opposite pole lengthen, and then the reverse takes place, and so on from year to year.

9. The alternation of the Seasons depends in like manner upon the inclination of the earth's axis in its yearly orbit. At one part of the year people in Europe and North America see the sun to be comparatively low in the sky: his rays are then but feeble; the climate is cold, and then comes the familiar sight of the frost and snow of *winter*. Six months later the sun appears much higher in the sky; that is, more directly overhead: his rays are then warm—even scorching, and we find ourselves in the midst of *summer*.

10. In reality, however, it is not the sun which has changed place, but our planet, which has reached different parts of its orbit, and consequently has presented itself in different positions towards the sun. In summer the days in the Northern Hemisphere (as the half of the globe between the equator and the North Pole is called) are long because the North Pole is turned towards the

sun. That hemisphere is consequently warmed. The rays of the sun come more directly down upon it; the long days allow much more heat to be received from the sun, and the short nights permit much less heat to be given off from the earth than in winter. While this state of things is going on in the Northern Hemisphere and brings summer there, precisely the opposite effects are taking place in the Southern Hemisphere. There, in proportion as the days lengthen and the warmth increases in the North, the days and the heat decrease, until at the midsummer of the Northern half of the globe, the people at the Antipodes, that is, on the opposite side, or Southern Hemisphere, are in their midwinter. Christmas, which in Europe and North America is always associated with the frost and cold of midwinter, happens at the midsummer of countries in the Southern Hemisphere, like Australia and New Zealand.

11. It is evident that the contrast between summer and winter must be most marked in the region round each pole, where indeed the year may be looked upon as consisting of only two seasons, one of daylight and summer, the other of night and winter. In the equatorial belt, too, that is on either side of the equator as far as the lines called the tropics, there is often a striking difference between the seasons according to the position of the sun in the sky. At each equinox the sun appears vertical over the equator. From March to June he seems to be travelling northward until about the 22nd of the latter month, when he shines vertically over a line called the Tropic of Cancer $23\frac{1}{2}^{\circ}$ north of the equator. He then turns southward, is again vertical above the equator at the September equinox, after which he travels southward until about the 22nd of December, when he shines directly overhead along the line known as the Tropic of Capricorn $23\frac{1}{2}^{\circ}$ south of the equator. The word tropic

means 'a turning' and is given because the sun after appearing to travel away from the equator begins to turn at these two limits on each side. The sun seems to be constantly travelling to and fro across the sky between the two lines of the tropics. This apparent movement of the sun is of course due to the real movement of the earth, and the limit within which the sun travels north and south of the equator is fixed by the angle of inclination of the earth's axis ($23\frac{1}{2}^{\circ}$).

12. In the belt between the tropics the sun appears vertical in the sky twice in each year. With each of these periods comes a rainy season, and between them a dry and hot season. So that if no other influences came into operation the equatorial regions would have two wet and two dry periods in each year. Owing, however, to the way in which the wind currents are interfered with by the masses of high land, it is in such regions as that of the broad Pacific Ocean that this arrangement of the seasons is most completely carried out.

LESSON III.—THE EARTH AND THE SUN.

1. Long before Astronomy or any science had arisen, the early races of man, seeing how greatly the earth is dependent on the sun, worshipped that luminary as the great parent of all the light, and heat, and life of the world. And surely of all the forms of idolatry none is so natural as this. Before they knew what the sun really is, men might well be content to prostrate themselves before it as a great and good Being, who when he rose at dawn, lighted up and warmed the world, and when he sank at dusk, left it in darkness and chill.

2. But how or why is it that our light and heat come from the sun? We know that the earth revolves round


the sun, but why should this be? Is it possible to know anything as to the real relation of the earth to the great luminary?

3. Let us try to find an answer to these questions by putting down some of the facts which have been discovered about the earth and the sun. In doing this we shall see how step by step the evidence grows, that the dependence of the earth upon the sun is so close because in reality the two bodies originally formed part of one great rotating mass of matter.

4. To begin with the earth. When a very deep bore is made into the ground in search of water, as is done round London and Paris, sometimes to depths of more than a quarter of a mile, it is found that the water which rises from so great a distance is warm. In like manner the air of deep mines is often so hot that the miners have to work with hardly any clothes on them. In many parts of the world springs of hot and even boiling water rise to the surface and have been doing so for hundreds of years without becoming perceptibly cooler. Again in all quarters of the globe volcanoes, or burning mountains, are found from which steam, hot gases, and molten rock, are from time to time given out, often with prodigious force and in enormous quantity (see Lesson XXII.).

5. From all these facts which have been observed throughout the world, in the coldest countries as well as in the warmest, the inference has naturally been drawn, that the inside of the earth must be intensely hot. And since the depth of even the deepest bores or mines is not in comparison to the whole mass of the earth so much as the thickness of varnish is to an ordinary school-globe, it would seem that the earth must be on the whole, a hot globe with a cool outer skin or crust.

6. Since its surface is cold the earth cannot give out any light of its own. And yet it undoubtedly shines in



the sky, as the moon does, because like that orb it receives and reflects light from the sun.

7. Passing from the earth to the nearest of the heavenly bodies—the moon, which as the earth's satellite or attendant revolves round our globe while the latter is revolving round the sun—we find some wonderful evidence that the earth is not the only orb which shines with light borrowed from the sun, and though cool on the surface bears traces of the influence of internal heat.

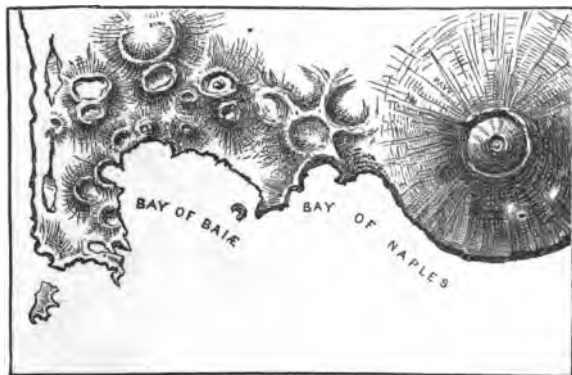


FIG. 4.—Plan of Volcanic Hills and Craters in the Bay of Naples.

8. The nature of this evidence may be followed by comparing the two drawings, Figs. 4 and 5. In Fig. 4, a sketch or plan is given of part of the neighbourhood of Naples. The numerous detached conical hills are volcanoes, either now active as Vesuvius is, or cold and silent like the group of cones to the west of Naples. The circular pits placed on the tops of these hills are the "craters" or vents from which dust, cinders, steam, and

molten rock, have at different times been copiously discharged. In Lesson XXII. some account will be given of volcanoes, and it will there be shown how the streams of molten rock now and then run down the outside of the cone, while vast clouds of steam and hot dust, cast up with great violence, bear further witness to the intense heat of the interior.

●. Now with this little bit of the earth's surface compare the drawing in Fig. 5, which shows a portion of the

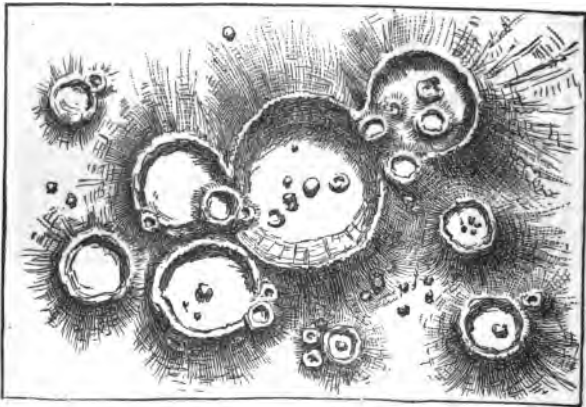


FIG. 5.—A Part of the Surface of the Moon showing Volcanic Craters.

surface of the moon. We observe in it a series of large "craters," with steep walls in the inside, and numerous smaller ones, sometimes in close-set rows. Several of the larger craters overlap each other, and in some cases their interior is crowded with the smaller ones, or the latter have broken through the ridges separating the great basins. So close is the likeness of this scene to such parts of the earth's surface as that shown in Fig. 4, that there seems

no reason why we should not call these conical elevations on the moon volcanoes, and look on their well marked "craters" as really the openings whence molten rock and other hot materials have been thrown out, as in terrestrial volcanoes. So far as can be seen they abound over most of the moon's surface. Indeed the lunar volcanoes have been measured by astronomers, and found to be far more numerous, and greatly more gigantic than the terrestrial ones. We may fairly infer that they mark a much more intense kind of volcanic action than anything we know of on the earth. So that though the chill surface of the moon now gives out no light of its own, but only throws back to us that which strikes upon it from the sun, the interior of that globe must once have been intensely hot, with a hardening crust through which the internal molten liquid was poured by those hundreds of cauldron-like craters and cracks.

10. No telescope has yet detected any actual volcanic eruption going on in the moon. It would seem, indeed, that the once prodigious volcanic action there has spent itself, and that the inside of the moon, though perhaps still warmer than the outside, has not heat enough left for further outbursts. Evidently the moon has cooled very greatly since the time when its craters were active volcanoes.

11. The earth must be cooling likewise. The mere difference of temperature between its surface and interior indicates this fact. The outer shell, or crust, would not be kept colder than the inner mass unless there were a continual loss of heat from the earth into space. Heat is steadily passing outwards through the crust and surface, but it does not make the ground on which we walk sensibly warmer, because as fast as it reaches the surface it is radiated into space.

12. Thus we see that the earth cannot have been

always as it is now. A million of years ago it must have had considerably more heat than it possesses to-day. A hundred millions of years ago it was, perhaps, in the condition of a globe of molten matter, with no land or sea, and of course with no life of any kind upon its surface. The present flattening of its form at the poles is just the kind of shape which such a liquid globe would necessarily assume under the influence of rotation, and probably the permanent flattening dates from that time. Earlier still the earth might have existed merely in the state of vapour.

13. That this has really been its history has been thought by many philosophers to be in the highest degree probable, when, turning from the evidence which the earth itself presents, attention is given to that furnished by the sun. That the sun is hot has long been familiar knowledge, but only in recent years has any adequate notion been formed of how intense its heat actually is. One fundamental difference between the sun and the earth and moon is that its light is given out by itself, while theirs is only borrowed sun-light. When sun-light is examined by means of the spectroscope¹, it is found to indicate a temperature in the sun so enormously high that nothing can exist there except in the form of gas or vapour. The light and heat which we receive from the sun proceed from glowing vapours, among which have been detected those of some of the metals found on the earth only as solid bodies very difficult to melt. Probably most or all of the simple substances composing the earth exist also in the sun, but in the form of vapour. Were our globe thrown into the sun it would immediately be dissipated into the same glowing vapour.

14. Observation of the sun's surface has shown that

¹ For an account of the spectroscope and spectrum analysis see Roscoe's *Spectrum Analysis*, or Lockyer's *The Spectroscope and its Applications*.

spots which appear there are carried steadily round from west to east. This points to a movement of rotation, which, however, is slower than that of the earth. The sun, or at least its outer luminous envelope which we see, takes about twenty-five of our days to turn round once on its axis; but the movement is in the same direction as that of the earth.

15. Besides the earth and its attendant moon, a small number of other heavenly bodies has been found to revolve round the sun. If we note carefully the positions of the stars in the sky we find that, though the whole sky seems to travel slowly westward, each star keeps the same place with reference to the other stars. But at certain seasons we may observe some which look at first like stars, but which, when watched more attentively, may be seen to shift their place and to travel across the other stars. These were called by the ancients, planets, that is, wanderers. They are now known to be really revolving round our sun at different distances and in various periods of time. So far as can be judged from what the telescope reveals, they closely resemble the earth in many points. They rotate on their axes. Some of them have a group of attendant moons. In some there are indications of an atmosphere with clouds and air-currents, and in one of them, called Mars, there appears to be a cap of snow and ice at each pole, as in the earth. Some are much larger, others much smaller than the earth; some are nearer to the sun than we, others greatly further away. This whole series of planets (including, of course, the earth), which revolves round the sun as its centre, is known as the Solar System.

16. Now let us sum up all these facts and see how they bear upon the present condition and probable history of our own globe. In the centre of the solar system stands the sun—an enormous globe of incandescent gas and

vapour, rotating on its axis, and sending forth heat and light far and wide through space. Round this central luminary, and depending on it for light and heat, a number of planets revolve in the same general plane, sometimes with minor planets or satellites revolving round them as the moon does round the earth. The planets show the same rotatory movement as the sun. The earth is one of the planets. Its present condition points to its once having been hotter than it now is, and possibly to its existence in a liquid or even gaseous state.

17. The nebular theory, which has been proposed to connect and explain these facts, maintains that at first the whole solar system existed only in the state of vapour, like one of the faint cloud-like nebulae disclosed by the telescope among the stars; that this nebula gradually condensed and threw off from its hot mass successive portions, which, by subsequent condensation and cooling, became planets, and that the present sun is the remaining incandescent, but still slowly condensing and cooling nucleus of the whole, round which the various detached portions continue to revolve.

18. If the theory thus briefly sketched be a true one, and the discoveries of modern science go far to confirm it, it gives an intelligible reason for the intimate way in which the earth and the other planets are related to the sun. It furnishes a satisfactory explanation, too, of the fact that the inside of the earth is still hot, though slowly cooling. That internal heat, manifested, as we have seen, in every deep bore and mine, as well as in hot springs and volcanoes, appears thus as the residue of the original heat of the great nebula out of which the planets and the sun have been condensed.

LESSON IV.—MEASUREMENT AND MAPPING OF THE EARTH'S SURFACE.

1: It would not have been possible for geography to make much progress without some means of accurately ascertaining the positions of places on the earth's surface. On a small scale we can measure with great exactness by means of carefully-constructed chains or rods, how far one place is distant from another. But evidently this laborious method could have been of little use in laying down the great features of the earth's surface, such as continents, islands, oceans, and the rest, or in discovering the actual dimensions of the planet itself. It was needful to obtain some other method which could be easily used, and would give the means of determining with precision the position of every part of the surface of the globe. Such a method is supplied by observing the position of the sun and different stars.

2. If we note the height of the sun in the sky at noon, we shall find that an hour after that time he seems to have moved a certain distance westward. In the course of another hour he will have traversed another similar space, and so on hour after hour till nightfall. Next morning, if we again watch his progress, we detect a similar advance hour by hour, until at noon he once more stands in the same position as at noon of the day before. In this interval the earth has made one entire rotation; we have been carried round a complete circle.

3. Every circle is divided into 360 equal parts, or degrees, and as we travelled round our circle in twenty-four hours, we must have passed over fifteen degrees every hour. Suppose that some place, say the Greenwich Observatory, is fixed on from which to count the degrees. It is clear that all places lying to the east

of Greenwich have their noon earlier, and all places to the west have it later than at Greenwich itself. As the sun takes an hour to travel 15° of the circle, any place which has its noon exactly one hour after noon at Greenwich must lie 15° to the west.

4. To determine, therefore, how far we are to the east or west of Greenwich, we may try to find out the difference between Greenwich time and the local time at the place where we may be. This has been done for short distances by flashing mirrors, exploding gunpowder, or in any other way communicating an instantaneous signal from one point to another. The most efficacious method is by electric telegraph. But for long journeys, especially sea-voyages, where no such communication is possible, carefully-constructed clocks called chronometers are used, which show Greenwich time. By comparing the local time, as fixed by taking the sun's position, with the time shown in the chronometer, it is easily seen how far a place lies to the east or west of Greenwich.

5. But as even the most accurately finished clock is apt to gain or lose, there is still another and more reliable, though less convenient kind of observation—that of determining the position in the sky of the moon or some of the planets. The places which these bodies will have with reference to each other and to the fixed stars at any moment is calculated for a long time beforehand at the Greenwich Observatory, and tables showing these positions are printed. By referring to these tables the traveller finds the precise instant of Greenwich time, for each position of the heavenly bodies; and the difference between that time and the time he observes at the place where he is shows him whether and how far he is to the east or west of Greenwich. This is called "finding the longitude" of a place.

6. If we mark each of the 360 degrees into which the

circumference of the globe is divided by a line passing from pole to pole, each line will be found to run due north and south, and a little reflection will show us that every place on the earth's surface lying along the course of the same line must have its noon at the same moment. These lines are termed meridians (Plate I., Fig. 1). Of course we may fix on any point from which to begin to count them. In England and in English-speaking countries they are reckoned from Greenwich Observatory. In France they are counted from the Observatory at Paris. But this is immaterial, though it would be a great convenience if all countries would agree to count from the same starting-point. The meridian line passing through Greenwich is marked by us as the zero or 0° . When it is twelve o'clock at Greenwich it is precisely the same time at every place through which the meridian of Greenwich passes. All the meridians lying west are West Longitude, those to the east are East Longitude, so that the two series meet on the other side of the globe, exactly opposite to Greenwich, at 180° , or half the number (360) into which the whole circle of the globe is divided. Having then fixed on our meridian as the starting-point, it is not difficult to see how the relative distances of places east or west of that meridian are determined and expressed. We say that a place which lies five degrees east or west from Greenwich is in 5° east or west longitude. Each degree is further divided into sixty minutes, and each minute into sixty seconds. Paris is situated two degrees, twenty minutes, and nine seconds east from Greenwich, which is abbreviated thus : $2^{\circ} 20' 9''$ E.

7. One degree of longitude is equal to four minutes of time. When we travel eastwards our watches seem to be going slower at the rate of four minutes for every such degree, because we are getting into longitudes where the time is earlier than at home, while, on the other hand

when we travel westwards, our watches seem to be going too fast by the same intervals.

8. But this effect of difference of longitude is brought out in by far the most striking manner in the sending of messages by electric telegraph. Though two places may be thousands of miles apart, a word sent from one is almost instantaneously received at the other. When a clerk in London telegraphs at noon to Calcutta (which lies in $88^{\circ} 30'$ E. longitude) his words, though they go with the speed of lightning, arrive about six o'clock in the evening by the clock in India; or should he send a telegram at the same hour to New York, which is situated in 74° W. it will find the clocks there pointing to seven o'clock in the morning.

9. But the determination of the longitude of a place would not be enough. We must be able to tell whereabouts that place lies on its meridian. To do this is what is termed "finding the Latitude." The first thing to strike us in this problem is the fact that we do not require to fix on any arbitrary point from which to begin our reckoning. The axis of the earth gives us two definite points at each pole, and midway between these lies the line of the equator. To determine the latitude of a place, therefore, we have to ascertain how far it lies to the north or south of the equator. Here again we must have recourse to the heavenly bodies.

10. If the axis of the earth were prolonged from the North Pole it would reach a point in the heavens close to the pole-star, called the Celestial Pole, round which in the northern hemisphere the stars, by reason of the earth's rotation, seem to revolve. So that though we cannot measure from the North Pole itself, we can determine our distance from it by observing how much the point of the heavens directly above us (that is, the zenith) is removed from the pole-star or from the celestial pole.

In the same way if the line of the equator were prolonged it would traverse the heavens as a great circle. Knowing exactly its position in the sky or the distance of any heavenly body from it, we observe how far our zenith is removed from it, and thereby learn our distance from the terrestrial equator.

11. Now the distance from either pole to the equator is exactly a quarter of a circle, or 90° . If we mark off the degrees by a series of lines on the surface of the earth, these circles will run round the globe parallel to the equator and to each other, but diminish in diameter as they advance to the pole. Such lines are called parallels of latitude (Plate I., Fig. 2). They are counted from the equator, which is the zero, or 0° ; and each degree is divided into minutes and seconds like the degrees of longitude.

12. If we find by observation that a place is situated fifteen degrees north and another twenty degrees south of the equator, we say that the one is in Latitude 15° North, and the other is 20° South. Each pole is hence in latitude 90° . Since the figures expressing the degrees increase in amount as they recede from the equator, it has become common to speak of "low latitudes," that is, regions or places lying towards the equator, and "high latitudes," that is, tracts situated near either pole.

13. In the degrees of latitude and longitude, then, we have two sets of lines, one running north and south, the other running east and west, by which the surface of the globe is supposed to be traversed as by a regular network. By means of these lines it is clear that we can fix exactly the relative positions of places on the earth's surface. But we must still determine what is the absolute length of one of those degrees in miles before we can ascertain the true distances of places and the real area which any country, or continent, or sea covers. When this is done

we have the materials for estimating the total bulk of our planet and making a correct survey of its surface.

14. Since the lines expressing the meridians of longitude converge to each pole, their distance from each other must vary according to latitude ; the parallels of latitude necessarily diminish in circumference as they recede from the equator. But were the globe a perfectly symmetrical sphere the length of each meridian from equator to pole should be exactly alike, and each of its ninety degrees should be of uniform length. The accurate measurement, therefore, of the length of one of these degrees should give us the means of easily computing the sum of the whole, and thence of ascertaining the true size of the globe.

15. This measurement of a degree of the meridian, as it is called, has been made with great care in different parts of the world. In India the length of a degree was found to be 362,956 English feet, or rather more than $68\frac{1}{2}$ statute miles ; a degree in Sweden measured 365,744 feet, or, in round numbers, about $69\frac{1}{4}$ miles. It has been found that besides little irregularities which show that the shape of the planet is slightly distorted, there is a progressive increase in the length of the degrees towards the poles as these two measurements in India and Sweden show. This could only take place by a flattening of the globe at the poles.

16. The sum of all the observations gives for the polar diameter of our globe a length of 7,899.17 statute miles, and for the mean equatorial diameter a length of 7,925.65 statute miles. As the difference is about $26\frac{1}{2}$ miles, each pole must be compressed to the extent of $13\frac{1}{4}$ miles.

It is not easy at once to grasp the full value of these figures. An express train travelling at the rate of thirty miles an hour would take about a month to go completely

round the earth at the equator ; and if it could go through the earth between the poles, it would take about eleven days to perform the journey. Upon the surface of a globe of such dimensions, the highest mountains and the deepest oceans are far less in proportion than the roughnesses upon the rind of an orange.

17. Astronomy has taught us not only the size of our own planet but has computed the dimensions of the others, whence we learn the comparative place of the earth in the solar system. The planet Jupiter, for example, is 1,400 times the size of the earth. On the other hand, our planet is seventeen times larger than Mercury, and greatly larger than certain small bodies called Asteroids. Again, the earth is neither nearest to nor furthest removed from the sun ; its mean distance, as already stated, is computed to be rather more than 91 millions of miles. But Mercury, at its greatest distance, is only about $44\frac{1}{2}$ millions of miles away from the sun, while Neptune revolves at the enormous mean distance of 2,862 millions of miles. The sun itself, the great centre of all the movement of the solar system, would contain 1,400,000 globes as large as ours.

18. Having found a means of accurately fixing the position of any place on the earth's surface and of determining the distance of places from each other, men could nevertheless have been able to carry in their minds but a vague notion of the general features of that vast and varied surface had they not devised a way of putting down these features upon paper so as to show their relative positions and shapes. Such a delineation of the whole or any part of the earth's surface is called a Map. If you look at the map of any country or continent you observe it crossed by two sets of lines, one of which, running from the top to the bottom of the paper, mark the

degrees of longitude, while the other set, running from side to side, show the parallels of latitude. It is usual in such maps to represent the ground as it might be supposed to appear could we ascend to a great height in the air and take in the whole area at one view, our heads being turned to the north. Hence the top of the map

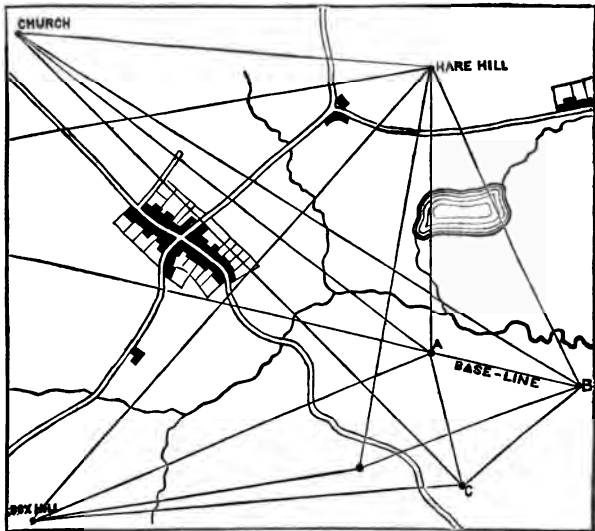


FIG. 6.—Measurement and mapping of a country by means of triangulation.

is north, the right hand being the east side, and the left hand the west.

10. When maps of a country are wanted on a large scale to express the features of the ground in great detail, another process of measurement called triangulation is used. First a base-line, a few miles in length, is

accurately measured with rods or chains as from A to B in the figure. Then from A an observation is taken with an instrument called a theodolite, to the point C, which may be a hill-top or church tower, or any distinct object, and the angle CAB is carefully noted. A similar observation is taken at B to determine the angle CBA. We can now construct our first triangle, and having found by actual measurement the length of its base, we find by simple calculation the length of its two sides, and consequently the exact position of C and its distance from A and B respectively. From such a measured base a system of triangles is observed all over the country, and the true positions of all its main landmarks are fixed without the necessity of further measurement.

LESSON V.—A GENERAL VIEW OF THE EARTH.

1. We proceed now to take a first general view of the parts of the earth with which further acquaintance is to be made in the following lessons. Our progress will probably be found to be more easy and pleasant, if we can carry with us from the outset such a broad but clear notion of the earth as a whole as will enable us to follow the necessary references to parts with which we may still be unfamiliar.

2. Beneath the outer envelope of Air with its winds, clouds, rain, and snow, the surface of the globe presents the two clearly marked, but very irregular divisions of **Sea** and **Land**. As the result of the surveys and observations taken in all parts of the globe, it is ascertained that the sea covers nearly three-fourths, and the land rather more than one-fourth, of the whole surface of the

planet ; or, more exactly, there are 275 parts of water to 100 parts of land.

3. For the sake of aiding our conceptions of such statements as these, globes and maps are prepared, of which Plate I. may serve as an illustration. Each of the two large circles there represents one side of the earth, and shows the manner in which sea and land are grouped. It will be observed, that while the amount of sea greatly exceeds that of land, the difference becomes more marked in the southern half or hemisphere, which is almost all water, while most of the land lies on the north side of the equator. But a school globe may be so placed, as to show nearly the whole of the land at one view, and the greater part of the sea at another. For this purpose, turn the globe so that the south-west of Britain shall be directly in the centre of the globe as you look at it, as shown in Fig. 3, Plate I. Britain is thus shown to lie in the centre of the habitable globe. Now turn your globe so that the point on the other side, exactly opposite to the south-west of Britain, shall be the centre. The islands of New Zealand are then not far from that position, and all round them lies the water hemisphere with comparatively few detached masses of land (Fig. 4, Plate I.).

4. While the great body of the land lies on the north side of the equator, it will be seen from Plate I. that it forms there two well-marked portions separated by two great tracts of sea. The larger of these portions includes Europe, Asia, and Africa, that is, all the regions which have been longest settled by a human population. Hence it is often spoken of as the Old World. The other portion embraces America, and is called the New World. As the Old World lies to the east and the other to the west, they are commonly described as the Eastern Hemisphere and Western Hemisphere.

5. One difference between the distribution of sea and

land cannot fail to arrest attention at the outset. The land is much broken up. Even in the Eastern Hemisphere, where it presents the most compact mass, long arms and inlets of the sea intersect it, and cut off large portions from the rest. The sea, on the other hand, is evidently one continuous whole. Even when penetrating furthest into the land, its most distant arms retain their connection with the main body. A vessel may sail over every sea and into every far recess and inlet, without ever needing to be dragged over an intervening mass of land. But no coach or form of carriage could visit every part of the land, without requiring to be transported across intervening tracts of sea. There are no isolated areas of sea completely encircled by land, answering to the abundant portions of the land which, encircled entirely by sea, receive the name of islands.²

6. Though the sea is one continuous liquid mass, it has been for the sake of convenience in description divided into different areas, termed Oceans. The limits of these are for the most part indicated by the position of the masses of land. Thus the largest, under the name of the Pacific Ocean, fills up the vast area between the western edge of America, and the eastern margin of Asia and Australia. The line of the equator divides it into the North Pacific and the South Pacific Ocean. Another longer but much narrower belt of sea, intervenes between the eastern coasts of America and the western coasts of Europe and Africa; divided in the same way by the equator into the North Atlantic and the South Atlantic Ocean. Owing to the broad mass of land forming Europe and Asia, no corresponding belt of sea can run there from pole to pole. But south of that land lies the wide Indian Ocean, between Africa and Australia. It is usual to call

² Reference will be afterwards made to one or two interesting exceptions to this rule, the Caspian Sea being the chief.

all that part of the sea within the Arctic Circle the Arctic Ocean, and the corresponding tract in the southern hemisphere, the Southern or Antarctic Ocean.

7. Besides these main sub-divisions, there are minor tracts of sea, more or less surrounded by land. Such names as sea, gulf, strait, channel, are given to them according to their size or the shape of the land-outlines by which they are bounded.

8. The upper surface of the sea forms a sharp, well-marked line, which is called the sea-level. It serves as the line from which the heights of the land and the depths of the oceans are measured. Slight differences of level have been detected between different oceans or seas, as, for instance, between the Atlantic and Pacific, on the two sides of the narrow part of America, and between the Mediterranean and the Red Sea. But such differences never amount to more than a few feet, so that for most practical purposes we may assume the sea-level to be uniform.

9. Passing next to the land, we observe from the map, that while massed on the whole in the northern half of the globe, it is far from forming there one solid block ; on the contrary, it is intersected by branches from the sea, so as to be easily grouped into a few great sub-divisions. These are termed Continents. Strictly speaking, there are only two continents, the Old World and the New World (Art. 5). More usually, however, they are grouped in three pairs, the first pair consisting of North and South America, the second of Europe and Africa, and the third of Asia and Australia.

10. Now in considering the general arrangement of the land on the globe, we soon see that one of the most marked general features of the continents is their tendency to get massed together towards the north, and to taper away towards the south to about

the parallel of 45° S. Look, for example, at the way in which the mass of Africa dwindles away southward to the Cape of Good Hope, and how the huge bulk of the Asiatic Continent and its prolongation in Australia tapers away to the headlands of Tasmania. Still more remarkable is the attenuation of the American Continent and its sharpened termination at Cape Horn.

11. Again, not only is the land much intersected by inlets, such as the great Mediterranean Sea and the Red Sea, but unlike the sea (Art. 5), large parts are completely cut off from the main mass, as in Australia, New Zealand, Japan, Britain, and the host of islands, large and small. Other parts nearly isolated are termed peninsulas. Of these Africa may be taken as the most conspicuous illustration. It is joined only by the little connecting neck or isthmus of Suez to the continent of Asia, so that were that strip of lowland to be cut through or sunk beneath the sea, Africa would become an island. This, indeed, has been in some sense accomplished by man in the making of the Suez Canal.

12. Another feature of the land marks it off in striking contrast from the sea. The surface of the latter, though liable to be roughened by ripples and waves, preserves everywhere the character of one vast level plain. But the surface of the land abounds in unevenness. Some parts, indeed, are flat, but most of it undulates into hills and valleys, and some portions mount up into vast ranges of rugged and pointed mountains.

13. This irregular distribution and inequality of surface give the land a peculiar character and influence in the physical geography of the earth's surface. Without inquiring at present how this influence shows itself, we can readily see that if either in the air or in the sea there is any movement of circulation, the currents, both of air and sea, must be greatly affected by the position

and form of the continents and islands which may lie in their course. A current in the sea moving westward, for instance, across the middle of the Atlantic, will strike against the long ridge of America, and be turned aside either to right or left, or both. A current of air, on the other hand, if it should take its rise from some centre in the ocean region, will be turned aside when it meets a mass of high land, and may be driven to ascend the slopes, discharge its moisture, and flow on at a much greater height, and at a very different temperature (Lesson X., Art. 31). The varying influence of sea and land upon the air, lies, as we shall see, at the foundation of all the changes of weather.

14. It has been already stated (Lesson III., Art. 4) that at many parts of our planet's surface, pipes or funnels exist, which descending deep into the earth, emit from time to time steam, hot vapours, dust, stones, and melted rock, so as in the end to form conical hills or mountains called volcanoes. It is observed that orifices of this kind are apt to occur in long lines, and especially along the back-bones of the continents, and long chains of islands (see Plate IX.). Over many parts of the globe, too, and particularly in those regions where volcanoes abound, the ground is frequently shaken by earthquakes, and sometimes permanently lifted above its previous level. Such appearances as these afford indications, not only of the nature of the earth's interior, but also of the way in which the interior affects the surface. They help us to understand how it is that the land has been ridged up above the general level of the sea.

15. Into these matters we shall enter more fully in later Lessons. Meanwhile, carrying with us this general outline of the several parts of the earth, we may proceed to consider these parts one by one in detail.

CHAPTER II.

AIR.

LESSON VI.—ITS COMPOSITION.

1. Above and around us, to what part soever of the earth's surface we may go, at the top of the highest mountain as well as at the bottom of the deepest mine, we find ourselves surrounded by the invisible ocean of gas and vapour which we call AIR. It must, therefore, wrap the whole planet round as an outer envelope. Considered in this light, it receives the distinctive name of the ATMOSPHERE, that is, the vapour-sphere—the region of clouds, rain, snow, hail, lightning, breezes, and tempests. In the study of the earth as a great habitable globe, this outer encircling ocean of air is the first thing to be considered. What is it? and what purposes does it serve in the general plan of the earth?

2. In early times men regarded the air as one of the four elements out of which the world had been made. It is not so very long since this old notion disappeared. But now it is well known that the air is not an element, but a compound of two elements—viz., the gases called **Nitrogen** and **Oxygen**. It is easy to prove this by burning a piece of another element, phosphorus, in a closed jar. We thereby remove the oxygen,

which unites with the phosphorus to form a compound substance, and the nitrogen is left behind.¹ In various ways chemists have analysed or decomposed air into its component elements, but the result is always the same, viz., that in every hundred parts of ordinary air there are by weight about seventy-nine of nitrogen and twenty-one of oxygen.

3. Air, when carefully tested, is always found to contain something else than nitrogen and oxygen. Solid particles, with various gases and vapours, are invariably present, but always in exceedingly minute, though most irregular, quantities, when compared with the wonderfully constant proportions of the two chief gases. Some of these additional components of air are not less important than the nitrogen and oxygen. That they exist may be easily proved, and some light may thereby be thrown on the nature and uses of the air.

4. The presence of vast numbers of **solid particles** in the air may be shown by letting a beam of sunlight or of any strong artificial light fall through a hole or chink into a dark room. Thousands of minute motes are then seen driving to and fro across the beam as the movements of the air carry them hither and thither. Such particles are always present in the air, though usually too small to be seen unless when, as in the darkened room, they are made visible against surrounding darkness by the light which they reflect from their surfaces when they cross the path of any strong light-rays. They are quite as abundant in the dark parts of the room, though for want of light falling upon them they are not seen there.

5. Could we intercept these dancing motes and examine them with a strong microscope, we should find them to consist chiefly of little specks of dust. But among them there sometimes occur also minute living germs, from

¹ See Roscoe's *Chemistry Primer*, p. 12.

which, when they find a fitting resting-place, lowly forms of plants or animals may spring. Some diseases appear to spread by means of the lodging and growth of these infinitesimal germs in our bodies, for they are so small as to pass with the air into our lungs, and thus to reach our blood.

●. It is difficult to catch these tiny motes from a sun-beam, but rain does this admirably for us. One great office of rain is to wash the air and free it from these impurities. Hence when rain-water is carefully collected, especially in large towns, it is found to contain plenty of these solid particles, which it has brought down with it in its fall through the air. The accompanying drawing, for example, shows what is seen when a small quantity of rain, gathered from an open



FIG. 7.—What is seen after some raindrops collected in a town are evaporated, and the residue is placed below a microscope.

space in a town, is evaporated to dryness, and the residue is placed under a microscope. Abundant particles of dust or soot are mingled with minute crystals of such substances as sulphate of soda and common salt. Hence we learn that, besides the solid particles, there must be floating in the air the vapours or minute particles of various soluble substances which are caught up by the rain and carried down with it to the soil. In seizing these impurities and taking them with it to the ground, the rain purifies the air and makes it more healthy, while at

the same time it supplies the soil with substances useful to plants.

7. But far more important than these solid ingredients are three invisible substances, two of them gases, called respectively ozone and carbonic acid gas, the third the vapour of water. After a thunder-storm the air may sometimes be perceived to have a peculiar smell, which however is more distinctly given off from an electric machine. This is **ozone**, which is believed to be oxygen gas in a peculiar and very active condition. It promotes the rapid decomposition of decaying animal or vegetable matter, uniting with the noxious gases, and thus disinfecting and purifying the air. It is most abundant where sea-breezes blow, and least in the air of the crowded parts of towns. The healthiness or unhealthiness of the air seems to depend much on the quantity of ozone, which is estimated by the amount of discoloration produced by the air within a certain time upon a piece of paper prepared with starch and iodide of potassium.

8. Consider next the **carbonic acid gas**. When a piece of coal is set on fire it burns away until nothing but a little ash is left behind. Or when a candle is lighted it continues to burn until the whole is consumed. Now, what has become of the original substance of the coal and the candle? It seems to have been completely lost; yet in truth we have not destroyed one atom of it. We have simply, by burning, changed it into another and invisible form, but it is just as really existent as ever. We cannot put it back into the form which it had in the coal and candle, but we can at least show that it is present in the air.

9. The substance of a piece of coal or of a candle is composed of different elements, one of which is called *carbon*. This element forms one of the main ingredients out of which the substance of all plants and animals is

built up. Our own bodies, for example, are in great part made of it. In burning a bit of coal, therefore (which is made of ancient vegetation compressed and altered into stone), or a candle (which is prepared from animal fat), we set free its carbon, which goes off at once to mix with the air. Some of it escapes in the form of little solid particles of soot, as we may show by holding a plate over the candle flame, when the faint column of dark smoke at once begins to deposit these minute flakes of carbon as a black coating of soot on the cool plate. The black smoke issuing from chimneys is another similar illustration of the way in which solid particles are conveyed into the air.

10. But the largest part of the carbon does not go off in smoke. It is in the act of burning seized by the oxygen of the air, with which it enters into chemical combination, forming the invisible carbonic acid gas. It is, indeed, this very chemical union which constitutes what we call burning or combustion. The moment we prevent the flame from getting access of air, it drops down and soon goes out, because the supply of oxygen is cut off.¹ All ordinary burning substances, therefore, furnish carbonic acid gas to the atmosphere.

11. The amount thus supplied is of course comparatively small, for the quantity of vegetable or animal substance burnt either by man or naturally must be but insignificant, when the whole mass of the atmosphere is considered. An infinitely larger quantity is furnished by living air-breathing animals. In breathing we take air into our lungs, where it reaches our blood. A kind of burning goes on there, for the oxygen of the air unites with the carbon of the blood, carbonic acid is produced and comes away with the exhausted air, which we exhale again before taking the next breath. Just as we put out

¹ For some simple experiments on the nature and production of carbonic acid, see Roscoe's *Chemistry Primer*, pp. 3, 5, 14-20.

the burning of a candle by inverting a glass over it, and closing it from the air, so we put an end to our own lives if we shut ourselves off from the air. When we reflect that every air-breathing animal is continually supplying carbonic acid gas to the atmosphere, we perceive how important this source of supply must be.

12. Living plants in the presence of sunlight have the power of abstracting from the carbonic acid of the air the carbon of which their framework is so largely made.¹ When they die, their decay once more sets loose the carbon, which uniting again with oxygen becomes carbonic acid gas, and is carried by rain into the soil, or taken up by the air. All decaying plants and animals which are freely exposed to the air furnish it with this gas.

13. Lastly, in many parts of the world, particularly in volcanic regions, this same gas is given out in large quantities from the ground. From all these various sources then, the atmosphere is continually replenished with carbonic acid gas, to supply the loss caused by the enormous demands of the vegetable world for carbon.

14. Nevertheless, the quantity of this gas present in the air is very small compared with the volume of the nitrogen and oxygen. It has been found to amount to no more in ordinary pure air than four parts in every ten thousand of air. Yet this small proportion suffices to support all the luxuriant growing vegetation of the earth's surface.

15. By the term **water-vapour**, or **aqueous-vapour**, is meant the invisible steam always present in the air.

¹ The familiar practice of growing cress-seed on moist cloth shows this power of living plants. Kept in the light the seeds soon begin to grow, and yield a crop of cress, the carbon of which the grown plants consist being derived not from the water, nor from the cloth, but from the air. See Roscoe's *Chemistry Primer*, pp. 16-18.

Everyone is familiar with the fact, that when water is heated it passes into vapour, which becomes invisibly dissolved in the air. A vessel of water, for instance, may be placed on a table in the middle of a room, heated by means of a spirit lamp till it boils, and kept boiling till the water is entirely driven off into vapour or evaporated. The air in the room shows no visible change, though it has had all this water-vapour added to it. But it may be easily made to yield back some of the vapour. Let an ice-cold piece of glass, metal, or any other substance be brought into the room. Though perfectly dry before, its surface instantly grows dim and damp. And if it is large and thick enough to require some minutes to get as warm as the air in the room, the dimness or mist on its surface will pass into trickling drops of water. The air of the room is chilled by the cold glass, and gives up some of its moisture. Cold air cannot retain so much dissolved vapour as warm air, so that the capacity of the air for vapour is regulated by its temperature. (See Lesson X.)

16. It is not needful, of course, to boil water in order to get enough of water-vapour in the air of a room to be capable of being caught and shown in this way. In a warm sitting-room, where a few persons are assembled, there is always vapour enough to be made visible on a cold glass. In frosty weather the windows may be found streaming with water inside, which has been taken out of the air by the ice-cold window-panes. Whence came this moisture? It has been for the most part breathed out into the air by the people in the room.

17. Each of us is every moment breathing out water-vapour into the air. As a rule, we do not see it, because the air around us is warm enough to dissolve it at once. But anything which chills our breath will make the vapour visible, such as breathing on a cold piece of glass,

or metal, when a film of mist at once appears on the object, or walking outside on a very cold frosty day, when the vapour of each breath becomes visible as a little cloud of mist in the air.


18. No matter, therefore, how dry the air may appear to be, more or less of this invisible water-vapour is always diffused through it. Every mist or cloud which gathers in the sky—every shower of rain, snow, or hail, which falls to the ground—every little drop of dew which at nightfall gathers upon the leaves, bears witness to its presence.

19. The importance of this ingredient of the atmosphere in the general plan of our world, can hardly be over-estimated. It is to the vapour of the atmosphere that we owe all the water-circulation of the land—rain, springs, brooks, rivers, lakes—on which the very life of plants and animals depends, and without which, as far as we know, the land would become as barren, silent, and lifeless as the surface of the moon. It is, likewise, to the changes in the supply of this same invisible, but ever present substance, that the rise of winds and storms is largely due (Lesson VII.).

20. The quantity of water-vapour in the air varies from day to day, and, indeed, from hour to hour. It is always comparatively small in amount, ranging from about four to about sixteen parts by weight in 1,000 parts of air.

LESSON VII.—THE HEIGHT OF THE AIR.

1. Though no actual upper limit has been found to the atmosphere corresponding at all to the sharply defined surface of the sea, we cannot suppose that the air extends indefinitely outwards from the earth. The atmospheric



envelope clasps the planet firmly, and moves along with it, both in rotation and revolution. Were this not the case, it is plain that the earth's movement through the air would be far more rapid than the most furious hurricane. No loose object could appear on the surface of the globe without being instantly whisked off. But the attraction of the earth retains the atmosphere in its place, so that it is carried with the rest of the planet through space.

2. There must then be some upper limit to the atmosphere. Beyond it lies the ether, which is supposed to fill all space, and through which all the heavenly bodies, and the rays of light proceeding from them are moving. How can it be known how far the atmosphere extends above us?

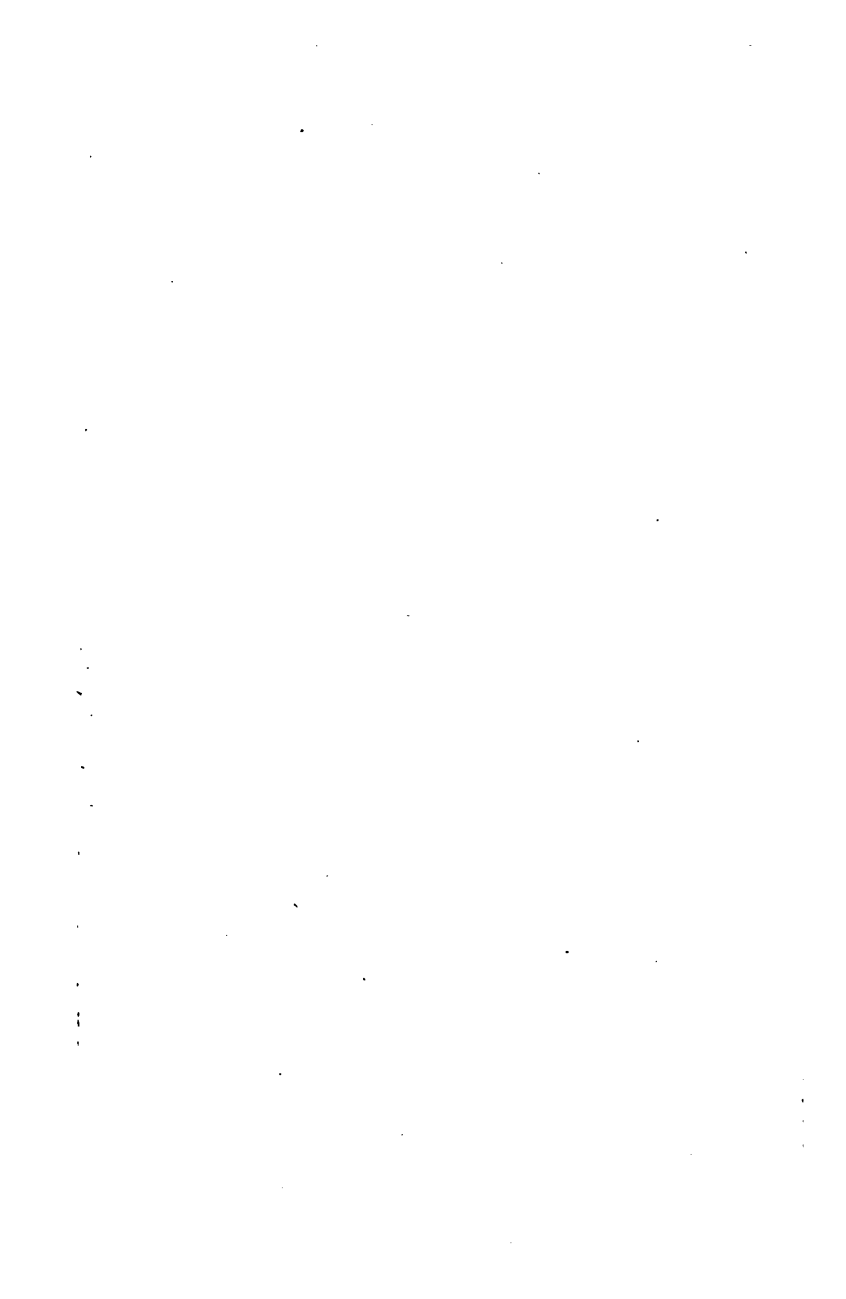
3. There are different ways of trying to answer this question. Let us look for a moment at one of them. Most of us have noticed, that on clear, dark nights, shooting-stars or meteors may be seen sometimes in considerable numbers. They suddenly appear, and after making a train or tail of light, quickly vanish. Sometimes they have been actually heard to explode in the sky, and fragments of them have been picked up. They have been carefully watched by astronomers. By observing their positions and directions of movement from, say, two different stations, the distance between which was known, it has been possible to determine how high they are above us, by a process very similar to that by which distances are estimated on the earth by taking angles to any object from the two ends of a measured base-line. They have thus been found to begin to be visible at from 70 to 100 miles from the earth's surface.

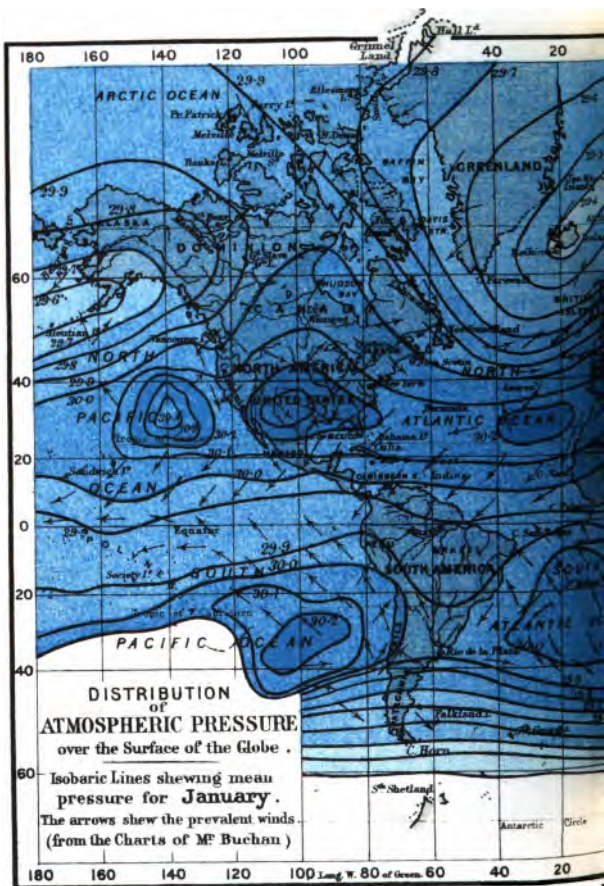
4. In themselves they are little fragments, usually not more than a few ounces or a few pounds in weight; but they revolve round the sun with the velocity of planets or

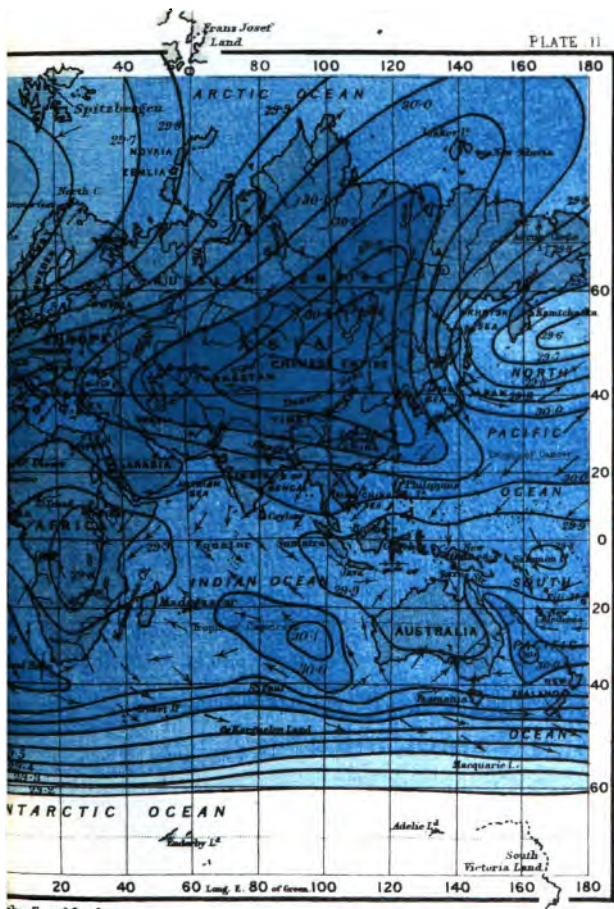
of comets. While still following their usual orbit, they are in themselves cold,¹ opaque bodies. The light which they emit, and by which they become visible, arises from the fact that, drawn out of their course by the attraction of our earth, and rushing into our atmosphere with an enormous velocity, they rapidly get heated by friction against the air, as well as by the heat liberated from the compression of the air in front of them. They soon become white-hot, and in most cases so intense is the heat to which they are raised, that they are dissipated into vapour, which shows as a tail or train of light, gradually fading out of the sky. From the height at which these shooting-stars begin to glow, it is inferred that the atmosphere must extend at least 70 to 100 miles above the general solid surface of the earth, and may even reach further than that.

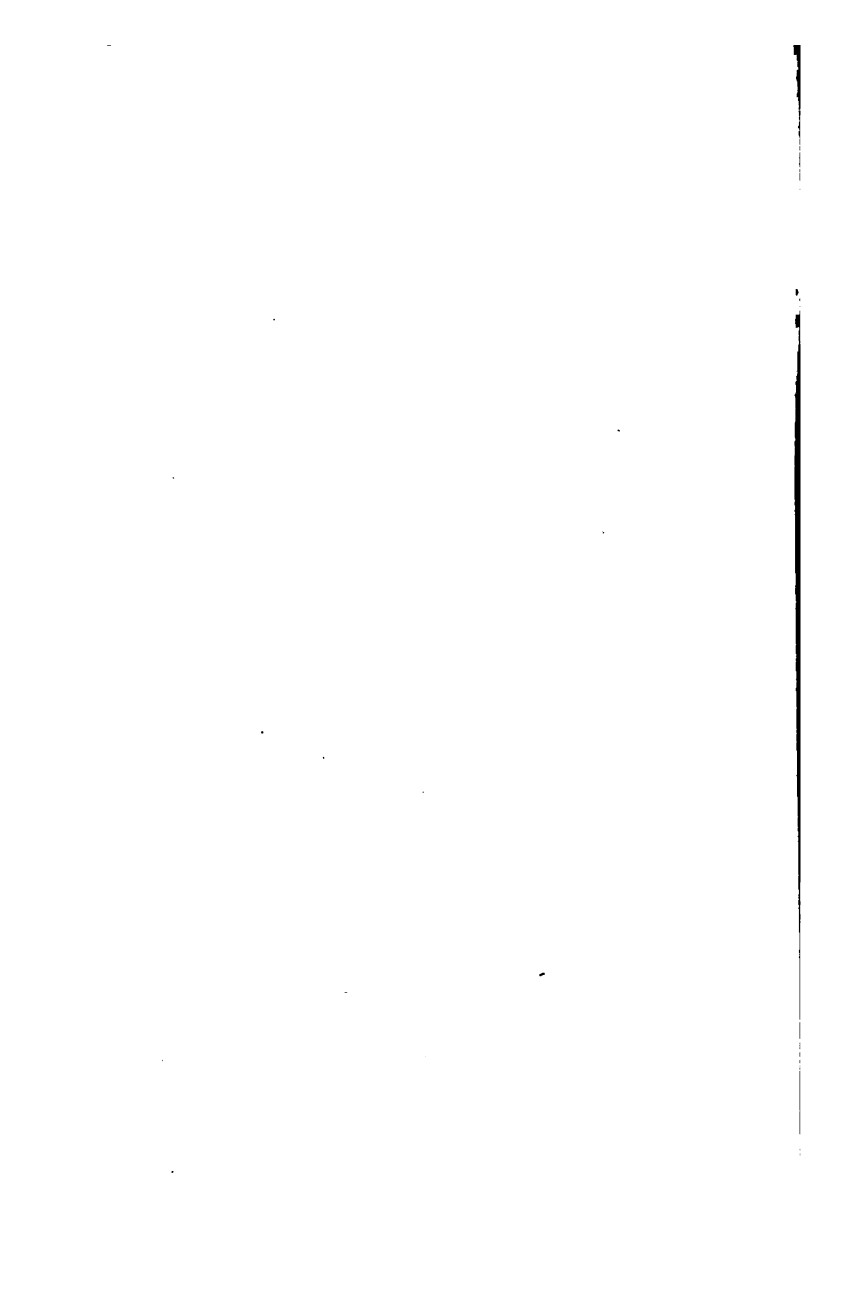
5. But the air at these great heights must be very different in many respects from the air next the earth. We could not breathe it. When, for instance, travellers ascend lofty mountains, it is said that they find an increasing difficulty in breathing as they advance. In the same way persons who have gone up to great heights in balloons have become insensible, and have nearly died from the difference between the air below and that above. The chief difference is in density, the air getting continually lighter or thinner as it recedes from the sea-level. Our bodies cannot endure the difference between the close, heavy air to which we are accustomed next the earth, and the thinner air further up. Breathing becomes impossible at a greater height than six or seven miles. Beyond that distance the air must get less and less dense, until it reaches an extreme attenuation at the furthest outskirts of the atmosphere.

¹ They probably have nearly the temperature of space, which is supposed to be about 271° Fahr. below the freezing-point of water.









LESSON VIII.—THE PRESSURE OF THE AIR.

1. Since the atmosphere becomes less dense as it recedes from the surface of the solid globe, we see that though invisible to us, and so light that we live and move in it without thinking of its existence, air nevertheless presses upon every part of the earth. The lower layers in which we live and move, must be pressed down by the weight of all the mass of air above them. This is what is usually spoken of as atmospheric pressure. It may be due not merely to the weight, but to other properties of the gases and vapours which form the air.

2. To prove its existence, take a little glass phial, and putting it to the mouth, suck out the air as well as possible, taking care to let the tongue press instantly back upon the opening. You feel the tongue driven into the phial, and perhaps even with pain, owing to the pressure of the air outside, and the absence of corresponding pressure from air inside. An effect of this kind is capable of being accurately measured. Accordingly, observation shows that, at the sea-level, this pressure is as much as about $14\frac{1}{2}$ pounds upon every square inch. Every one of us, therefore, bears a weight of 12 or 14 tons of air. Yet we do not feel this pressure, because it is exerted equally on all sides, and because the air within our bodies has the same pressure outwards which the air outside has inwards. If we could withdraw the air from all the cavities and passages in a human body, the weight of the air outside would crush the body in and cause immediate death.

3. Bearing in mind that each part of the atmosphere has to bear the weight of all the air which lies above it, we can understand why, as we ascend and have a decreasing mass of air above us (Lesson VII., Art. 5), we should encounter air of less and less density, that is, why-

the atmospheric pressure should diminish with altitude. If by any means the amount of pressure at the sea-level can be accurately determined, and at the same time, a method can be devised for measuring the rate at which the pressure lessens according to elevation above that level, it will not be difficult to measure the heights of mountains.

4. Now this can really be done by the use of the instrument called the Barometer. The principle of this instrument is that the weight of the atmosphere will balance the weight of a column of any other fluid, the height of that column being determined by the relative weight, or what is called the specific gravity of the fluid employed. A glass tube about thirty-three inches long, closed at one end, is filled with mercury, and then inverted with its open end in a cup of the same fluid metal. The mercury is observed to fall in the tube until, if near the sea-level, it stands at a height of about thirty inches above the surface of that in the cup. The column of mercury thirty inches in height is balanced and kept from descending further by the pressure of the overlying column of the atmosphere upon the surface of the mercury in the cup. The more the atmosphere presses upon the latter the further does the mercury rise in the tube, while on the other hand the less it presses, the more the mercury sinks in the tube.

5. With such an instrument, variations in atmospheric pressure may be detected even when so small and slow that we should never otherwise have been sensible of them. If the height of the mercury in the tube is accurately noted at starting, and the barometer is carried to a higher level, the mercury will be observed to fall because of the diminished pressure, while it will rise again when the instrument is brought back to lower ground. So regular and delicate is this action that the barometer

has often been employed for measuring heights. Were there no other cause for variations in the pressure of the atmosphere than merely elevation, the barometer would be most useful for that purpose, but, as we shall immediately see, would be of no service in the way it is now chiefly employed.

6. The diminution of atmospheric pressure according to height in the air is regular and constant. But besides this the pressure is liable to continual changes at every level, sometimes sudden and great, at other times slow and slight. We are made sensible of these variations when they are accompanied by changes of weather. But they are most accurately measured by the movements of the barometer. If from any cause the pressure should decrease, the mercury will fall, should it increase, the mercury will rise, the rapidity or slowness of the movement in the column of mercury affording us, as it were, a mirror of the amount and the rate of change in the balancing atmospheric column.

7. Suppose by way of illustration that on looking at the barometer some morning we should find the mercury to have fallen a whole inch during the night; the mercury column would thus indicate that during a few hours it had lost a thirtieth part of its whole length, and we should be justified in believing that in some way or other the column of air pressing on the mercury in the cup, had in like manner lost a thirtieth part of its pressure or weight. Some of its upper portions must have flowed over into surrounding regions so as to diminish the pressure to such an extent. No such sudden and great change however, could fail to produce a violent hurricane. The fall of the barometer in almost every case occurs in time to prepare us for the coming storm.

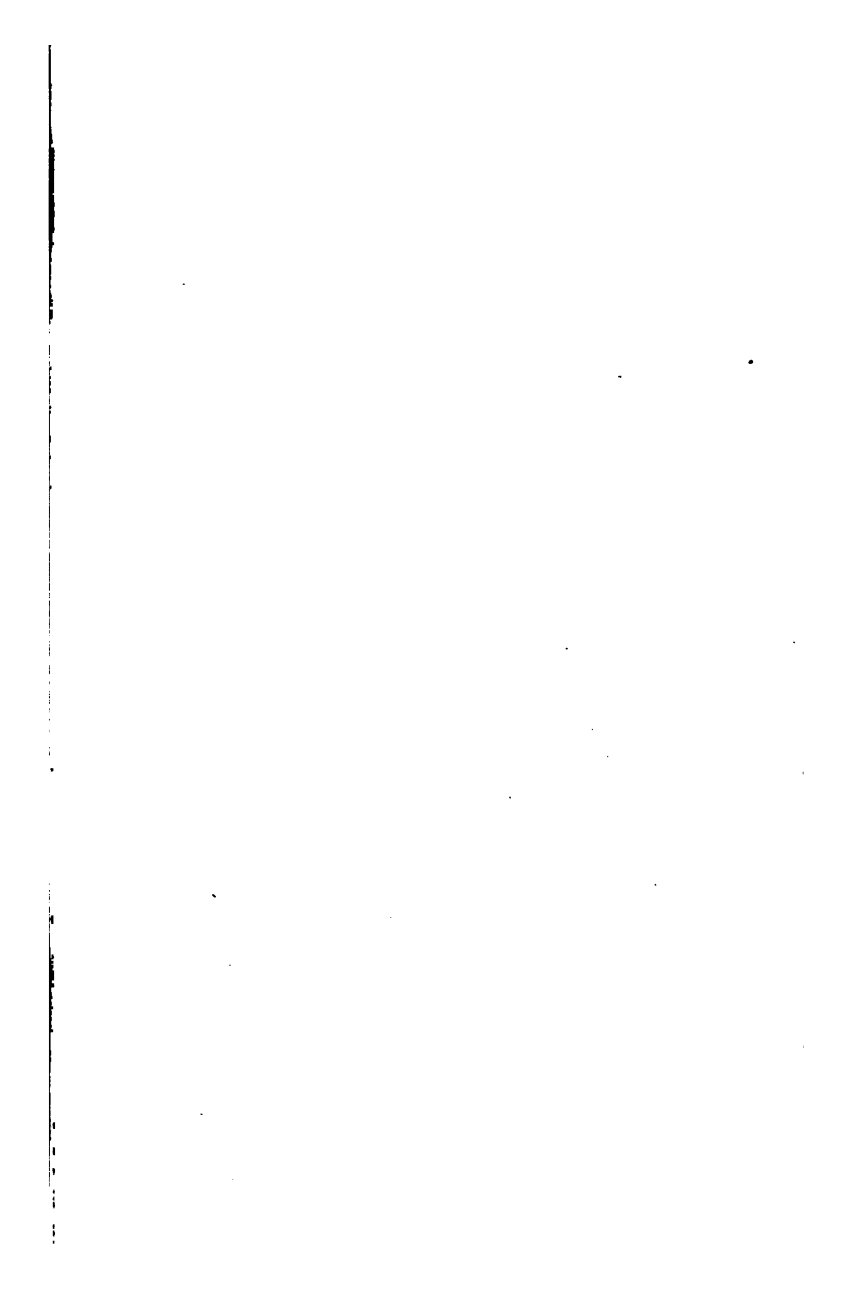
8. The barometer tube is divided into inches, and these into tenths and hundredths, so that the position

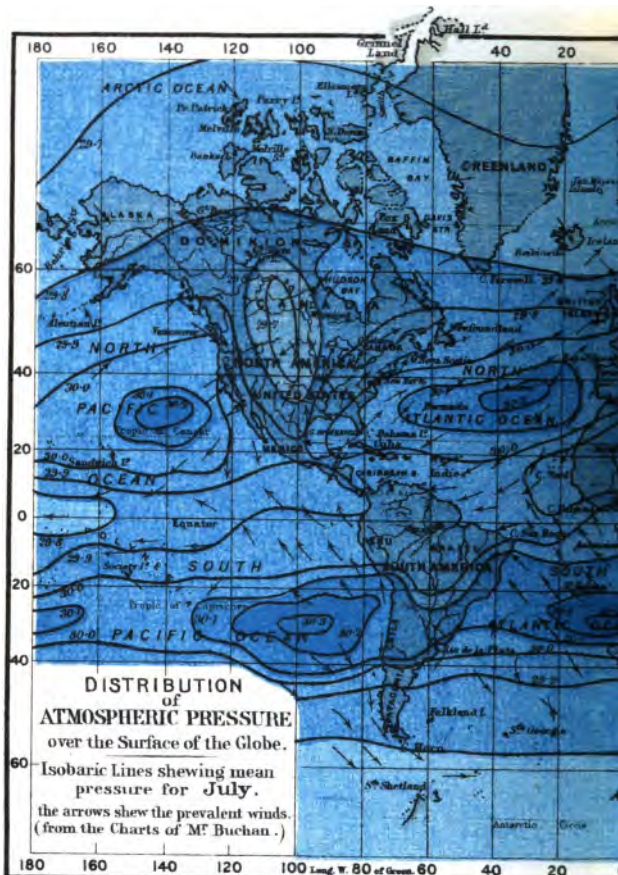
of the mercury is noted to the hundredth of an inch. When the pressure of the atmosphere just balances a column of mercury thirty inches in height, the barometer is said to mark or stand at 30. If the mercury falls half an inch the barometer stands at 29'50. If it then rises a tenth of an inch it is read as 29'60, a further rise of a hundredth of an inch would make 29'61. The height of the mercury in the barometer at the level of the sea has been found to be very nearly thirty inches as a mean over the whole globe. In different regions, however, the actual average height for the year varies considerably from that mean. In the Pacific Ocean, for instance, some distance to the westward of California, the mercury stands on the average at 30'30 inches. On the other hand, in the north-west of Iceland it stands at a mean height of 29'60 inches, while within the Antarctic circle the average is even considerably lower. When the mercury has fallen below its average height, it indicates a low pressure, when it rises above the average it marks a high pressure.

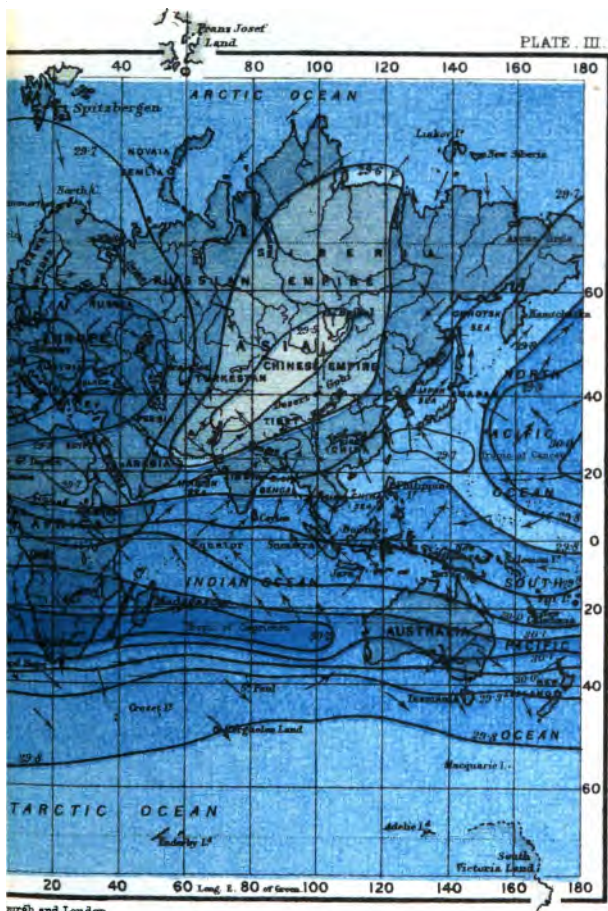
9. The importance of noting carefully the variations in atmospheric pressure will be apparent from the fact which has now been proved by observation in all parts of the world—that it is differences of pressure which give rise to winds, storms, and in short all the movements of the air which are intimately connected with changes of weather.

But what is the cause of these variations in pressure? Why should the air be liable to such increasing and often sudden, as well as great changes? The only certain answer which can at present be given to these questions is, that the pressure is affected, 1st, by temperature, and 2nd, by aqueous vapour.

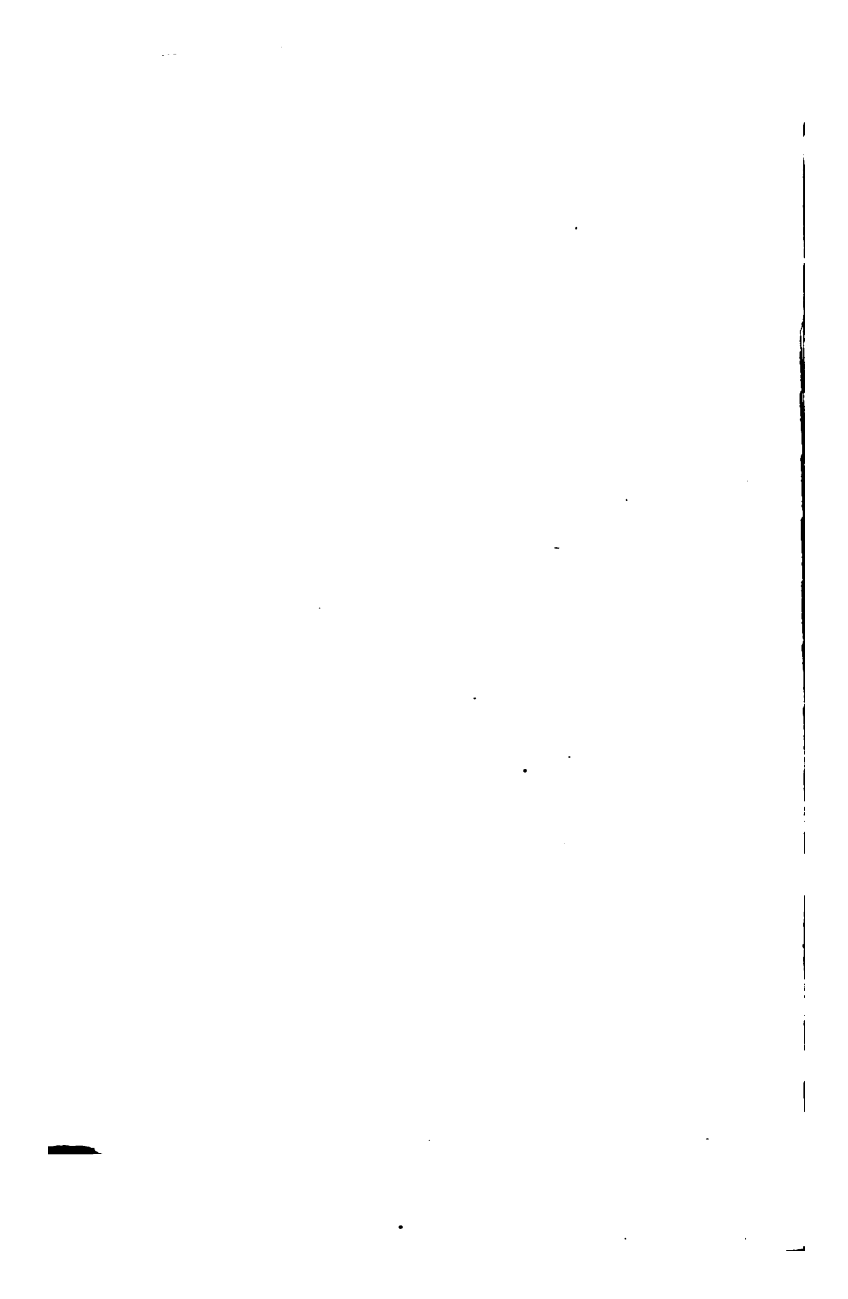
10. (1) **Temperature.** It is not difficult to see how this influence acts. When air is heated it expands, when cooled it contracts, behaving in this respect like other







burgh and London.



substances. Cold air is therefore denser than warm air, so that the latter ascends, while the former descends. The ascent of warm air must necessarily diminish atmospheric pressure. When a broad tract of the earth's surface, such for instance as the centre of Asia, is greatly heated by the sun's rays, the hot air in contact with the ground rises and flows over into the surrounding regions. Hence the atmospheric pressure there is lowered during the hot months of the year.

11. (2) **Aqueous Vapour** is however found to be a still more important agent in affecting the pressure of the air. We have already considered how universally present this invisible vapour is, and how easily it may usually be made visible by cooling the air; for then the vapour is at once changed into visible water. In Lesson X. an account is given of the vapour of the atmosphere, and the way in which it is continually passing into the air, and as constantly passing out of it again into some visible form of water. Let us consider here how this unceasing process affects the pressure of the atmosphere.

12. Suppose that we take two empty glass vessels each capable of containing exactly one cubic foot of any substance. By means of an air-pump we remove the air from them as completely as possible; into one of them we admit water-vapour, at a temperature of, say, 50° Fahr. (Lesson IX., Art. 1), till it can contain no more; the other, at precisely the same temperature, we fill with perfectly dry air, that is, air from which the water-vapour has been removed as thoroughly as possible. We then weigh them both, and after allowing for the weight of the glass-vessel in each case, we find that the vapour weighs only 4·10 grains, while the air weighs as much as 546·81 grains.

13. Now, without entering into the question how far atmospheric pressure is due to mere weight or to other

causes, what conclusion must we draw from this experiment? Evidently, that water-vapour is vastly lighter, or has less pressure, than air. At the temperature of 50° it is about 133 times lighter, and though the difference would be rather less at higher temperatures, still, under all the temperatures to which the atmosphere is ordinarily subject, the weight or pressure of the air is always far greater than that of the vapour.

14. Suppose, again, that we take six glass vessels, each holding exactly a cubic foot, and fill them as follows:—three with air saturated with vapour, each at a different temperature, the first, say at that of freezing-water (32°), the second at that of an ordinary English spring morning (50°), and the third at that of a warm English summer noon (80°), and three with perfectly dry air at the same three temperatures. In the first three the air of each vessel contains as much vapour as its temperature will allow it to retain, and as we have seen that cold air cannot hold so much as warm air, we know that in the warmest vessel there must be a good deal more vapour than in the coldest. We weigh them all carefully as before, and find that the ice-cold moist air weighs about a grain and a quarter (1.27 grain) less than the perfectly dry air at the same temperature, that the moist air at the medium temperature is about two and a half grains lighter than the dry air, and that the warmest moist air is about six and a half grains lighter than the warmest dry air.

15. What lesson does this experiment teach? Plainly that the addition of water-vapour makes air lighter or diminishes its pressure, and that this change is greater the warmer the air, for more vapour can be dissolved in warm than in cold air.

The vapour which rises so abundantly from sea and land into the atmosphere diffuses itself through the air,

pushing the-air atoms aside ; and having far less weight and more elasticity than the air, it necessarily lessens the density of the atmosphere, or in other words, lowers the atmospheric pressure. A mixture of air and vapour is lighter than the same volume of dry air would be, and, of course, the larger the proportion of vapour the greater will this difference be. •

16. The quantity of vapour in the atmosphere is constantly varying from day to day, and from season to season (Lesson X.). We perceive, therefore, that this cannot fail to be at least one grand cause of the ceaseless oscillations of pressure which the barometer reveals. The addition of a large volume of vapour to the atmosphere lowers the atmospheric pressure, and the mercury in the barometer therefore falls. The removal of this vapour either by condensation into rain, or otherwise, restores the pressure and the mercury rises again. Sometimes these fluctuations are very slow, extending over days or weeks, sometimes a great oscillation may take place within a few hours.

17. How it is that these vast changes in the volume of vapour in any part of the atmosphere are brought about is still unknown. -It is well ascertained, however, that they determine movements of the air. When sudden and extensive they are accompanied by storms of rain and wind. When less violent they still show their influence upon the winds and weather.

18. All movements of the atmosphere arise from differences of pressure, caused, as far as we can tell, by variations in temperature, and water-vapour. We shall consider these in the next two lessons.

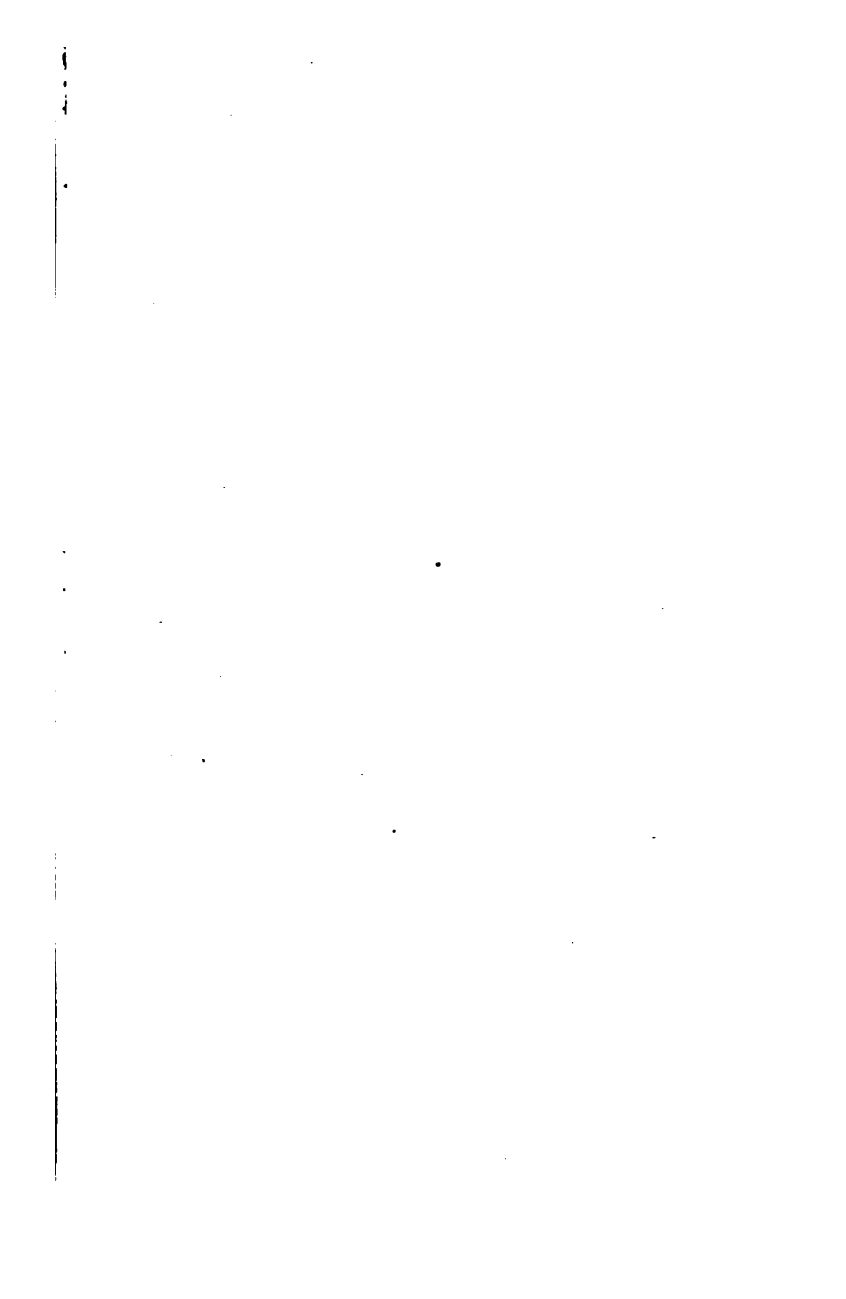
19. Observations with the barometer carried on for many years in all quarters of the globe, have enabled meteorologists to prepare charts showing, by means of lines of equal pressure called *Isobars*, the general dis-

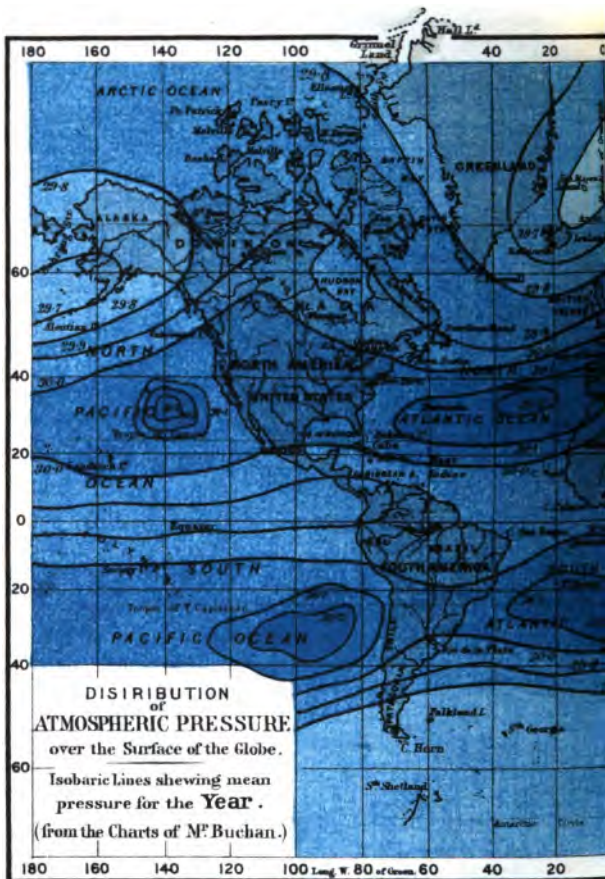
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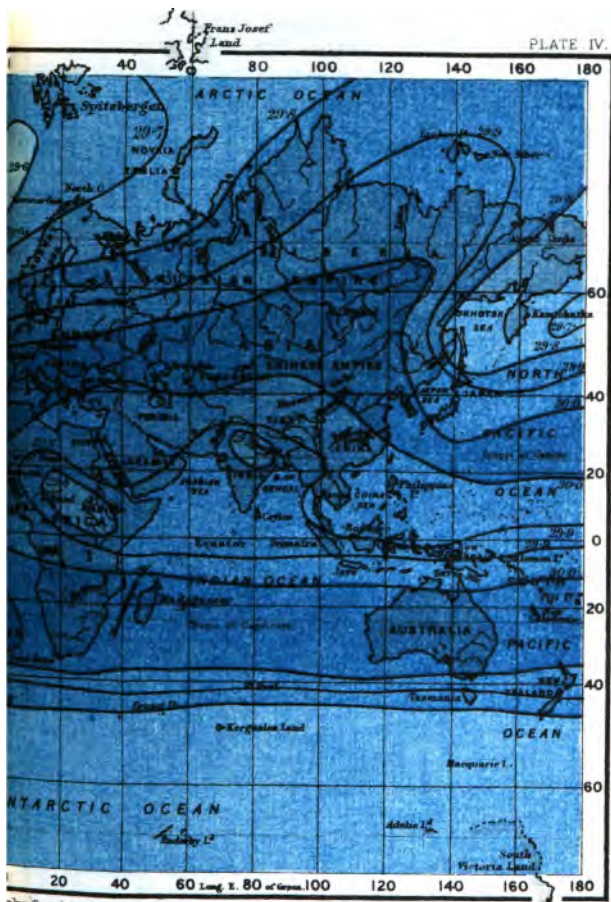
tribution of atmospheric pressure over the earth's surface for each month or season, or for the whole year. Plates II., III., and IV., are examples of such charts. It is found that, taking a very broad view of the subject, there are three great areas of low pressure. One extending as a broad belt round the equatorial regions, the other two lying about each pole; and two areas of high pressure extending parallel to the equatorial belt on either side, and separating it from the polar area of low barometer. These areas are most continuous in the southern hemisphere; but even there, and therefore still more in the northern, they are apt to be broken up into detached portions, owing to the irregular distribution of sea and land. They change position too with the seasons, as is well shown by comparing the distribution of pressure in January with that in July, as is done in Plates II. and III.

LESSON IX.—THE TEMPERATURE OF THE AIR.

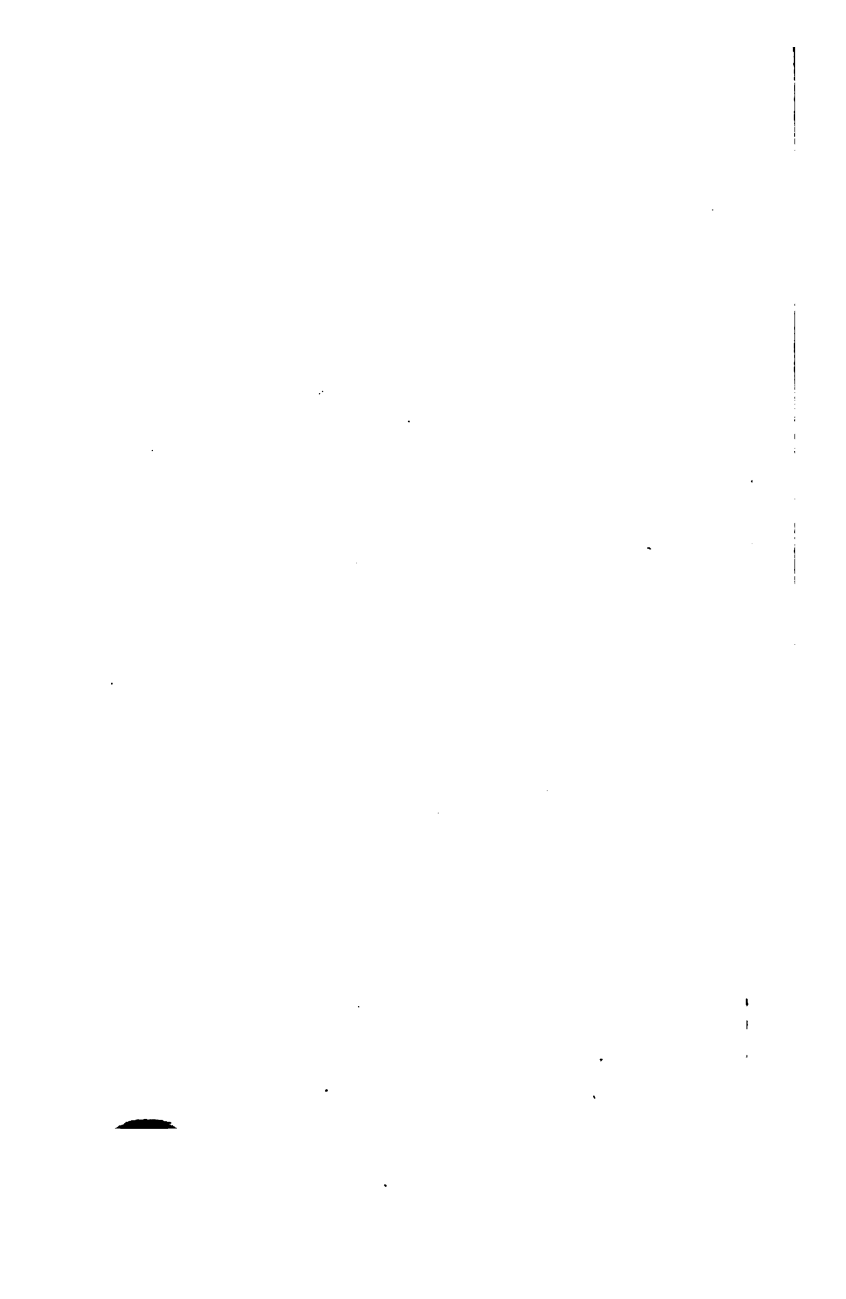
1. As in the study of the pressure of the air so in that of its temperature, it would not be possible to make much progress without some means of accurately measuring the fluctuations, for it is only the more marked of these of which we are sensible from the comfort or discomfort they bring to us. Happily, in this case, too, the instrument for accurate measurement is exceedingly simple both in its construction and use. It is called the Thermometer, or heat-measure, and consists of a small glass tube closed at both ends, the lower of which is expanded into a bulb. The tube, previously deprived as completely as possible of air, is partially filled with mercury or spirits of wine, and attached to a flat piece of ivory, wood, or other substance, on which is placed a graduated scale. Under the influence of heat the fluid







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in the tube expands and rises; when the heat is withdrawn it contracts and descends. The temperature is expressed by the figure on the scale of degrees, opposite to which the upper end of the mercury column stands. The thermometer scale used in this country, called after the original maker, Fahrenheit's, is so graduated that when the instrument is placed in melting ice or in water just about to freeze, the mercury marks 32° .¹ This is the *freezing-point* of fresh water. On a pleasant summer day in England, the mercury may mark 70° . On a hot noon in India it would rise to 90° or more, while on the burning sands of an African desert, at the hottest time of the day, it might sometimes stand as high as 150° , or even higher. Under ordinary circumstances the thermometer indicates 212° as the temperature at which water boils. When the mercury is low in the tube, it marks cold, or what is called a low temperature; when it stands high in the tube it indicates heat, or a high temperature.

2. By means of the thermometer it is possible to measure very minute changes of temperature, and to compare the range of temperature in different places. Observations of this kind have now been carried on for many years in all parts of the world, with the result of making known the general distribution of temperature over the globe. To show this distribution it is usual to construct maps on which lines are drawn through all places having the same temperature (see Plates V., VI., and VII). Such lines have received the name of Isothermal lines, or lines of equal temperature. Each of them is named after the degree of the thermometer which it expresses, as, for example, the isotherm of 60° , which shows that all the

¹ The zero (0°) of Fahrenheit's scale is 32° below the freezing-point of fresh water. When the temperature sinks below zero it is marked by prefixing the *minus* sign, thus -5° , -10° : that is five, ten degrees below the zero-point.

places through which it is drawn on the map have the average temperature of 60° .

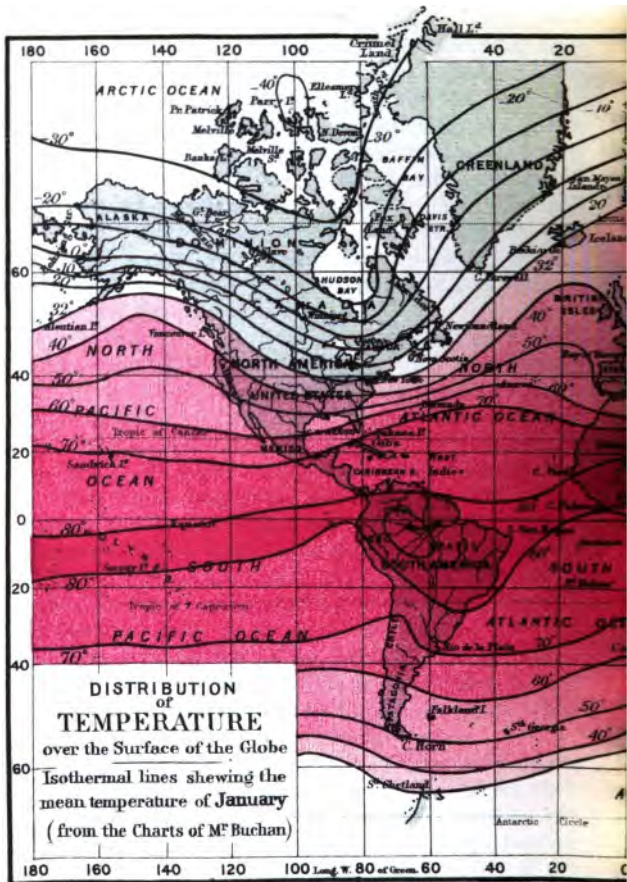
3. Whence does the earth receive its heat, and why does the temperature of one part of its surface vary so much from that of another?

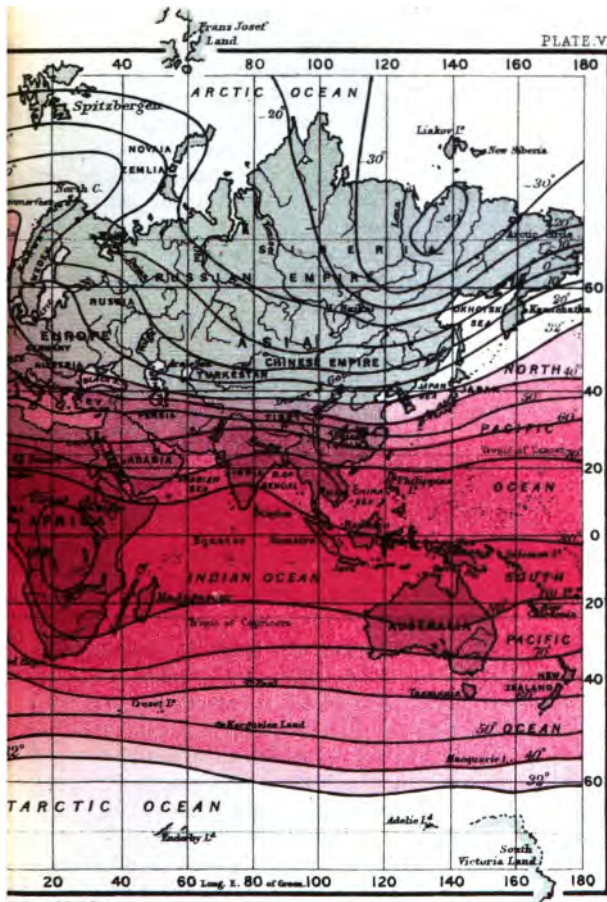
4. Although, as we have seen (Lesson III.), our planet was once probably a molten globe, and retains even now a vast amount of heat in its interior, its surface temperature is not materially affected thereby. Were it left without any other source of heat than its own, its surface would become so intensely cold as to be utterly uninhabitable by at least the races of plants and animals which now live upon it.

5. It is from the Sun that our supply of heat comes. At noon, when the sun shines brightly, we feel his rays to be warm. At night, when the sky is clear, we find it cold, because the sun's rays no longer fall upon our part of the earth, and the heat absorbed during the day is given off again into cold space. The sun is continually radiating heat from his glowing mass, and it is according to the greater or less amount of this heat which different parts of the earth receive that they vary in temperature.

6. It must be remembered, however, that the air is only to a very slight extent warmed by the passing through it of the heat-rays of the sun. No matter how hot these rays may be felt by us, they do not of themselves sensibly warm the air. It is when they have heated a part of the earth's surface that the air resting upon that surface becomes heated by contact with it.

7. The heating power of the sun's rays has been ascertained to be dependent upon the angle at which they reach the surface of our planet. Wherever they fall vertically upon the earth, as at B in Fig. 8, their heating power is greatest. It diminishes as the direction of the rays recedes more and more from the vertical, until, when





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they become horizontal, as at A and C, it reaches its lowest point. Hence, though powerful at noon, they are comparatively feeble in the morning and in the evening.

8. In considering, then, the manner in which the temperature of the atmosphere is distributed over the globe, we see at the outset that those countries must necessarily be warmest where the solar rays are vertical or most nearly so, and that those regions must be coldest where the rays are most oblique. Between the tropics (Lesson II., Art. 11,) the sun is vertical twice a year. That belt of the globe must needs therefore have the highest

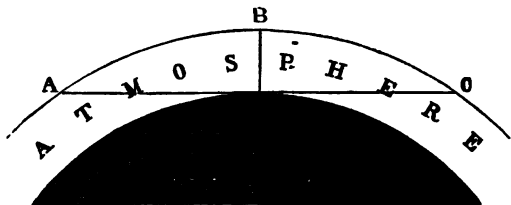


FIG. 8.—Diagram showing the influence of the varying thickness of the atmosphere in retarding the Sun's heat. A. Line of Sun's rays in the morning. B. Line of the rays at noon. C. Line of the rays at sunset.

temperature. Round the poles the sun does not shine for six months in winter, and never gets very high in the sky even in summer. These must consequently be the coldest regions. Hence the first conclusion we may draw regarding the temperature of the atmosphere is that it depends upon distance from the equator, or, expressed more shortly, (1) **temperature is regulated by latitude.**

9. If no counteracting influence came into play there would be a regular diminution of temperature from the equator to the poles. Every latitude should then have its own temperature, so that a determination of the

average temperature of a place should give us the latitude ; or, knowing the latitude, we should know also what must be its temperature.

10. But this correspondence holds only to a very limited degree. The places which may chance to lie on the same parallel most frequently do not enjoy the same temperature. London, for example, coincides in latitude with Labrador. But in England the summers are not hot, and the winters are not very cold, while in Labrador the summers are mild, and the winters rigorously cold. There must, therefore, be some other cause at work modifying the influence of latitude. What is this secondary cause ?

11. When Isothermal lines are drawn round the globe so as to connect together all the places which have the same general temperature, these lines, instead of following the parallel of latitude, are found to bend up and down, and these bendings are observed to have reference to the place of the continents and oceans. Any temperature may be chosen. We may select for example a temperature of 32° Fahr. or of 40° , 50° , 60° , or any other arbitrary point on the thermometer, it being understood that over every part through which the line is drawn the general temperature during the year or one of the seasons or months has been found by prolonged observation to be the same, and amounts to the number of degrees fixed upon. In this way we may represent the mean temperature of the whole year, or that of winter or of summer.

12. By referring to the maps in Plates V., VI., and VII., which represent the distribution of temperature over the globe for January, for July, and for the average of the whole year, it will be seen how these lines are drawn and how simply they tell the story of the arrangement of temperature upon the earth. They are least irregular in the southern hemisphere over the wide expanse of ocean ;

they show the greatest deflections across North America, the Atlantic, Europe, and Asia. In this way they indicate that temperature is more uniform and more directly dependent on latitude in the oceanic parts of the globe than in the continental, or where the oceanic and continental come together, as they do in the basin of the Atlantic.

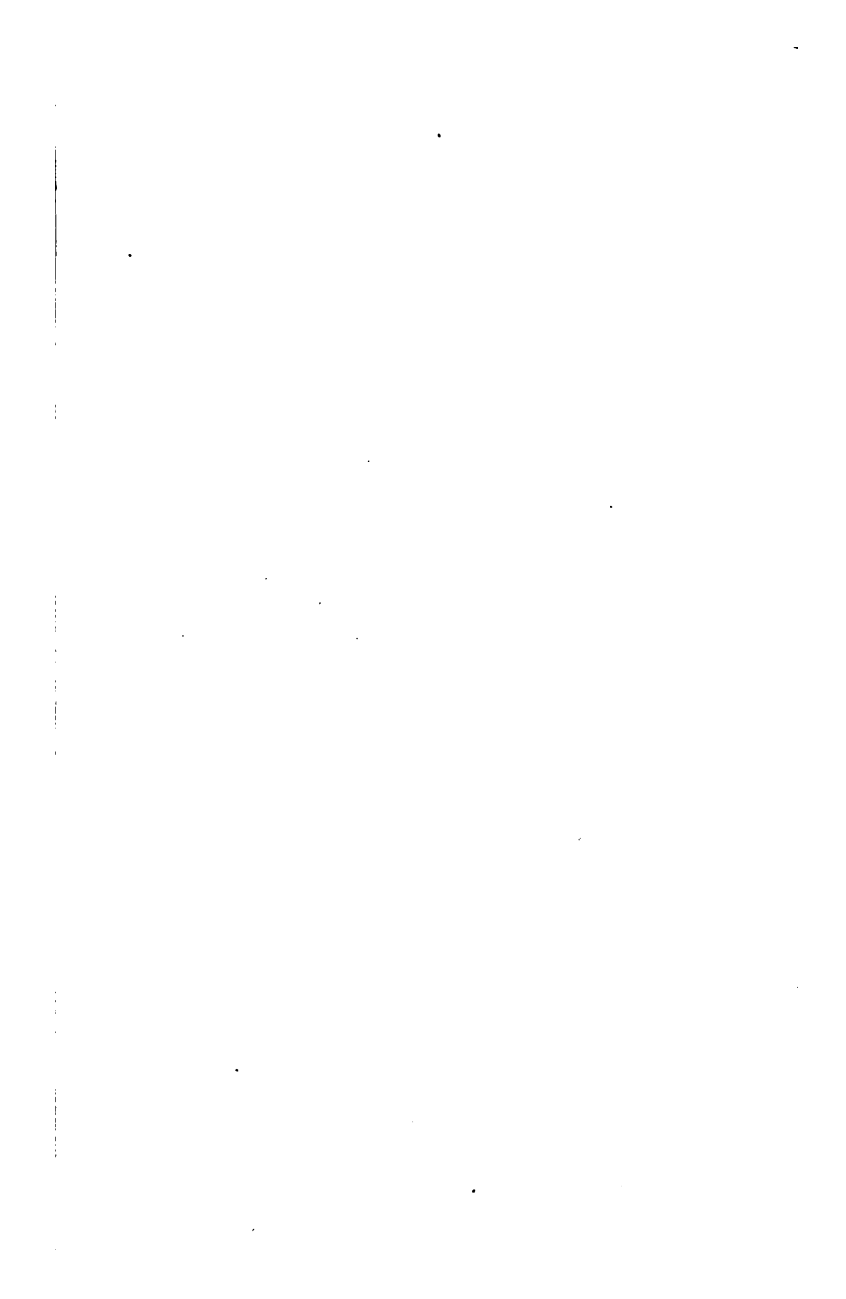
13. In order to see the full use and meaning of these isothermal lines, let us take by way of illustration the line in the northern hemisphere marking a mean annual temperature of 50° Fahr. Traced over Britain, this line runs from about London, across the centre of England and the north of Wales: that is to say, all the parts of the country lying along that line have a mean annual temperature of 50° , while the districts to the north-east are a little colder and those to the south-west a little warmer. The line bends south-westward and crosses to the western shores of Ireland. If now we turn to the opposite side of the Atlantic to see where places are to be found having the same average temperature for the year, we do not meet with them on the same parallel of latitude as in the British Islands. They lie much further south, so that the line or isotherm of 50° makes a bend in crossing the ocean, and reaches the American coast not far from New York. The mean annual temperature of London and New York is the same. And yet New York is about as far south from London as Madrid.

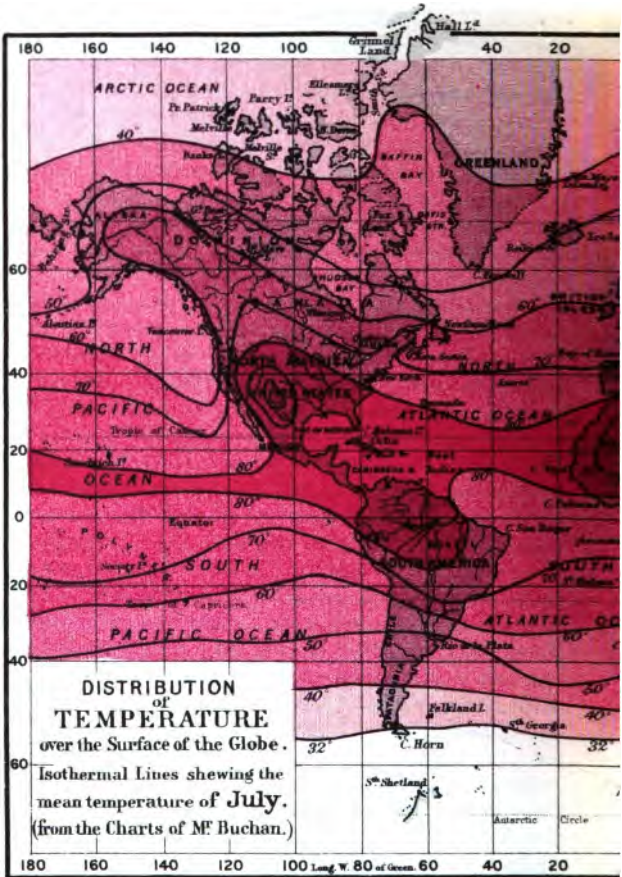
14. This isothermal line of 50° between Europe and America and still more the other lines lying to the north of it, illustrate how far the belts of equal heat are from coinciding with the parallels of latitude. It will be seen from the maps how this divergence is determined by the way in which the areas of land and sea are grouped.

15. Land gets sooner heated by the sun's rays than the sea and also gives off its heat again sooner. The

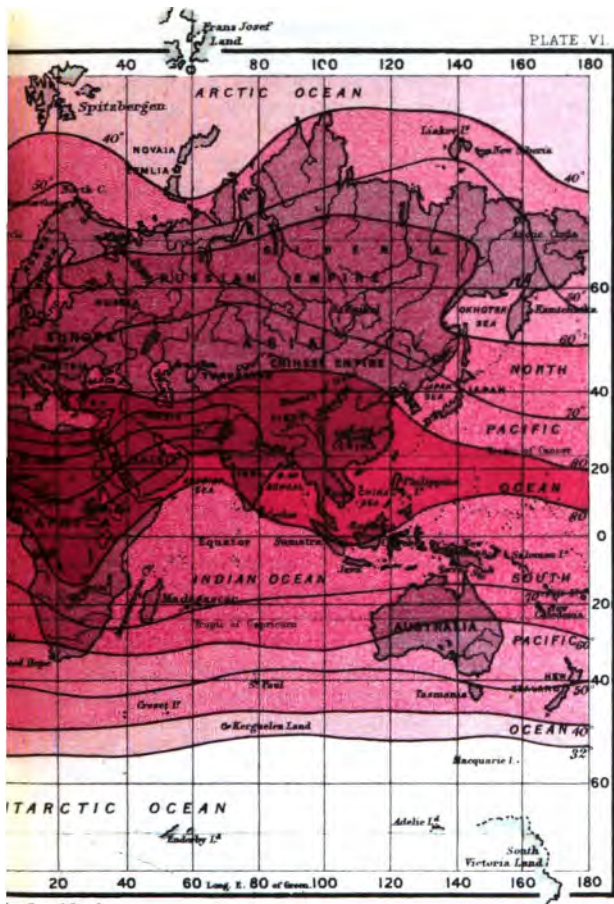
sea, though it does not get so hot as the land does, retains its heat longer, and is enabled by virtue of its liquidity and motion to diffuse it. Hence, the influence of the sea tends to mitigate both the heat and the cold of the land. Its warm currents heat the air resting on them, and so give rise to warm winds which blow upon the land, while its colder waters in like manner temper the air, which reaches the land in cooling breezes, or it may be in cold damp winds and fogs. Thus, in the basin of the North Atlantic, a warm ocean-current called the Gulf Stream issues from the Gulf of Mexico, and, augmented by the surface-drift of warm water which is driven onward by the prevalent south-west winds, flows across the Atlantic to the shores of Britain and even of Spitzbergen. It brings with it the supplies of heat which make the climate of the west of Europe so much less cold than it would naturally be. On the other hand, an icy stream of water, coming out of Davis Strait, brings a chill to the coasts of Labrador and Newfoundland. The ocean, therefore, by its cold currents is depressing the temperature in America along the same latitudes where in Europe by its warm currents it is raising it.

16. A large body of land placed in high latitudes by accumulating snow and ice, gives rise to a lower temperature, and a similar mass of land in low latitudes, by exposing a broad surface to the tropical rays of the sun, produces a higher temperature than would be found if the region were occupied by the sea. In illustration of this statement it will be seen from the map (Plate V.) that the January isothermal lines marking temperatures below the freezing-point, come a good way southwards over Northern Asia, and again over Greenland and North America, though in the water passage between these regions they swell a long way northwards. Again, the map of the July temperature (Plate VI.) shows that over





DISTRIBUTION
of
TEMPERATURE
over the Surface of the Globe.
Isothermal Lines shewing the
mean temperature of July.
(from the Charts of MF Buchan.)



burgh and London.



the equatorial parts of America and the great mass of Africa and Southern Asia, the space inclosed between the isothermal lines of 80° swells out greatly, so as to include a much larger area than where these lines cross the oceans.

17. It by no means follows that two places which have the same average yearly temperature enjoy the same climate. For instance, Reykjavik in the south of Iceland (Lat. $64^{\circ} 40'$ N.), has a mean annual temperature of about 38° Fahr. while the city of Quebec has one of about 40° ; but in July the average temperature in the former locality is 51° , in the latter it is 70° ; while in the one case the January temperature is 30° and the other 12° ; so that in winter Quebec is usually so intensely cold as to have 20° of frost, while the south of Iceland is often without frost. In summer, on the other hand, Quebec is 19° warmer than the south of Iceland. Canada is chilled by the cold land and sea lying to the north and north-east of it. Iceland is warmed by the ocean-current (Art. 15) which summer and winter sweeps past its shores.

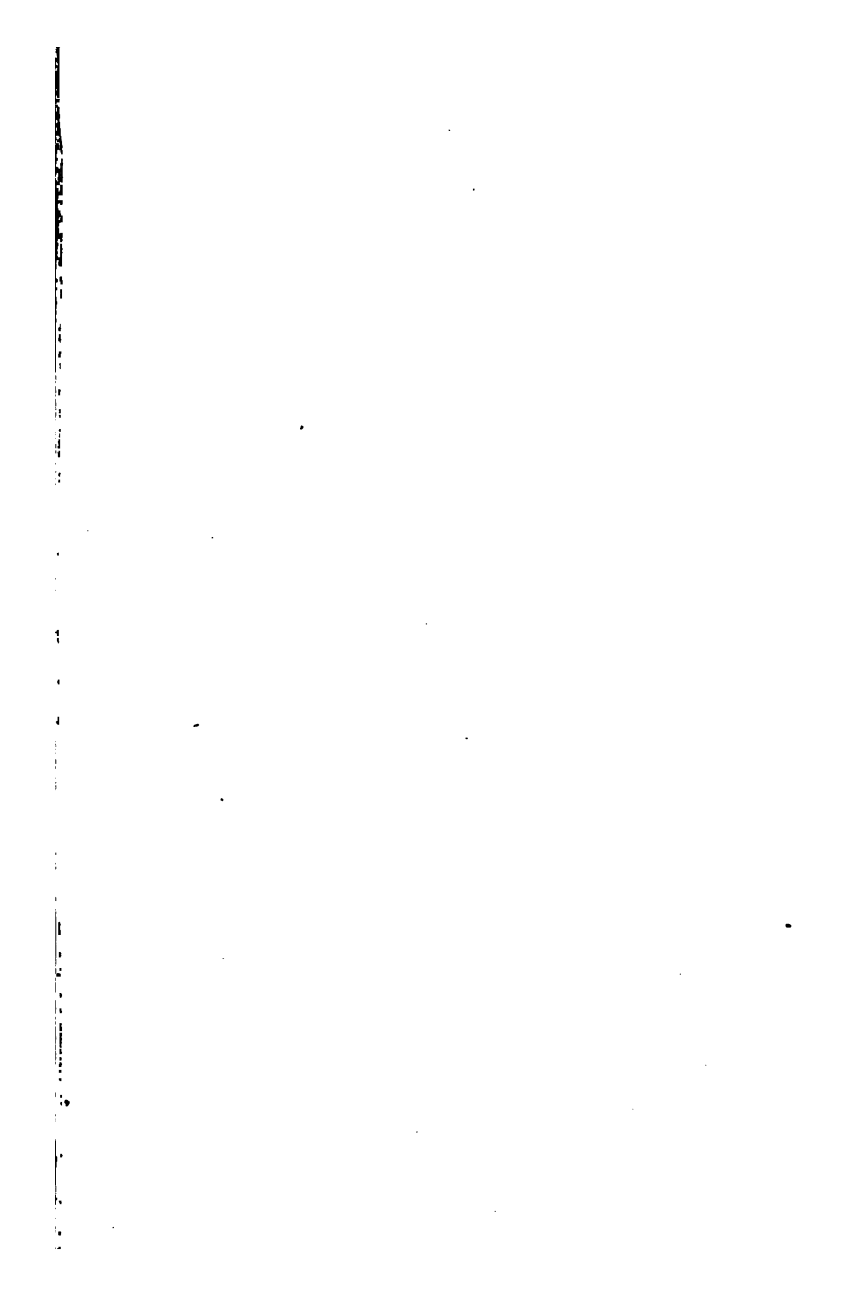
18. In order to compare the climates of two places, it is necessary to know how the temperature is distributed through the different seasons. To aid comparisons of this kind, charts are prepared like those in Plates V. and VI., showing the average distribution of temperature for each month, or for summer and winter, as well as charts like Plate VII., constructed to show the mean temperature of each part of the globe for the whole year.

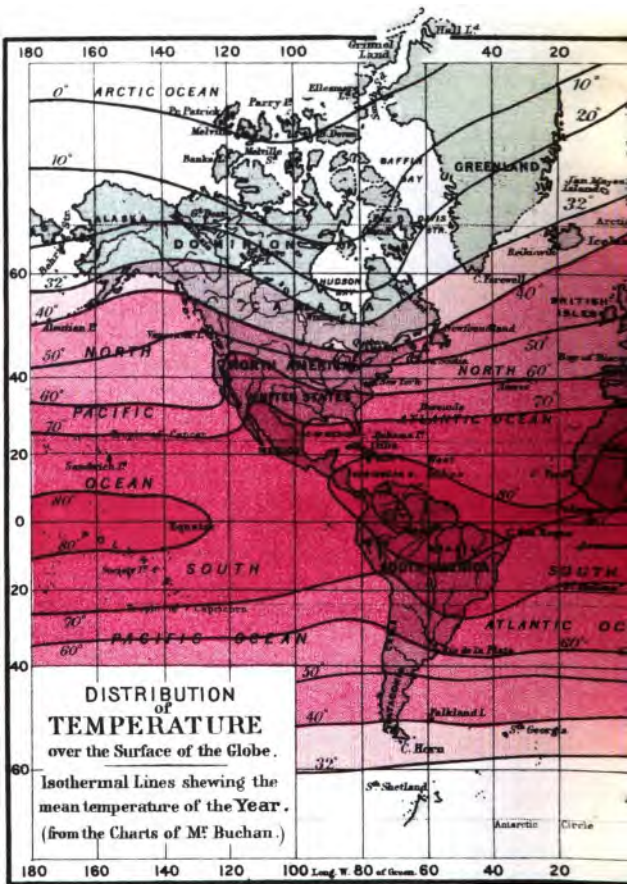
19. From the facts embodied on such charts, gathered as they have been, from all parts of the world, we conclude that (2) **temperature is regulated by the distribution of sea and land.**

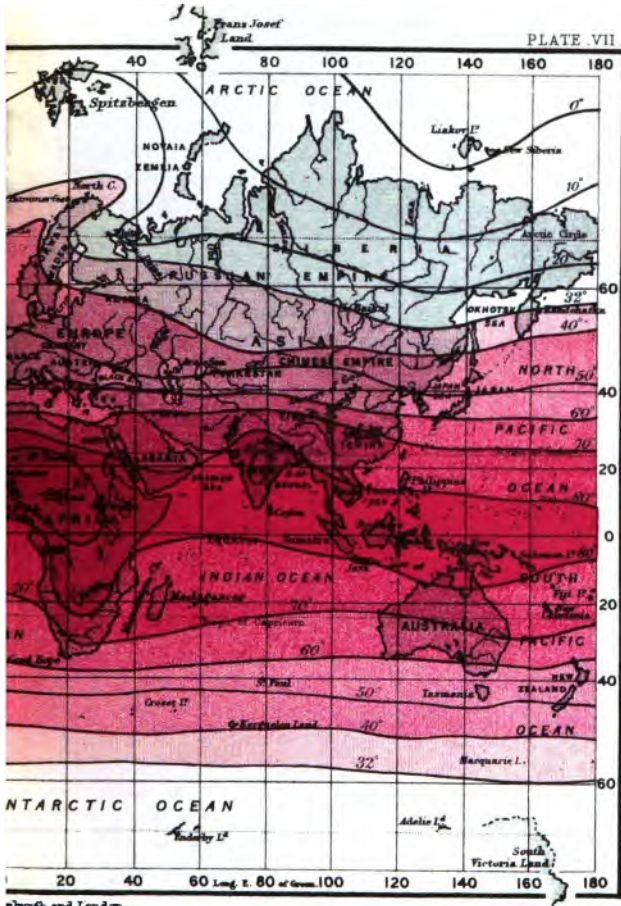
20. But there is still another cause on which the temperature of any particular place on the earth's surface is dependent. The two influences already given, viz., latitude

and the grouping of land and sea, come into play *horizontally* over the globe ; the third influence acts *vertically*. We are familiar with the fact that the air on low-lying land is warmer than it is on the tops of hills. Even on many of the British mountains, which are, comparatively speaking, low, the summits are so cold that in sheltered places screened from the sun and wind, snow remains unmelted upon them through the summer. In the upper parts of the Alps, the Himalaya, the Andes, and all the higher mountains of the globe the cold is such that the snows of winter never wholly disappear, but remain as a perpetual covering. There is a gradual cooling of the air perceptible as we rise above the sea level in every part of the world. The rate at which this fall in temperature takes place varies much, but it is commonly taken to be on an average 1° Fahr. for every 300 feet. Since within the tropics, while the low lands are lying under a burning heat, the mountains, if lofty enough to reach the upper cold air, have their summits covered with snow, we perceive that an elevation of a few thousand feet above the sea gives as great a change of temperature as a journey of many thousand miles away from the equator would do. So that from these examples we are taught that (3) **temperature is regulated by height above the sea.**

21. Since the earth is continually receiving such vast stores of heat from the sun, we may be tempted to ask whether it is not getting warmer. So far as observations of temperature have yet been made, they do not indicate any sensible increase or diminution of heat. It would seem that the earth radiates back into space as much heat as it absorbs. The quantity of heat received from the sun, may be looked upon as remaining on the whole constant from year to year ; though careful observations of the sun's surface, and especially of the recurrence of







burgh and London.

the black spots visible there, may yet show that the quantity varies from time to time, and that the variation affects the temperature and climate of our planet. It has been satisfactorily established that there is a coincidence between heavy rainfall, including storms, and those periods when the sun's face is most marked with spots.

22. Radiation, or the giving off of heat by the earth, is felt most at night, and especially when the sky is clear. At such times we may learn how rapidly the heat of the day is passing away from the earth into cold stellar space, and how entirely our globe depends on the sun for its present surface temperature. Objects which in the day were warm to the touch, now grow more and more chilly. The air gets colder by contact with the cooled ground, and our own bodies, like everything else around, are at the same time radiating heat, and adding to our sensation of cold.

LESSON X.—THE MOISTURE OF THE AIR.

1. One of the constantly present ingredients in air described in Lesson VI. was the vapour of water. We have found how important this component is in determining differences of pressure, and consequently in giving rise to changes of weather. It may now be considered rather more at length in reference to its sources of supply, and the different forms in which it is taken out of the air, and restored again to land and sea.

2. First then let us ask whence does this widely-diffused and all-important vapour come? It is all evaporated, or given off in an invisible form, from the surface of every sea, lake, river, and spring; in short, from every water-surface on the face of the earth, and even from ice and snow. Nothing is more familiar than the rapidity

with which water dries up on streets and roads after rain. Every accumulation of water, indeed, which is freely exposed to the air, and is not replenished with water, is seen to diminish, and finally to disappear. It is not that the water sinks into the ground. Part of it no doubt does so, but we find the water to disappear even from a saucer or any other vessel where there can be no loss underneath.

3. The air is usually busy absorbing vapour. When it can hold no more, it is said to be saturated, or to reach its point of saturation, and then evaporation ceases. This limit varies according to temperature, warm air, as was proved in Lesson VIII., being able to contain more vapour than cold air.¹ Evaporation is greatly helped by wind. Damp places and pools of water, for instance, dry up in a breeze sooner than they do in still air, because the wind removes the vapour as it is formed, and brings other and drier air to drink up and carry away renewed supplies of vapour.

4. Evaporation, therefore, must take place chiefly during the day, and especially the warm parts of the day. It must be more active in summer than in winter. It is feeble in amount when the air is moist and still, but goes on briskly when a fresh wind blows. It takes place far more abundantly in warm tropical regions than in those with a temperate or polar climate.

5. It has been calculated that the amount of water which is annually condensed out of the atmosphere upon the surface of the earth would, if collected together into one mass, cover to the depth of one mile an area of about 200,000 square miles, or a space nearly equal to the size of the whole of France. This vast liquid mass, is as it were pumped out of the sea and of the waters of the

¹ Air at 50° Fahr. when saturated with water-vapour contains 9·74 grammes of it in every cubic metre of air.

land by the sun's heat. But the amount of water which passes off into the vapour of the atmosphere is best realized when we consider the enormous quantity discharged into the sea by rivers. All over the world, rivers, great and small, are continually pouring their burden of water into the ocean. All that water was supplied from the atmosphere to them either directly by rains and snows, or indirectly by springs. But they obtained originally even more than that vast quantity which they discharge, for while they are coursing down from the mountains to the sea, vapour is continually rising from their surfaces and their volume is consequently growing less.

6. If, then, from every ocean, lake, and river on the surface of the globe water-vapour is continually passing into the air, what becomes of all this vapour, and why do not the waters of the globe grow less? The reason is, that the conversion of water from its common visible liquid form into the invisible gaseous state is only one-half of a gigantic system of circulation. The vapour is not allowed to accumulate indefinitely in the atmosphere; it is changed back again or condensed into water, and then appears in such forms as dew, clouds, rain, or snow.

7. The two processes of evaporation and condensation seem on the whole to balance each other; that is, so far as the larger features of the earth's system are concerned, there appears to be about as much water returned to the sea and land as is taken from them by the atmosphere. It is from this circulation of water that all the manifold phenomena of clouds, rain, snow, rivers, glaciers, and lakes arise. But more than this, if we reflect that now evaporation and now condensation must from time to time be predominant at any place, we perceive how greatly these processes must affect the pressure of the air; and since variations

in pressure largely determine the various movements of the atmosphere (Lesson VIII.), we come to see how all-important this water-vapour must be in the present arrangements of our globe.

8. So far as we can tell, though the quantity of vapour may sometimes be greatly reduced, it never disappears wholly from the air at any place. On the other hand, the air round about us is comparatively seldom so saturated with moisture as not to be able to hold any more, though in damp weather we may see from the extreme slowness with which wet places dry up, how little appetite the air then has for further vapour.

9. One office of the vapour of the atmosphere is to keep the earth much warmer than it would be were the air quite dry. It interposes as an invisible screen between the earth and the rays of the sun, which would otherwise be intensely hot. And it is the same screen, often condensing into visible form as clouds, which at night prevents the earth from giving off too rapidly its heat into cold space. If it could be removed for a little from around us, we should be burnt up by day, and frozen hard at night. Clouds would cease to form, rain to fall, and streams to flow, and the condition of the planet as a habitable globe would be brought to a stand.

10. When evaporation takes place, heat is abstracted by the vapour from the surface which evaporates. If you spread a drop of water over the back of the hand, you feel the skin to be chilled a little, because the water in passing into vapour abstracts heat from the hand. It is a common practice to cool liquids by placing moist flannel round the vessels in which they are contained. The moisture of the flannel evaporates, and in so doing takes some of the heat away from the vessel inside. It is clear, therefore, that the invisible vapour which rises into the air so abundantly carries heat along with it. But this

heat is not sensible so long as the vapour remains uncondensed, and it is therefore termed *latent*.

11. When, on the other hand, condensation goes on, the heat which the vapour had held in its grasp is given out again and becomes *sensible*, as the vapour passes into water. It has been pointed out, for example, that every pound of water which is condensed from vapour liberates heat enough to melt five pounds of cast iron. We can well understand, therefore, that when the process of condensation takes place in nature on a large scale, the conversion of vapour back again into the state of water materially warms the air.

12. The act of condensation always occurs when air is cooled down to its *dew-point* (Art. 14); but this does not always happen at the same temperature, nor does it appear in the same forms. Sometimes it issues in a light mist, or beads of dew, or drops of rain, or, if the temperature should be low enough, in flakes of snow or pellets of hail.

13. The substance which we call water is thus shown to exist in three forms, according to temperature. At ordinary temperatures, or from 32° to 212° Fahr., it is abundant as a *liquid*, and this is, of course, its most familiar condition. If at any temperature, it is allowed to stand exposed to the air, or if by having its temperature raised to 212° , it boils, it passes off into invisible *vapour*. If again the temperature be depressed to 32° , the water begins to become *solid*. It crystallizes into the brittle, colourless substance called ice. This act of crystallization has received the name of *freezing*, and the temperature (32°) at which it takes place, is called the *freezing-point*. According, therefore, to the temperature at which condensation takes place, the vapour is made either to assume a liquid or a solid form.

14. *Dew*.—On a summer evening, if the sky is clear, condensation takes place on the leaves of plants, on stones and

other objects. These become covered with fine drops of water, which is known by the name of Dew. If the sky is cloudy, dew either does not form or only in small quantity, and on such cloudy nights the air is not so cold as on those which are clear. The cause of the appearance of this moisture is the same as that of the mist on a tumbler of very cold water brought into a warm room. The dew comes out of the vapour of the atmosphere, and not out of the objects on which it forms. When the sky is cloudless, radiation goes on rapidly from the earth, and the surfaces of those substances which part with their heat most readily become much colder than the air. Grass, for example, cools about twice as much as common garden-soil. As the process goes on, the air next the cooling surfaces is so chilled as to be no longer able to keep all its vapour, part of which is condensed and appears as dew. Hence the grass soon becomes quite wet with the copious gathering of dew upon its blades. The temperature at which this condensation takes place is the point of saturation or dew-point.

15. It has been pointed out (Art. 10, 11,) that the vapour of the atmosphere contains a large amount of latent heat, and that when the vapour is condensed this heat becomes sensible. The formation of a film of dew upon the surface of the earth lets heat escape back again into the air. But as radiation proceeds, the warmed air is again chilled down to the dew-point, then more dew is abstracted, and more heat given out. In this way the nights are kept from getting colder than the temperature of the dew-point. But sometimes, when the temperature is low, as in winter, or when radiation is great, the surface of the ground is so chilled that the dew freezes in forming, and appears as the white rime, or **hoar-frost**, which we see at early morning on the grass. Clouds hinder dew from forming because they check the passage of heat from the earth

into space while they themselves radiate heat towards the earth ; hence it is that cloudy nights are warmer than those which are clear and starry.

16. Mist and Fog.—When a mass of warm and moist air meets colder air, or comes in contact with cold ground or in any other way is cooled down below its dew-point, the excess of vapour which it can no longer retain condenses into minute particles, and is made visible in the form of Mist or Fog. In winter we have a familiar illustration of this phenomenon in the condensation of our breath into mist as it passes from the mouth into the cold air. In summer, mists frequently form in the evening over rivers and sheets of still water. As radiation goes on, the ground round the water becomes several degrees colder than the water itself, and the vapour rising from the water is consequently chilled by the air and condensed into banks or sheets of fog. Again, when a warm wind strikes upon a hill or mountain, and is forced to ascend the slopes, its temperature begins to fall, and if this cooling is carried below the temperature above which it can retain its vapour, that is its dew-point, the excess of vapour takes the form of a mist.

17. Clouds.—Dew and mist are formed on or near the ground, whether low plains or high mountains. But whenever the vapour is carried up into the cold upper parts of the atmosphere, it there becomes visible in another form of condensation as Clouds. A cloud is only a mist hanging in the air instead of resting on the ground. When the ground rises into the upper air, as it does in mountains, it reaches or even surmounts the layers of the atmosphere where clouds are commonly formed. We see wreaths of cloud gathering on the sides or summits of mountains, and yet if we climb up to these clouds we find them to be mere masses of mist, such as are formed down among the valleys.

18. Owing to the great changes which are constantly taking place in the temperature and evaporation at the earth's surface, currents of air must always be ascending and carrying vapour up with them. But besides this vertical movement, the atmosphere, as far upward as can be observed, includes many different layers or horizontal currents one above another, moving in various and even opposite directions. We may often notice the existence of these high aerial currents by watching the movements of the clouds. A lower layer of thick, woolly clouds may for example be rolling along, but through the rifts, a far higher layer of thin, light, white cloudlets may be drifting in a contrary direction. People who have made ascents in balloons, have found abundant evidence of the great number and various motions of these atmospheric currents.

19. When an ascending current of warm, moist air rises high above the surface of the earth, it meets with air much colder than itself, and it likewise loses sensible heat from the expansion consequent on the diminished atmospheric pressure at these elevations. Chilled down to its dew-point, it forms a cloud in the air. The cloud may be thought to remain perfectly motionless and unchanged in form, yet when watched attentively, it may be seen to alter, somewhat in the same way as a wreath or cap of mist on a mountain, and for the same reason as explained in Art. 23.

20. It is often easy to watch how clouds are formed and disappear. In summer, for instance, when the sky is quite clear in the morning, we may observe that as the day advances white clouds make their appearance, small at first, but growing by degrees from a more or less level base into huge piles with every variety of fantastic outline. As evening comes on these masses diminish slowly. At sunset, perhaps, only a few thin flaky cloudlets are to

be seen, and at last when night sets in, the sky is once more clear and cloudless. In such cases, it is to the warming of the earth by the sun during the day, and the consequent ascent of moist air-currents, that the clouds are due. Each growing mass of cloud forms the top or capital, as it were, of a column of up-streaming warm air laden with vapour. The air, expanding and cooling as it rises, at last reaches a point where it can no longer retain its vapour, and there the cloud begins to form. But after the heat of the day is over and the warm moist air no longer streams upward, the cloud ceases to grow larger. It then begins to sink earthwards, as radiation goes on, and coming into warmer air, it gradually melts away so as to leave the sky clear and starry.

21. As the atmosphere is traversed in all directions by currents of air of various temperatures and degrees of moisture, the meeting of these currents must often give rise to clouds, and also melt clouds already formed. A warm moist wind, for example, coming in contact with a cold wind will part with some of the vapour as cloud. On the other hand, when a belt of cloud is invaded by warm dry air, it will be evaporated and disappear. As a rule, clouds which are ascending in the air, increase in size as they get higher, while those which are descending get smaller, because the air in which they are suspended gets chilled in the one case, and warmed in the other. It is the continual movements of the atmosphere, which produce the never-ending comings and goings of the cloud-world above us.

22. When clouds get into one of the upper steady air-currents they are borne along sometimes for great distances and at a great rate. On a breezy spring day they may be seen sailing across the sky at what may seem a leisurely pace, which, however, is proved to be often more than eighty or ninety miles an hour by

the rate of their swiftly-moving shadows over hill and plain. If you watch them you see that they are continually changing shape and size as they move along, rolling in huge folds over each other, sometimes lessening and sometimes increasing, and in all these movements testifying to the ceaseless turmoil of the atmosphere in which they are suspended.

23. In hilly countries another curious feature of cloud-formation may often be seen. When a strong wind blows,



FIG. 9.—Clouds condensed by the cliff of the Noss Head, Shetland, with a clear sky and a S. E. wind. Sketched on June 14th, 1876.

sending leaves and dust high into the air, a cloud may be observed to remain persistently on a mountain-top, and only get bigger the more the wind rises. Fragments of the cloud may be torn off now and then, and carried in the direction of the wind, but these, after travelling some way, gradually disappear. In such cases, the wind comes along with its warm vapour, which is quite invisible until it strikes against the colder hill-side and undergoes expansion as it is forced up the slope into a higher level in the atmosphere. Chilled there, it passes

into a fine mist, which, seen from a distance, takes the shape of a cloud capping the summit. This cloud is stationary, but the little particles of which it consists are not. The wind continues to blow up and over the mountain-top, and the vapour which it carries along is made visible as mist or cloud, while it sweeps over the chill ground. When it has passed the mountain and mixes again with the warmer air beyond, the visible vapour is redissolved, and the cloud therefore melts away on the leeward side of the hill as fast, perhaps, as it formed on the windward side. The cloudlets, too, which are occasionally detached and carried away from the mountain-top by the wind, are gradually absorbed into the air again and so disappear. (Fig. 9.)

24. Considerable variety exists in the forms which clouds assume, from the thin flaky cloudlets seen in the high upper air down to the huge rolling rain-clouds which come down low upon the hills, and the dull grey sheet of cloud which sometimes overspreads the whole sky. To these various forms special names have been given which, however, it is not necessary to enumerate here. Each kind of cloud is formed in particular conditions of the atmosphere, and hence a study of the clouds furnishes valuable information regarding the weather. This subject properly belongs to the study of Meteorology.

25. The chief function which has to be considered in the mechanism of the clouds is the way in which they supply the earth with moisture. The vast amount of water raised into the atmosphere as invisible vapour returns to the surface of the earth again to feed the springs and rivers of the land and to replenish the sea. In this constant circulation the clouds act the part of condensers. They collect in a visible form the ever-present invisible vapour of the air, and allow it to fall back again upon the earth.

26. Rain.—By far the larger part of the vapour of the atmosphere descends to the earth in the form of rain. The little water-particles of which cloud is composed run together as the condensation proceeds. As the drops thus formed increase, they become too heavy to float in the air, and begin to fall earthwards as rain. At first they are very small, as we may feel if we happen to be on a mountain where the mists are gathering into rain-clouds. But as they descend through the air they increase in size until they reach the ground in well-marked rain-drops.

27. Rain, therefore, is a further stage in the condensation of mist or cloud. If the chilling of the cloud is prolonged, rain falls from it. This may be caused in several ways. For instance, where a warm moisture-laden wind comes against a range of high mountains, and is consequently forced to ascend, it may not only be condensed into mist (Art. 19), but if cooled still more, will drop its moisture as rain. Or a cold wind which is heavy and keeps next to the ground, may wedge itself in below a warm, moist layer of air, and so chill it as to form cloud and bring down rain.

28. The fall of rain being dependent upon the amount of evaporation is greatest in tropical regions, where the largest quantity of vapour passes up into the air, and diminishes with the gradual sinking of the temperature towards the poles. But this general law is subject to some important qualifications by the distribution of sea and land, and by the direction of the great aerial currents.

29. (1.) Although evaporation is more abundant from the surface of the sea than of the land, condensation is more active over land than over sea. Hence the rainfall is also greater over land than sea, and over the northern hemisphere, much of which is land, than over the southern hemisphere, most of which is water.

(2.) As the ocean furnishes most of the vapour of the atmosphere, the condensation of vapour into rain upon the land is greatest near the coast-line. The sea-board of a country may be rainy while the interior is comparatively dry.

(3.) The fall of rain is greatly affected by the form of the surface of the land. Mountains act as condensers (Art. 23), and are consequently much wetter than plains.

(4.) Places which lie in the path of any of the regular air-currents are wet when they cool the current, and dry when they warm it. Hence winds blowing towards the equator, since they come into warmer latitudes, are not usually wet winds; but when they blow towards the poles, they reach colder latitudes, and are chilled and therefore rainy.

30. Some of these laws are well illustrated in the British Islands, where the rains are chiefly brought by the south-westerly winds which have come across the Atlantic. The coast-line facing that ocean is more rainy than the east side looking to the narrow North Sea. In the former part of the country away from the hills, the amount of rain which falls in a year would, if collected together, have a depth of from thirty to forty-five inches. On the east side, however, the average annual rainfall does not exceed twenty to twenty-eight inches. Where the western coast happens to be mountainous, an excess of rain falls; hence the wetness of the climate along the north-west coast of Scotland and in the lake district of England, where the annual rainfall ranges from eighty to one hundred and fifty, and sometimes even more than two hundred inches.

31. Great differences are observed at different points of the earth's surface in the annual amount of rainfall. Between the tropics, where the enormous evaporation

which there takes place raises a constant stream of vapour into the atmosphere, rain is heavy and frequent, so much so that this belt of the earth's surface is known as the Zone of Constant Precipitation (Lesson XI., Arts. 10, 11). When in this rainy region a high mass of land lies in the path of the warm moist air-currents, the rainfall is still further increased. Thus in India the range of the Khasi hills stretches across the course taken by the winds called the south-west monsoons, which bring up their burden of warm vapour from the Bay of Bengal (Lesson XI., Art. 34). The result is that the winds as they slant up the hills into the higher and cooler air have their moisture at once precipitated as rain, of which as much as from 500 to 600 inches fall there in the year.

32. Again, a tract of country situated behind a high mass of land to which vapour-laden winds blow, may have little or no rain. Thus in India, while the range of the Western Ghats, lying in the pathway of the warm moist monsoon from the Indian Ocean, intercepts a heavy rainfall, amounting on the tops of the range to 260 inches annually, the country on the east or lee side receives comparatively little rain. Puna for instance, lying at the foot of the hills, has a yearly rainfall of only 26½ inches.

33. In South America the high chain of the Andes takes out the last of the moisture which blows from the east across the Continent, and the winds descend upon Peru so dry that rain is almost unknown there. Another and much greater rainless tract is seen in the desert regions from North Africa across Arabia and far into the heart of Asia. In these parts of the earth's surface, the dry, sandy soil is raised to an intense heat during the day. There is little or no water to evaporate. The hot, dry air ascends, but the winds which blow in upon the deserts cannot deposit any moisture, for instead of being

cooled, they are heated and driven up in the ascending currents.

34. In some countries (Lesson XI., Art. 21) the wind blows for part of the year in one direction, and for the rest in the opposite direction. These periodical winds are generally accompanied with rain when they come from warm to colder regions, and with dry weather when they come from a cooler to a warmer climate. In such regions there are consequently rainy seasons and dry seasons. During June and July, for example, the southerly wind, referred to in Art. 31, brings the "rains" that refresh the surface of India after the scorching heat of April and May. During November, December, and January, on the other hand, a cool dry wind gently streams down from the northern mountains into the plains of Hindostan, and brings calm dry and settled weather. In north-western Europe, and generally in those parts of the earth which lie in temperate and arctic regions, great irregularity of rainfall occurs. At the end of autumn the weather begins to get more rainy, and continues so during the winter until spring. But long-continued and heavy rains occur also sometimes in summer.

35. Rain falls to the earth as water nearly pure. It is, indeed, natural distilled water. But nevertheless it seems to be never quite pure, and sometimes it contains a noticeable amount of impurities. These have been already referred to in Lesson V. Thus it absorbs a little air, with carbonic acid, other acids, and some gases and vapours which are present in minute quantity in the air. These it brings down to the earth, together with floating particles suspended in vast quantities in the air. The decay of animals and plants furnishes a large supply of infinitesimal particles of organic matter, while over and above these, millions of microscopic living organisms

float all through the lower parts of the atmosphere. In ordinary good air these various impurities are of course in exceedingly small proportions. They are least abundant in the clear air of the mountains, and most in the bad air of towns. Now the rain, as it were, washes these out of the air, and thus purifies it, and makes it healthier, while at the same time it carries down into the soil substances such as ammonia, which are diffused in the air, and are of service to the growth of the plants. So that in addition to all the benefits which rain confers upon us by moistening and fertilizing the soil, filling the rivers, replenishing the springs, and thus keeping the face of nature fresh and green, it cleanses for us the very air which we breathe.

36. Snow.—When from any cause, and in any place whether on the land, on the sea, or in the air, water is cooled down to 32° Fahr., it usually no longer remains in the liquid form, but passes into ice (Art. 13).

Ice is formed in the air by the freezing of the particles which are condensed from the water-vapour. It may descend to the ground as Snow, Sleet, or Hail, according to the circumstances in which it is formed or the condition of the successive layers of air through which it has to pass in its descent. Since the temperature falls in proportion as the air recedes from the surface of the earth, the freezing-point will be reached at no great height above the ground. We might, indeed, imagine a line passing through the air above our heads, from pole to pole, and marking the isotherm of 32° , that is, the limit above which the water-vapour must condense into ice, and below which it will condense into water. Such a line would be liable to oscillate up and down, according to the latitude, the season of the year, and the varying currents of the air. In England, for instance, it would descend to the very ground in winter when the ponds and streams are crusted with ice during the cold period, called a frost,

while in summer it would be a mile and a half above our heads. In India it would be three miles overhead.

37. Let us imagine, then, such a variable line or limit in the air, and consider the forms in which the frozen moisture exists there, or descends thence to the ground. It is highly probable that many of the delicate white cloudlets which we see far up in the air in summer are made of fine snow. At a lower level, flakes of snow are formed and fall earthwards, though from their feathery forms and lightness they are liable to be driven about by every feeble breath of wind.

38. If you look at a snow-flake formed in calm air, you find that it has a regular structure, that in fact it is built up of minute needles or crystals of ice symmetrically arranged. These are grouped into a star with six rays,



FIG. 10.—Snow-flakes.

each of which has a feathery surface, from the abundance of minute crystals of ice ranged along its sides. The accompanying woodcut represents different varieties of snow-flakes. But it will be seen that they are only modifications of the same six-rayed star. The rays strike off from each other at an angle of 60° , and no matter how complicated the form of the snow-flake may be, you will always notice that this angle is maintained between all the rays. All ice, even in the solid sheets which form on rivers and lakes, during the winter of cold climates, is built up of particles which tend to range themselves in hexagonal crystals, although it is only in the snow-flake that this structure is usually seen.

39. Snow is white, but if one of the flakes be taken by itself it will be seen to be a little crystal or group of crystals of transparent ice, glittering with the prismatic colours. So that the white hue of a mass of snow arises from a combination of all these colours reflected from the immense number of minute surfaces of ice. In the same way a dish of salt appears white, and yet any one of the little crystals of which it is made up will be found to be transparent and colourless.

40. When the air is intensely cold, that is, much below the freezing-point, the snow which happens to fall appears in the form of minute flakes like a white powder. The largest flakes fall when the temperature is nearly at the freezing-point. The heaviest snow-falls do not take place during severe frosts, but before or after them. This arises from the fact that in proportion as the air gets colder it loses the capacity to retain watery vapour, so that when it becomes intensely cold it grows also comparatively dry.

41. Over by much the larger part of the globe snow never falls. It can only appear when the isothermal limit of 32° , spoken of in Art. 36, descends to or near to the surface of the earth: it falls, therefore, in those countries where the temperature sinks to 32° . Again, wherever the land rises above that limit, the moisture which falls on it comes down as snow. Thus the Himalaya mountains, though they stand in one of the hottest quarters of the globe, are yet so lofty that their upper portions tower far into the cold upper air, and are therefore covered with snow. On the south side of that high range the lower limit of the snow descends to 16,200 feet above the sea, or about 1,200 feet lower than on the north side,¹ because the cold mountain slopes condense into

¹ These figures are those given by the brothers Schlagintweit: the usual heights assigned are 19,560 for the snow-line on the north side, and 15,500 on the south side.

snow the moisture which is blown from the Indian Ocean and allow the air to pass comparatively dry over to the north side, and because the dry air from the heated plains of Thibet evaporates the snow on the north side.

42. The snow-line, or limit of perpetual snow, is the name given to a line below which the summer heat suffices to melt all the snow, but above which more snow falls than the warmth of the summer months can thaw (Fig. 11.) We may picture it to ourselves as a great invisible arch, of which the centre rises high over the equatorial regions while the extremities come down to the sea-level within the Arctic and Antarctic circles. Under the centre of

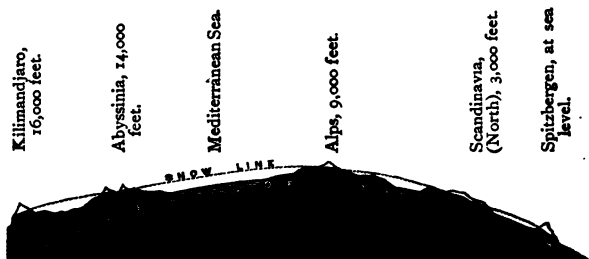


FIG. 11.—Position and height of the Snow-line between Equatorial Africa and the North Polar Seas.

this arch the temperature is so high that snow is only seen on the loftiest mountains, about a height of 15,000 to 20,000 feet over the sea-level, while towards the north and south where the temperature is greatly lower, the snow remains not wholly melted, even down to the margin of the sea.

43. Snow is of great use in winter, as it protects vegetation from being nipped by severe frost. Being a bad conductor of heat, it keeps the soil and plants on which it lies from parting readily with their warmth.

During a frost, therefore, the soil and the plants under a few inches of snow will be found soft and uninjured, while on those places where the snow has been blown away the soil is frozen stiff, sometimes to a depth of eighteen inches.

44. When snow accumulates above the snow-line it gets pressed into ice, and creeps down into the valleys in the form of Glaciers. But this further stage in the history of snow will be considered in a later Lesson when we come to look at the circulation of water over the land.

45. **Sleet.**—When snow is much driven about by the wind its delicate fretwork of crystals is greatly broken. If this takes place when the temperature is rising, or if the driving snow falls through a warmer layer of air, it begins to melt, and in this half-melted state reaches the ground as sleet.

46. **Hail** is the name given to pellets of snow and pieces or concretions of ice which fall from clouds. Usually hail-stones are small, white, rounded, conical, or irregular. Sometimes, though rarely, they assume more definite crystalline shapes. They are occasionally as large as eggs, and sometimes, when several of these come together in their descent through the air, they freeze into large irregular lumps of ice, and in that form reach the ground. Hail is more frequent in summer than in winter, and in hot than in cool weather. It is generally supposed to be caused by the contact of cold currents of air with warm wet ones, though we do not yet quite know how this change takes place. Hail frequently accompanies thunderstorms, and is thus connected with electrical changes in the atmosphere.

47. Hailstorms are sometimes remarkably destructive. When the fragments which fall are of large size they break branches from trees, destroy growing crops, maim and kill cattle and human beings, and injure

buildings ; the course of one of these storms can thus be traced across a country by the devastation which it leaves behind it.

LESSON XI.—THE MOVEMENTS OF THE AIR.

1. How rarely does the air seem to be perfectly motionless ! Even when we say that “not a leaf is stirring,” that “the cottage-smoke mounts up straight,” or that “the clouds are at rest,” we can find, if we look for it, proof enough that the air is not so stagnant as it seems, but is ever moving and mixing. Usually the motion is easily recognised, sometimes in a mere light air or in a gentle breeze, or a strong wind, or a boisterous gale, or, it may be, in a destructive tempest.

2. Why should the air be so restless ? We have already found the answer to this question : because, owing to the unequal heating of the earth’s surface by the sun, and the ever-varying amount of water-vapour poured into or withdrawn from the air, the density or pressure of the atmosphere never remains long stationary at any one place. All movements of the air arise out of differences of pressure. The law governing the direction of these movements may be stated thus, *air always flows in spirally from areas of high pressure into areas of low pressure.* That this must be their direction will be apparent if we reflect that low pressure indicates a deficiency, and high pressure a surplus of atmosphere. The column of air is heavier in the latter case than in the former. Consequently, obeying the universal law of gravitation, the heavier column must necessarily flow out at the base to supply the deficiency in the lighter one. The air does not flow from all sides straight into a low-pressure area.

It circles round it, but always getting nearer to the centre, until it is drawn up and passes away into higher regions of the atmosphere. This inward circling of the air is termed a *cyclone* or cyclonic movement; while the outward flowing from a region of high pressure is called an *anticyclone* or anticyclonic movement. The rate of motion of the air is usually much higher in the former than in the latter. Storms of wind and rain are associated with cyclones; while calm air or light breezes accompany anticyclones.

3. We are accustomed to think only of the horizontal movement of the wind along the surface of the earth. But to form a proper idea of the circulation of the air, we must look upon that horizontal movement as only one part of a much wider system of change. Each cyclone or area of low pressure has at its centre an ascending current, and each anticyclone or area of high pressure has a descending current of air, and it is the position and extent of these upward and downward currents which govern the direction of the winds at the surface.

4. The rate of motion of the surface-currents or winds must evidently be determined in each case mainly by the amount of difference in the pressure and the distance between the centres of the two areas of higher and lower pressure. The greater that difference and the shorter the distance within which it occurs the more rapid will be the flow of air. This is referred to again in connection with storms in Arts. 24—29.

5. It will occur to you, then, that if there be this constant relation between atmospheric pressure and atmospheric movement, it is only needful to know how one of these is distributed over the globe to have a good understanding of what must be the distribution of the other. This is really true. If we know the average readings of the barometer over a continent or over the whole surface of the globe

for any given period, we can readily tell what must be the direction of the chief aërial movements at any particular place. Plates II., III. and IV., illustrate this relation between the distribution of atmospheric pressure over the globe and the general circulation of the atmosphere.

6. The principal causes of differences in atmospheric pressure were stated in Lesson VIII., Art. 18, to be temperature and water-vapour. It is not always possible to say how far any particular movement of the air is due to the one or the other of these causes, or to both combined. Directly or indirectly, indeed, everything might be referred to temperature, for evidently the process of evaporation on which the addition of water-vapour to the air depends is really regulated by temperature, being greatest when the temperature is high, and least when it is low. But it is useful to distinguish between the effect produced by the mere heating and cooling of the air and that due to the changes in the amount and distribution of its vapour.

7. Let us take some simple illustrations of the influence of temperature in producing movements in the air. A fire, whether in a room or out of doors, furnishes an excellent example on a small scale of what takes place in nature. When a fire is kindled in a grate the air overlying it is heated, and ascends; but at the same time there is an in-draught of air at the bottom of the grate. A circulation is established. The air from all parts of the room and from outside, through the crevices of the doors and windows, is drawn towards the fire, warmed, and driven up the chimney. If in any way this free circulation is interrupted, the fire does not burn well, while if it is stopped, the fire goes out. Again, when a house or any large building, or a prairie or forest, takes fire, the heat may be so great as to cause a very rapid ascent of a considerable mass of air above the burning materials. This

is necessarily accompanied by a flow of air inwards from all sides along the ground towards the fire as a centre, and this in-draught may be sometimes such as to amount even to a violent wind. Rushing round towards the blazing centre, this wind feeds the flames, and, as it gets heated, ascends with great force, carrying clouds of smoke, sparks, and even, it may be, large fragments of burning wood along with it.

③. The heat of the sun is not concentrated upon a limited spot like a burning building or even a wide forest on fire. But the heating of a mass of land during the day produces effects quite similar in kind to those we have just been considering, wherever the heated land lies near some other area (of water, for example) which has not been heated so much.

Nowhere can this be seen to more advantage than in the maritime districts of countries where the days are warm and the nights cool. During the day the surface of land becomes much warmer under the sun's rays than the sea does, and the air resting on that surface becomes hotter than that resting upon the sea. The result of this difference of temperature (and therefore of pressure) is to set in motion a light breeze, which, increasing in force during the day, moves from the sea to the land, in order to take the place of the hot air which is always streaming up from the heated surface of the land. This is a sea-breeze. It dies away in the evening. When the nightly radiation begins, however, the land parts with its heat more readily than the sea does, and consequently cools the air more. The air over the land, being denser than that over the sea, moves seaward to take the place of the ascending current from the sea, and hence arises a land-breeze. This movement in turn increases in force, as the difference of temperature between sea and land is augmented by the continued radiation, till it may rise to

a good stiff breeze, which, however, dies away before morning, when the sun once more begins to take effect upon the land, and the temperature of land and sea again becomes uniform.

In mountainous countries where the higher ground rises far up into the colder layers of the atmosphere, another beautiful illustration of these changes of movement in the air may be watched. During the day the air, warmed on the mountain-sides, ascends, and a breeze blows up the valleys towards the heights. During the night the cold heavy air on the mountains flows down as a cool breeze into the valleys.

9. As regards the water-vapour in the atmosphere and its effects, let us consider where these should be looked for in their greatest intensity. Evidently they must be most marked where the temperature is greatest. Hence the broad belt of low atmospheric pressure between the tropics (Lesson VIII., Art. 19), should be the part of the earth's surface where this feature in the motions of the air is best studied. There the sun's heat is greatest and the evaporation greatest. The tropical belt, indeed, with its wide ocean surfaces, may be looked upon as the great evaporating cauldron of the globe, whence comes most of the moisture which is distributed by the winds as rain and snow. It is necessarily a region of low pressure, and therefore air, from north to south, must be constantly pouring into it. Were there no interrupting masses of land the regularity of this movement would be continuous all round the globe. A constant wind would blow towards the equator from either side, and between the two opposite currents there would be a zone in which the currents would meet and ascend as an up-streaming mass of air along the centre of the tract of low pressure. The position of this zone would vary according to the position of the sun. In July it would be

about the position of the Tropic of Cancer. Thence, following the sun in his southward passage, it would travel to the line of the Equator in October, and to the Tropic of Capricorn in January.

By the present arrangement of land and water, however, this regularity is greatly interfered with. It is best shown in the wide Pacific Ocean, where there is least land. Also, it is more distinct in January than in July, because the belt of low pressure then lies nearly along the line of the equator with more sea in its track than in July, when it goes 10° or 15° north of the equator and crosses the great mass of land forming the continents of Asia and Africa.

10. Putting aside for the present these irregularities, let us think of the air round the tropical belt of the globe as constantly receiving an enormous volume of water-vapour from the oceans. From this great influx of vapour, as well as from the high temperature along that belt, there is a continued ascent of heated and humid air into the higher regions of the atmosphere, and a consequent in-draught from north and south. It is called the Belt of Equatorial Calms—a somewhat misleading name, because, although there are no constant winds, the air is in ceaseless disturbance as it gets heated and streams upward.

11. Constant Winds and Aerial Currents.—The hot moist air which mounts up from the equatorial belt expands and cools as it ascends, and in so doing parts with much of its moisture, which condenses and falls as rain. Hence the belt is distinguished not only by its great heat, but, in many parts, by its almost continual rains and thunderstorms, so as to be spoken of as the Zone of Constant Precipitation (Lesson X., Art. 31). Reaching higher parts of the atmosphere the equatorial hot air divides, one part flowing northward, the other southward, and each

as an upper current taking a direction exactly opposite to that of the in-draught at the surface. These upper currents must be many thousand feet above the sea at the equatorial belt. They travel pole-wards until reaching the belt of high pressure they descend and at last reach the surface in temperate latitudes. Thence the air streams partly back along the surface to the equator and partly towards the pole.

12. Here, then, we have the fundamental aërial circulation of the globe—a great system of surface currents continually streaming out of the bands of high atmospheric pressure towards the belt of low pressure round the equator on the one side and towards the poles on the other, and of upper currents constantly flowing away from the low pressure areas—a system set in motion by the greater warming of the equatorial regions by the sun, and the greater moisture and consequent lower pressure of the atmosphere there than in the regions on either side.

13. It might be supposed that these currents flowing in uninterrupted circulation should have a direct north and south course. And so they would, were it not that the earth, instead of being at rest, is always rotating on its axis. Owing to the motion of rotation, an object at the equator is carried along with a far higher speed than an object nearer the pole, just as the rim of a wheel moves through much more space in the same interval of time than the parts near the axle. Hence, when a current of air travels away from the equator northward or southward, it moves from a region where the velocity of the surface is greater to one where it is less. At the equator the air, having the same rate of rotation as the rest of the surface, is borne along with the rest of that surface, and if it blows eastward or westward, rotation does not affect its direction. But when it blows northward or southward from the equatorial tracts,

it carries with it part of its equatorial rapidity of rotation, and goes a little faster than the surfaces over which it successively moves. Now, as the earth rotates from west to east, all the air which is trying to flow straight from the equator towards the pole is bent slightly to the east. On the other hand, as the air which travels from higher latitudes towards the equator is always getting into regions where the speed of rotation increases, it seems to lag behind, and instead of flowing straight north and south, it is deflected towards the west.

14. The lower or surface currents, turned more and more westwards as they advance, become on the north side of the equatorial belt the North-east Trade-winds, and on the south side the South-east Trade-winds. These are constant winds. They always stream towards the equator because of the lower atmospheric pressure there. They received the name of "Trade-winds" in the early days of navigation, when it was found that they blew so steadily in one direction that they could always be reckoned on for the purposes of commerce.

The Trade-winds are chiefly felt between each tropic and the equator, but they begin in some places considerably beyond the tropics, their constancy being better exhibited over the oceans than over the land, seeing that the more marked heating and cooling of the land by day and night tend to disturb their regularity. Beyond the limits of the Trade-wind region, along the belt of high atmospheric pressure, there is in each hemisphere a belt of calms and variable winds, that on the north side being called the Calms of Cancer, that on the south side the Calms of Capricorn.

15. Out of each high pressure belt the winds blow towards the equator on the one side and towards the pole on the other. As the direction of the Trade-winds is deflected to the west by the influence of the earth's

rotation, so the winds which emerge from the polar side of the tropical calms acquire an easterly direction. In the northern hemisphere therefore they come from the south-west. In Britain, and the west of Europe, which lie considerably to the north of the Calms of Cancer, the prevalent winds are the familiar south-west winds. In these countries it is common to find that the large towns grow towards the west, and that the term "west-end" usually means the quarter where the best streets or dwelling-houses are situated. The reason is evidently to be sought in the direction of the prevalent winds, for the west and south-west are the quarters which escape most of the town-smoke that is blown eastward.

16. Much interesting information may often be gathered respecting the movements of the upper currents when still high in the atmosphere, as well as of other aerial currents, by watching the clouds. In Britain, for instance, fine light fleecy clouds may be noticed at a great height, sometimes with other heavy clouds lying below them. They may be seen to move and change their shape, sometimes affording by these movements valuable prognostications of coming weather.

17. But the fact that there are upper currents in the atmosphere is now and then brought home to us still more strikingly. Volcanoes have been already referred to as mountains from which melted rock, steam, and gas are given out, and from which fine dust is sometimes thrown for an enormous height up into the air. With such force is this dust or ash ejected, that it sometimes enters a high aerial current blowing steadily and strongly in one direction. It is then borne along to vast distances, and descends perhaps to the earth in regions where nobody ever saw such a thing as a volcano. In this way the inhabitants of the northern parts of the British Islands have more than once been visited with showers of dust which came from

Iceland. In the year 1783, during a great eruption of Skaptár Jökul, one of the volcanoes of that island, the fine impalpable dust fell in such quantities between the Orkney and Shetland Islands during north-west winds, that vessels sailing there required to have the deposit shovelled off their decks every morning. In that season, too, the crops along the north coast of Caithness proved a failure from the quantity of volcanic material which fell on the ground. The inhabitants still speak of the year of "the ashie." The distance from Iceland to the coast of Caithness is upwards of 600 miles.

18. Instances are on record where the erupted ashes have been borne over a much greater space. In the year 1835, Coseguina, a volcano in Guatemala, broke out in eruption, and though the regular east wind was then blowing, the fine ashes were carried eastward by an upper current, and fell on the island of Jamaica, 800 miles from the point of emission. They had been four days on the journey, and their rate of progress must have been 200 miles a day. Again, in 1815, during a calamitous volcanic outbreak in the island of Sumbawa, lying to the east of Java, the ash was carried for 800 miles to the east, as far as the islands of Amboyna and Banda, although the south-east wind was then at its height.

19. In countries bordering the Mediterranean, and even as far west as the Cape de Verde and Canary Islands, the air is occasionally filled with a peculiar reddish, or brown dust, which sometimes falls in such quantities as to cover the decks of ships at sea, though far out of sight of land. When rain falls at the same time it brings down this dust with it, and having then a peculiar colour, is known as "blood-rain." It seems that in most cases this "red-fog," "sea-dust," or "sirocco-dust," as it is called, has come from the hot deserts of Africa, whence, swept up by whirlwinds from the burning soil in great

clouds, it has been borne away by the overflow of air at the top of the hot ascending column, and after being carried for hundreds of miles by the upper currents, has at last descended to the surface again.

20. Clouds, ashes, and dust, therefore, sometimes give important evidence as to the direction in which the upper aërial currents are moving, and help to prove how regular and constant is the circulation of the atmosphere.

21. Seasonal or Periodic Winds.—When the two charts of the world in Plates II. and III. are compared it may be seen that the larger masses of land in the northern hemisphere interfere a good deal with that regular distribution which, as shown by the southern hemisphere, a broad unbroken expanse of ocean favours. In January, for instance, the high and cold table-lands of Central Asia become the centre of a vast area over which the pressure of the air is high. Consequently from that elevated region the wind issues on all sides. In China and Japan it appears as a north-west wind. In Hindostan it comes from the north-east. In the Mediterranean it blows from the east and south-east. But in July matters are reversed, for then the centre of Asia, heated by the hot summer sun, becomes part of a vast region of low-pressure, which includes the north-eastern half of Africa and the east of Europe. Into that enormous basin the air pours from every side. Along the coasts of Siberia and Scandinavia it comes from the north. From China, round the south of the continent to the Red Sea, it comes from the Indian Ocean, that is, from south-east, south, or south-west. Across Europe it flows from the westward. Hence, according to the position of any place with reference to the larger masses of sea and land the direction of its winds may be estimated.

22. On the shores of the Indian Ocean the summer and winter winds are known as Monsoons—an Arabic

word signifying any part or season of the year—but now generally applied to all winds which have a markedly seasonal character. Since the air is drawn in towards the heart of Asia in summer and comes out from that centre in winter, the direction of the monsoon at any place depends upon geographical position. In India the winter wind is the N.E. Monsoon, which corresponds to the N.E. Trades of the North Atlantic and North Pacific Oceans ; the summer wind is the S.W. Monsoon, which is a complete reversal of the natural course of the Trade wind owing to the enormous in-draught caused by the low summer pressure over Asia. On the Chinese coast the winter wind is a N.W. monsoon, and the summer wind a S.E. monsoon. Similar but not quite so strongly contrasted monsoons occur in North America. In the Southern States, for instance, the winter wind comes from the north-east, the summer wind from the south-west.

23. Local Winds.—Many winds, often of a destructive character, occur in different countries or in different districts of the same country to which local names are given. When they come from tracts where the pressure is high and the temperature low to where the pressure is lower and the temperature higher, they are felt as cold blasts, whereby the humidity of the air in the low pressure area is condensed into torrents of rain. When around hot desert regions like those of Africa, Arabia, or the interior of Australia, a low atmospheric pressure occurs, its effect sometimes is to draw in towards it the hot air lying over these burning sands, which in the countries where it blows is extremely unhealthy. In Italy it is known as the *sirocco*—a hot moist wind which raises a haze in the air, and produces a sensation of extreme languor both in man and beast. In Spain, where it receives the name of the *solano*, it sometimes comes across the narrow part of the Mediterranean laden with fine hot dust from the vast

African deserts. In Africa and Arabia it appears as the dreaded *Simoom*—a hot suffocating wind which sometimes rushes across the desert with such violence as to raise clouds of sand, and sweep them in whirling masses for many miles. It thus heaps up vast mounds of sand under which caravans of travellers may be completely buried. One of the armies of Cambyses, 50,000 in number, is said to have been engulfed in the sand when on its way to attack the oasis and temple of Jupiter Ammon. Again on the coast of Guinea during December, January, and February, a hot wind, called the *Harmattan*, blows from the interior out to sea. The north-west provinces of India have likewise their hot-winds, which sometimes produce violent whirlwinds sweeping up the dust and carrying it in tall whirling columns into the upper air, whence it gradually finds its way to the earth again.

24. Storms.—But besides the winds and currents already noticed, most of which blow continuously for weeks or months, there are sudden and violent commotions in the atmosphere often very disastrous in their effects. These may be classed under the general name of storms. Let us see whether they can be accounted for by the general principle of aërial circulation which holds good for the constant and periodic winds.

25. A violent storm often bursts upon a country with very little warning. The wind rises rapidly into a gale and continues to rush along at a rate of often 70 or 80, sometimes in gusts even of 120 and 150 miles an hour. After some hours it slackens its speed and may even die away as rapidly as it rose. But again it springs up from a different or even from the opposite quarter, and possibly equals if it does not exceed its first fury and destructiveness. All parts of a country have not the storm simultaneously. In Europe, for instance, it is usual for the

storms to come from the westward. Britain and the western shores of France and Portugal catch the first fury of the tempest, which then travels eastward or north-eastward, and not unfrequently dies away in Russia. The rate at which the storm moves from place to place does not exceed on an average from fifteen to thirty miles an hour—a much more leisurely rate than the rapid and destructive rush of the wind would have led any one to expect.

26. Suppose that such a storm has passed over Europe, and that, after it is gone, we obtain information as to the readings of the barometer over all the countries traversed by the storm as well as those around its track. We find that the centre of the storm has been an area of very low atmospheric pressure, 600 miles or more in breadth, and that the pressure has been much higher immediately outside of that area. The great law of atmospheric movement, therefore, holds good here also. The air has moved as a cyclone inward from surrounding regions of high pressure upon a central area of low pressure.

27. Observations of this kind have proved, that the greater the amount of difference between the low pressure in the centre of the storm and that of the high pressure outside, and the nearer the extremes of pressure lie to each other, the more violent is the wind. Suppose, for example, that while the barometer stood at 30 inches at one side of a country, it should rapidly fall to 28 inches at some district on the other side, three hundred miles away : so great a difference in so short a space would almost certainly be accompanied by a violent gale. The air in such circumstances rushes in with an in-moving spiral and ascending motion. It whirls round the circumference of the storm-ring, always getting further in, until meeting in the middle from all sides and forming in consequence a calm, it is carried upward in the central vortex as a

vast up-streaming current, which overflows at the top and passes off into other regions.

28. Hence it is evident why, though the wind rushes so furiously, the whole body of the storm does not travel so fast in many cases as an ordinary passenger train. There are two motions in the storm, that of the whirling air as it is borne inward, and that of the whole rotating mass or storm. When the storm-ring bursts upon a place the wind may be blowing from the south-east. It dies down when the centre of the storm reaches the place, but rises again from the opposite quarter when the remaining side of the storm-ring advances, and continues until the whole whirling mass of air has gone by.

29. A good illustration of this vorticose movement of a storm is afforded by the little dust-whirlwinds often seen upon a dusty road in dry weather. The air in each of these, as shown by the movements of the dust, takes a rapid inflowing spiral course, drawing in air on all sides below and whirling it rapidly round and upward, till it expands at the top and flows over into the surrounding air.

30. Office of the Winds.—There are two kinds of work done by the winds and currents of the atmosphere (1) The distribution of temperature ; and (2) the distribution of moisture. This work has been referred to already, but may again be stated in a summary way here.

31. (1) When a wind blows from a warm or mild quarter it raises the temperature of the places to which it comes. The west and south-west winds of Britain, for instance, warmed by the heated waters of the North Atlantic, keep the air of that part of Europe much milder than from its latitude it ought to be. When on the other hand a wind blows from a cold to a warm region it lowers the temperature as it moves along. Thus, owing to the vast expanse of cold land in Asia during winter and the

high pressure there, the outflowing winds are cold, and blowing as they do westward into Europe, they make the weather there in winter and spring colder and drier than when the west winds blow. A careful study of the charts in Plates II., III., and IV. will show this more clearly than a description in words can do. By noting the areas of low and high pressure we learn in which main direction the wind must blow in any part of the globe; then a reference to the isothermal charts will indicate to us the temperature of the high pressure tracts from which the wind comes, and therefore whether the wind in any particular case must be cold or warm.

32. (2) The winds are the great agents by which the moisture of the atmosphere is distributed over the globe. Were it not for their help the condensed vapour would be discharged over the same tracts from which it had been evaporated. And therefore, as by far the most of the vapour rises from the ocean it would fall back again into the ocean, and the land would only get back the comparatively small quantity evaporated from rivers and lakes. But since the winds bear the vapour along with them, and since areas of low atmospheric pressure often pass over continents, the vapour is carried away from the sea and condensed into rain over the land. On the whole, therefore, the wetness or dryness of any place will depend on the direction from which its prevalent winds come (Lesson X., Art. 29-34). If they blow from a wide and warm expanse of sea they will be often laden with moisture and ready to drop it when they are chilled by passing over the land. If on the other hand, they blow from the interior of the continent they will be dry, their heat or cold depending upon the temperature of the region whence they come.

33. In short, winds are wet when they travel from a warm vapour-laden tract to a colder one: because they are

cooled below their dew-point and must let fall the surplus moisture. Winds are dry when they travel from a cold to a warm tract, because, instead of parting with their vapour, they are ready to take up more ; they are dry too when they issue from a hot and dry region, and before they have an opportunity of crossing any sea and greedily drinking up vapour from its surface.

34. When a mass of high land stretches across the track of a warm moist wind, it forces the wind to flow up and over its ridges. The air is thereby expanded and cooled, and as a consequence deluges of rain descend. The most notable instance of this kind is that of the Khasi Hills already mentioned (Lesson X., Art. 31), which front the Bay of Bengal and interpose as a great steep buttress against the advance inland of the warm vapour-laden south-west monsoon. Driven up the slopes, the wind has to discharge its moisture in rain, of which, during the seven months when that monsoon blows, a depth of as much as 500 inches, or about 42 feet, falls to the ground. But, after passing over the hills, the wind, having lost so much of its moisture, is comparatively dry. Hence the rainfall rapidly diminishes northward. As the wind journeys in that direction, however, it is once more forced to ascend into the higher regions of the atmosphere by the giant chain of the Himalaya. Its moisture is further wrung out of it, as rain on the lower slopes, and as snow among the higher ridges. And thus it descends on the north side as a cool dry wind into the wide plains of Thibet.

35. In countries like India, therefore, the dry season and rainy season depend upon the changes of the monsoon and the quarter whence the monsoon blows, whether dry land or vapour-giving sea.

36. In regions where the winds are irregular the rains are found to be the same. In the west of Europe, for

example, the wind shifts continually, but the prevalent winds are from the west and south-west, though during part of the year east winds become frequent. The westerly winds come from the warm Atlantic, laden with vapour, which they discharge plentifully on the western parts of Britain and the Continent, leaving the eastern districts comparatively dry. The easterly winds on the other hand come from the heart of Asia, and are so dry as often to wither up vegetation as much as a hard frost or even a scorching fire would do.

37. There is yet another kind of work performed by the winds, which, though of very limited and local character compared with their other services, deserves notice here. Storms by sweeping over the surface of the land sometimes produce great changes there, uprooting trees and



FIG. 12. - Sand-dunes—ridges of dry sand blown inland off the shore by the wind.

even prostrating large woods, destroying fields and houses and generally carrying ruin with them in their track. But even where the wind does not rush with such fury over the land it sometimes effects remarkable changes there when its prevalent direction is from the sea to the shore. On sandy beaches exposed to winds which usually blow

from the sea, the dried sand is often driven inland, where it forms mounds and ridges called sand-dunes, sometimes 60 or 100 feet high, fronting the beach for, it may be, many miles. A small streamlet, flowing behind these ridges may arrest their advance by carrying away the loose sand as fast as it creeps on. But where no such check exists, the sand may gradually extend inland, and completely bury the cultivated soil over whole farms or parishes.

38. Many examples of these effects are to be found along the coasts of the British Islands. Within the last few hundred years, thousands of acres of valuable land have been overwhelmed by the drifting of sand from the sea. Along the shores of the Moray Firth and of Aberdeenshire, for example, several parishes have been wholly or in great part buried. On the coasts of Norfolk and Suffolk, villages and thousands of acres of land have been covered with blown sand during the last two centuries, and on the Cornish coast similar inroads have taken place. The shores of the Bay of Biscay furnish still more striking instances of the power of the winds to alter the surface of a district. The dunes which are there heaped up by the westerly gales, march inland at a rate of 60 or 70 feet in a year. No barrier, natural or artificial, is able to withstand their progress. Fields, woods, and villages are buried in succession. Nor is this all. The sand ridges interrupt the drainage from the interior, and the water collects among and in front of the ridges. Ponds and lakes are formed, which, unable to find an exit, are driven inland along with the sand barriers which dam them up. Large tracts are thus first inundated with water, and then finally overwhelmed under the advancing sand. Roman Roads, and many villages which existed in the Middle Ages, have disappeared. And the same destruction is going on still.

39. It is not always along the sea-margin that these effects of the action of wind are to be traced. In the interior of many countries there are wide tracts of barren sand or Deserts (Art. 23), over which the wind raises the sand into ridges, as it does along the shore. Such, for instance, are the great desert of Sahara and much of the interior of Arabia.

CHAPTER III.

THE SEA.

LESSON XII.—THE GREAT SEA-BASINS.

1. From the outer envelope of air which encloses the earth we now pass to the underlying envelope of water called the Sea. At the outset some obvious differences between these two coverings may be noticed. For example, while the atmosphere completely wraps round the whole planet, rising to a height of many miles above its general surface, the water envelope is pierced in many places by masses of the underlying solid part of the earth, which rise above it to form land. As the result of observations in all parts of the world it has been ascertained that the sea covers not quite three quarters, and the land a little more than one quarter, of the entire surface of the earth ; or more accurately, the sea spreads over 144,712,000 square miles, and the land over 52,000,000 square miles, the total area of the surface of the globe being computed at 196,712,000 square miles.

2. Again, we know nothing about the upper surface of the atmosphere, and cannot say precisely how far it lies above us, but the surface of the sea forms a great plain, and the line between it and the air is sharply defined. Though we speak of the sea as forming a plain, we know

this plain to be really curved, and that from its wide extent and its freedom from inequalities, it shows the curvature of the earth's surface better than can be seen on land. (Lesson I. 2.) The amount of the curvature is about eight inches in a statute mile; that is, an object eight inches high above the sea-level sinks out of sight when looked at from the same level at a distance of more than a mile. The line of meeting between the sky and the surface of the earth is termed the horizon. Its distance from us evidently depends upon the elevation at which we may happen to stand. Thus, at the sea-shore with our eyes exactly six feet above the sea-level our horizon out to sea is three miles off. If we ascend so that our eyes are about ten feet and a half above the sea-level, our horizon is extended to four miles. If we climb to some adjoining height, say to a lighthouse top, about ninety-six feet above the sea, our horizon is increased to a distance of twelve miles.¹

3. Another evident contrast between the air and the sea lies in the fact that while every inhabitant of the earth is familiar with the one, only a comparatively small part of mankind has ever seen the other. Even in a small and populous island like Britain, a large proportion of the inland population has never been within sight of the sea. On the continents the proportion must be much greater, for, except the people dwelling along the sea-margin, the great mass of the inhabitants, having little or no communication with the coasts, have no acquaintance with any larger sheet of water than their own native river or lake.

4. It is hardly possible that any one who has not seen the sea can realize what it is from descriptions in books.

¹ In estimating the extent of the visible horizon, however, when the distance exceeds half a mile, we need to take into account the effect of atmospheric refraction which tends to make distant objects seem higher than they are. The allowance to be made for this effect varies from day to day; it commonly requires a deduction of about one-seventh from the apparent height of an object.

Let us suppose, however, that some intelligent dweller in the inland regions were taken for the first time to the sea-coast, and after recovering from his first impression of wonder and admiration, were to begin to look attentively at those features which would be most likely to attract his attention. He would observe that the solid ground, with which he had been familiar all his life, gives place to a seemingly boundless plain of water, at first sight level and motionless, but soon found to be in a state of perpetual unrest, and answering to all the movements of the air above, heaving or rippling when the air is gently stirred, but rising into waves and foam-crested breakers along the shore when the wind blows strongly. Should he taste some of this blue sparkling water he would find it salt and undrinkable, even though clearer perhaps than his own river at home. Gathering the shells and other remains of the living things of the sea from the sands of the shore, he would find every one of them different from anything he had ever found on the land or in the fresh-water. If he watched by the shore from day to day, he would notice that twice every day the water advances slowly and as slowly retreats, and that this regular movement takes place whether the water be smooth or rough. Were he to set sail upon that seemingly boundless plain of water, he would watch the land behind him gradually sinking as it were into the water, till at last its highest point had disappeared, and the same long level line of meeting between sea and sky would then sweep around him on every side. With no land in sight and no other vessel perhaps to be descried; with only the sky overhead and the heaving water beneath and around, he would learn better than from any map or description the vastness and solitude of the great deep. And yet from the deck, or even from the mast-head of his ship, only a comparatively small patch of the sea could be seen at once. The horizon or sky-

line, which he thinks so immeasurably distant, is only a few miles off. When standing on the sea-level he cannot see an object on the same level as high as himself at a greater distance than about six miles in any direction. And he might sail for weeks together, passing over thousands of miles, with all the time the same limited horizon and the same monotony of sea and sky.

5. The traveller who finds himself for the first time on the open sea is not only impressed by the vastness of its surface, but by a vague notion of its profound depth. He has read of the "unfathomable abysses of ocean," and naturally feels some awe at the thought that they are actually lying beneath him. In recent years, however, Expeditions have been fitted out to measure the depth of the sea in all parts of the globe, and the result has been to show that even the deepest of the so-called abysses probably does not much exceed five miles, while the average depth of the sea may be taken at about half that amount.

6. Had the earth been a perfectly smooth globe like a ball of polished steel, we may suppose that the sea would have completely covered its surface. There would then have been an outer shell of air and an inner shell of water, within which the solid planet itself would have lain. The average depth of such an universal sheet of water as may be inferred from what has been made known by numerous soundings in all parts of the present sea, would have been rather more than a mile. Probably the water envelope never completely wrapped the globe round in this way: at all events, owing to the irregular projections and depressions of the solid part of the earth, the sea is now broken up into separate basins or oceans by intervening masses of land.

7. Looking at a map of the world we may be disposed to ask whether any reason can be given for the shape which the great ocean-basins have taken. We may wonder whether the sea, by the force of its own waves

and currents, could have hollowed out the great depressions in which it now lies. But if we try to follow out this subject we soon find that the sea could have had very little to do with the formation of its basins. It is, as we shall learn in Lesson XVIII., only the upper parts of the sea which eat away the solid part of the earth ; the water in the deep abysses does not even stir the fine mud, which, like dust in a deserted room, slowly settles down upon the bottom. Where the waves wear down the land, the materials which they remove are not destroyed, but are carried away into some still part of the sea, where they accumulate on the sea-floor. So that even if we could suppose it possible that the sea might have scooped its basins out of the solid surface of the globe, it could only remove the *débris* from one part of its bed to another, and, in short, fill up its hollows after it had formed them.

8. It is not the action of the sea, then, but the sinking and uprising of the solid surface of the globe, which has determined the shape and size of the great sea-basins. Consider what was said in Lesson III. about the probable history of the earth. If our planet has slowly cooled from a condition of intense heat, it must have gradually contracted in bulk. We may compare it to an apple which is kept a long while, until it begins to dry up, and consequently to shrink. The apple does not diminish equally all round. Some parts of the skin sink down, others get puckered into wrinkles and foldings, so that in the end the once smooth surface becomes shrivelled and uneven. In the case of the earth, the contraction has likewise been irregular. Some portions have sunk down, others which shrank up soonest have become more and more wrinkled, and rise far above the rest. The depressed parts have formed the basins for the oceans. Those portions of the larger wrinkles which rise above the sea-level form the islands and continents.

9. Until a comparatively few years ago, the abysses of the sea were as unknown as the interior of the earth. Their depth was only guessed at, for the few attempts to measure it had led to great exaggeration. The nature of their floor was equally a matter of conjecture. It was supposed, indeed, that the water at these profound depths lay dead and motionless, that no living thing could possibly exist at the bottom, and that no change went on there save the slow accumulation of fine mud, carried by ocean currents from far-distant land.

10. But much of this uncertainty has now been dispelled. Expeditions to measure the depth and explore the bottom of the sea have been sent to all parts of the world, and have collected such a mass of information that we now know more about the depths of the ocean than about some parts of the continents. Three instruments have been used in these researches : the Sounding-line, the Dredge, and the Thermometer.

11. The Sounding-line consists essentially of a rope or cord with a weight at the end, so arranged that when dropped from the side of a vessel to the sea-bottom the amount of line which runs out from the reel is accurately registered and at once indicates the depth. The weight may have its bottom covered with lard, or other soft material, which, when it strikes against the bottom, will bring up some of the mud, sand, gravel, or other indication of the nature of the ground below. The Dredge consists of an iron frame with a strong net or bag attached to it ; the object of this instrument being to bring up a more considerable quantity of the materials which form the sea-floor, as well as any of the living or dead plants and animals to be found there. The Thermometer has been used to such good purpose that the temperature of the vast Atlantic and Pacific Oceans has

been determined at many points from the surface down to the deepest parts of the bottom.

12. At present, the tract of sea which has been most thoroughly explored is that of the Atlantic Ocean. As shown upon a map of the globe, this Ocean runs as a long and winding belt of water between the New World and the Old. Towards the north it is rather closed in by the convergence of the American and European continents with the islands between them, though a free, open communication is maintained between the Atlantic and the Arctic Ocean. Southward from the equator it gradually widens out and merges eastward into the wide Indian Ocean. Hence, though narrow in comparison with the vast basin of the Pacific, the Atlantic is really longer than that ocean, for there is no such close approach of the coasts of Europe and America as there is between those of America and Asia, which bound the Pacific to the north, and divide it from the cold Polar waters beyond Behring's Strait. It has been estimated that the whole area of the Atlantic Ocean amounts to 35,165,000 square miles, which is nearly a fifth part of the entire surface of the globe.

13. Stretching thus from pole to pole, it crosses all the various zones of temperature which girdle the globe. Its central parts lave the shores of equatorial America and Africa. Beyond these, it reaches the temperate climates of North America and Europe. At either end it enters the icy regions of the Pole.

14. Again, owing to the arrangement of the continents which bound it on the west and east, the Atlantic receives a far larger river-drainage than any other ocean. The map shows that down the whole length of America all the large rivers flow eastward, and therefore fall into the Atlantic—the St. Lawrence, Mississippi, Orinoco, Amazon, and La Plata. In Europe and Africa, if we include

the rivers which enter the Mediterranean and Black Sea, by far the largest proportion of the drainage finds its way into this ocean by such important rivers as the Rhine, Rhone, Danube, Dnieper, Nile, Niger, and Congo. The Atlantic basin is thus the great reservoir into which the largest rivers of the globe discharge their waters.

15. That the floor of the Atlantic cannot be a vast plain like the surface of the water is shown by the islands which here and there rise even from the middle of the ocean. Tracing the positions of these upon the map, we see that there must be several great elevations in the bed of the North Atlantic. One of these has its highest points in the islands of the Azores; another rises into the group of the Bermudas. In the South Atlantic a long submarine ridge must run down the centre of the ocean, finding its way to the surface here and there, as in the islands of St. Paul, St. Helena, Ascension, and others.

16. But we are not left to make such general inferences regarding the shape of the Atlantic floor. From the numerous soundings which have been taken throughout the entire length and breadth of this ocean its form can be laid down with considerable accuracy upon a map. The British Government has fitted out between 1870 and 1872 several vessels, with the express object of exploring the depths of the sea, and in particular of the Atlantic Ocean.

17. As the result of these soundings it has been found that few parts of this ocean have a greater depth than 3,000 fathoms; the deepest sounding was obtained at about 100 miles north from the island of St. Thomas, and gave 3,875 fathoms, or rather less than $4\frac{1}{2}$ miles. This seems to be the greatest depth which has yet been obtained in any part of this ocean by accurate and trustworthy measurement. The wider parts of the Atlantic

have a depth of from 2,000 to 3,000 fathoms, or from about 2 to $3\frac{1}{2}$ miles. It would seem, therefore, that these submarine tracts form a vast plain, rising in some parts into ridges, such as that of which the Azores form the highest elevation.

18. To the west of the British Isles for 230 miles the slope of the ocean-bottom is very gentle, being only six feet in the mile. But beyond that distance the ground descends so rapidly, that in the next 20 miles there is a fall of 9,000 feet, down to the level of the great submarine plain, which stretches for hundreds of miles to the west, with little variety of surface.

19. The Pacific Ocean has not been so diligently examined as the Atlantic. But some recent soundings, taken during the great cruise of the ship *Challenger*, show a still greater depth than those in the latter ocean. One observation, taken in the North Pacific between Japan and Admiralty Island, gave the enormous depth of 4,475 fathoms, or rather more than five miles. Another gave 3,950 fathoms, or about $4\frac{1}{2}$ miles. These are the deepest ocean-abysses which have been accurately measured.

20. But these are exceptional depths. The greater part of the North Pacific, like the North Atlantic Ocean, seems to average from 2,000 to 3,000 fathoms. The vast Pacific basin is traversed by many ridges, as may be seen on the map, from the chains of islands which dot its surface. One vast ridge rises along the northern margin of the basin, and stretches from the American coast through the Aleutian Islands to Japan, thence by the Philippine Islands, New Guinea, and the New Hebrides to New Zealand. In mid-ocean the tops of other vast ridges are indicated by the scattered archipelagoes, such as those of the Caroline, Marshall, Sandwich, Marquesas, and Society Islands. (See Lesson XIX.)

21. The map of the world shows that although the

great water envelope of our planet may, for the sake of convenience, be parcelled out into separate oceans, these are all united into one vast continuous sheet of water. Here and there, however, owing to the way in which the land has been ridged up, portions of the water have been almost separated from the main mass. These are called *mediterranean*, the best example being that which has long been known as the Mediterranean Sea. The Black and Baltic Seas in Europe, Hudson's Bay and the Gulf of Mexico in North America, the Red Sea, Persian Gulf, and Seas of Japan and Okhotsk in Asia, are other illustrations. But sometimes the uprise of the land has taken place in such a way as to cut off completely some outlying parts of the oceans. These are termed *inland seas*. The Caspian Sea and Sea of Aral are the most important instances of this result. The former sea covers a larger space than the British Isles. Its surface is about 85 feet below sea-level, and its greatest depth amounts to nearly 3,000 feet. Its waters are inhabited by seals and other marine animals, while in the tracts of land which now inclose the Caspian and Sea of Aral and separate them from the Black Sea on the one hand, and the Arctic Ocean on the other, beds of dead sea-shells are found. Reference will be made in Lesson XXIII. to the proofs that the land along the coast of Siberia has in comparatively recent times been raised out of the sea. The extent of the upheaval is generally indicated on the map in Plate IX. Before this uprise the Arctic Ocean seems to have extended in a long arm between Europe and Asia as far as the hill-range which is now cut through by the narrow channell of the Bosphorus, but not to have communicated with the present Mediterranean Sea. By the rise of the land towards the north all that part of this vast inlet lying to the south of the parallel of 50° or 52° N. was cut off from the main

ocean. The present abundant salt lakes and marshes, as well as the two large basins of the Caspian and Aral, must be looked upon as the mere shrunk remnants of this old Mediterranean sea. In some yet unexplained way the intervening ridge between the Black and Mediterranean Seas has been trenched, so that the former sea no longer forms part of the Caspian basin, but sends its surplus waters through the Bosphorus and Sea of Marmora into the Mediterranean. There seems also to be less rain now than formerly in the region of the Caspian and Aral, so that these sheets of water are still further shrinking.

LESSON XIII.—THE SALTNESS OF THE SEA.

1. The aqueous vapour of the atmosphere is in great part derived from the evaporation of the surface of the sea, and as the sea does not get sensibly lower in level notwithstanding the enormous volume of moisture which is annually driven off from it into the air, it must receive back again from the rivers which descend to its shores, and from the rain which falls upon its surface, as much water as it supplies. There is thus a ceaseless coming and going of water between the sea, the air, and the land.

2. When the water, however, is examined in different parts of its passage, it shows great varieties. Rain is water condensed from natural distillation, and is nearly pure. Nevertheless it often has a distinct taste, and when evaporated leaves behind, as we have seen in Lesson VI., various salts and particles of dust, which it has taken out of the air. The water of springs and rivers always leaves a still larger proportion of substances after evaporation, but is usually quite fresh and tasteless. But sea-

water is always strongly salt to the taste, and this is true of the sea in all quarters of the globe.

3. Another distinction between the water of the air or of springs and rivers, and that of the sea, is shown in their different weights. A bottle filled with sea-water weighs more than when filled with fresh-water. Perfectly pure water being taken as the unit, sea-water has an average relative weight, or specific gravity, of about 1.026; in other words, if a certain quantity of perfectly pure water weighs 1,000, the same quantity of sea-water will weigh 1,026. The water of the different oceans varies slightly in density. That of the Atlantic, for example, varies more and is generally heavier than that of the Pacific. In the path of the trade-winds in the North Atlantic, where evaporation must be comparatively rapid, some of the heaviest and of course saltiest sea-water occurs, its specific gravity being 1.02781. In the Antarctic Ocean, on the other hand, among the broken ice, the specific gravity sinks to 1.02418.¹

A striking proof of the greater lightness of fresh-water may sometimes be seen after heavy rain has fallen in still weather upon the sea. The rain floats on the surface of the sea, and so does the water brought down by streams from the land. Fresh drinkable water may then be taken from the surface of the salt-water which lies below, until, unless more quickly mingled by wind and waves, the fresh-water slowly diffuses itself downward.

4. What is it that makes the sea-water heavy and salt? To answer this question for yourselves, get, if you can, a small quantity of sea-water. Evaporate it till only a little of it remains, which tastes extremely salt and bitter. Stirring this remaining liquid, so as to mix it completely, put a drop or so upon a piece of thin glass and place the glass under a microscope. Now watch what takes place.

¹ Buchanan, *Proc. Royal Society* (1876), vol. xxiv.

It is one of the simplest and most beautiful experiments you can make. The drop continues to pass off into vapour, but the substances which gave the water its salt and bitter taste cannot do so. They remain behind, and as the water leaves them you see them one after another shooting into the most perfectly symmetrical crystals. Wonderfully interesting is it to watch how the little particles rush to each other, and unite to form those exactly



FIG. 13.—What is seen when a drop of concentrated sea-water is evaporated under a microscope.

proportioned outlines. First you observe some oblong, pointed forms (gypsum, or sulphate of lime), which probably began to appear before you took the drop out of the concentrated salt-water; then, as the drop rapidly evaporates, a crowd of little squares marshal themselves all round the edges, the last of the water passes off, and every particle and crystal becomes fixed and motionless. You have driven away the water, and what you see on the glass are the **salts** which give the sea-water its characteristic taste and weight. If you live far away from the sea you may see a part of this experiment by taking a drop of water into which as much common salt

has been put as it will dissolve, and placing it under the microscope. The square crystals will be seen growing round each other with great vigour and in perfect symmetry.

5. While the different forms of crystals indicate the different saline materials present in the sea-water, each substance which crystallises keeping its own form of crystals, the relative abundance of each form shows in a general way the proportions in which the substances occur in the water. We see, for example, that by much the most abundant are the square or cubical crystals. These belong to what we call common salt, or chloride of sodium. Moreover, they come last in the order of appearance of the crystals, whence it is clear that they can remain longer dissolved in the water than the other substances. They are therefore said to be more soluble. Now while the dry film left by the drop is still below the microscope, breathe gently on it, and watch what takes place. The moisture of your breath dissolves the crystals one by one, until, if you have supplied moisture enough, they wholly disappear in the little drop of water which now replaces them. The cubical crystals of salt vanish first. Should there be much vapour in the air, you may not be able to keep the crystals from attracting it and dissolving, unless you heat the glass before a fire.

6. When sea-water is analysed it is found to contain about three-and-a-half parts by weight of these various salts in every hundred parts of water. As our evaporated drop indicated, common salt forms by much the largest proportion, at least three-fourths, of the whole amount of salts. The other saline substances are called chlorides of magnesium and potassium, sulphates of lime (gypsum) and magnesia (Epsom salts), and some others in still smaller quantity. The proportions do not greatly differ throughout the whole extent of the ocean.

7. The next question that arises is, Whence does the sea get its salt? This is not quite so easily answered as the previous one. We cannot tell what the earliest condition of the sea may have been. Probably its waters have always been salt ever since they condensed out of the original atmosphere of gas and vapour, and carried down the saline vapours which were no doubt at first diffused abundantly through that atmosphere. But even if the sea had been quite fresh at the beginning, it would have been perceptibly salt now. In Lesson XXVII. you will find an account of the way in which the fresh-waters which flow off the land, all, without exception, carry saline substances to the sea. As this process goes on without ceasing, and in every quarter of the globe where land exists, no one can fail to understand how vast must be the yearly tribute of this kind received by the sea from the land. Probably, if the methods of analysis were delicate enough, it would be possible to detect in sea-water traces of every substance which is dissolved in the waters of the springs, rivers, and lakes of the land. Some ingredients, for example, which are present in too minute proportion to be detected by the chemical analysis of a small quantity of sea-water, are absorbed by marine plants and animals, and found in their ashes or skeletons. The copper bottom of a ship which has been sailing for months gathers out some of the dissolved arsenic or other metals which are diffused through the waters of the sea, and these substances are likewise found in the crust which forms inside the boilers of ocean-steamers where salt-water is used.

8. Among the ingredients present in sea-water in such minute proportions as almost to escape notice, there are two of such importance that we must note their names in passing—**carbonate of lime** and **silica**. The former of these is the material of which all the familiar forms of shell, coral, sea-urchin, and star-fish, are mainly built up.

When embodied by these sea-creatures as part of their skeletons, it becomes the white chalk-like substance which we recognise as the material composing the common shells of the sea-shore. Silica is secreted by lowly forms of plant and animal life in the sea. The hard compact stone called flint is a common form of ancient sea-grown silica.

9. Sea-water likewise contains air. That this must be the case is evident from what takes place in windy weather at sea. The surface of the water is tossed up into waves, and these are blown into spray. The sea becomes a sheet of white foam, which is a mixture of air and water. Some of the air remains dissolved in the water after the waves have disappeared, and is slowly diffused even to the depths of the ocean. The quantity of air present in sea-water varies from time to time; sometimes it has been found to amount to no more than one-hundredth of the volume of water, at other times to as much as one-twentieth.

10. Does this seem a small proportion? At all events we cannot but look upon it with interest, for it is this included air which enables the animals of the ocean to breathe. They inhale it, and exhale carbonic acid. Hence that gas, besides the proportion of it in the dissolved air, is likewise present in sea-water. In the higher parts, which come under the influence of the waves, the carbonic acid is partly removed, and fresh supplies of air are furnished in its place.

11. There is still one other substance in sea-water deserving of notice here. It is called **organic matter**, or **protoplasm**, the material which serves the simpler forms of marine creatures as food. It seems to be derived from the substance of plants and animals both in the sea and carried by rivers from the land.

12. It appears, then, that the sea is salt, partly probably from the vapours of the original atmosphere from which

it condensed, partly because many different salts are continually carried into it by brooks and rivers ; that it contains in solution more or less of all terrestrial substances which water will keep dissolved ; that it is aërated by the waves, the air being diffused through its mass, and serving for the respiration of marine animals ; and that from the dead bodies of animals and the remains of plants, both on land and sea, the fundamental substance or protoplasm of which plant and animal forms are built is diffused through the whole mass of the ocean.

LESSON XIV.—THE DEPTHS OF THE SEA.

1. From what has been said in Lesson XII. it is plain that the bottom of the sea must be uneven and undulating, like many parts of the surface of the land ; stretching out in some parts into vast plains, rising here and there into broad and long ridges, shooting up into vast precipitous slopes, like that of the Atlantic floor to the west of Britain, sinking into deep hollows, which descend far below its general level, as lake-bottoms do below the general level of the land, and traversed by mountain ranges, of which the oceanic islands are only the unsubmerged peaks. But what is the nature of the surface of the sea-floor? Is it solid rock, or soft mud, or barren sand,—or are the submarine plains covered with an oceanic vegetation, as the plains of the land are with grass and herbage?

2. Standing by the margin of the sea we observe that the water breaks upon sand, gravel, mud, or strewn fragments of stone or shells, and that these materials pass down beneath it. If the shore is rocky, pools of the salt-water may be noticed, from which some idea may be formed of the nature of the bottom of at least the shallower parts of the

sea. Each of these pools forms as it were a miniature sea. Its sides are hung with tufts of delicate sea-weeds and bright with clusters of sea-anemones, while many a limpet and periwinkle stands fixed to the stone, or creeps cautiously over its surface. The bottom of the water abounds in shady groves of sea-weed, through which many tiny forms of marine creatures dart or crawl. As we look into one pool after another we find them all to be more or less full of plant and animal life.

3. Turning from these shore pools to the edge of the sea itself when the tide is low, we mark that the ledges of rock support a thick growth of coarse dark green or brown tangles and sea-wrack, among which, if the water is still enough, tiny crabs, sea-urchins, jelly-fish, and other bright-coloured marine animals may be seen. If the water is examined from a boat, this forest-belt of large dark seaweed is found not to extend to a greater depth than a few fathoms. Beyond it the bottom, whether rocky, sandy, or muddy, can be seen through the clear water, or may be examined by means of the dredge. Delicate scarlet sea-weeds with corallines and deeper-water shells inhabit these tracts. The sea-weed belt which fringes the land has an average breadth of about a mile. Beyond it, as we gradually get into deeper water, the common plants and animals of the shore are found one by one to disappear, and other kinds to take their place. The dredge may be dragged along some parts of the sea-floor and bring up only sand or mud, while at a short distance off it may come up full of many and varied forms of marine life, thus showing that there must be bare tracts of sand, mud, or stone on the sea-floor, and other patches where plants and animals are crowded together.

4. Down to a depth of about fifty fathoms the waters of the ocean round the margin of the land abound in plants and animals. This upper and marginal belt

supports indeed by far the most abundant and varied marine population. Until recently, naturalists supposed, that as they found fewer and fewer traces of life as they dredged in deeper and deeper water, the great abysses of the sea must be barren of life. They even fixed the limit at three hundred fathoms, below which, as light ceases to penetrate further, they thought the waters of the ocean must be lifeless.

5. But within the last few years, researches, carried on across the Atlantic and other oceans, have shown that no such limit really exists, that, on the contrary, an abundant and varied assemblage of marine animals covers many parts of the sea-floor at depths of more than two miles, where total darkness must prevail, and where the temperature is very low. The species of animals obtained from these depths, however, differ from those of the shallower tracts. Plant-life is much more restricted in its range. Sea-weeds abound from high-water mark down to about fifteen fathoms; they become rare beyond a depth of fifty fathoms, and are not found at all in water deeper than two hundred fathoms, although some other kinds of marine plants occur at still greater depths.

6. In such comparatively shallow water as that which surrounds the British Islands, where the action of sea-currents extends to the bottom, banks of sand, with patches of broken shells and gravel, and wide flats of mud, cover much of the sea-floor. Some of these sand-banks, where oysters, clams, and other shells live, furnish good feeding-ground for cod, haddock, turbot, and other familiar fishes, and are therefore well known to fishermen. They seem to retain their place and form for many years. But in the shallower parts of these seas they are apt to be altered by storms; some of them, such as the Dogger Bank, come so near the surface that vessels are often driven upon them and wrecked.

7. The mud, sand, and gravel which are strewn over these tracts of the sea-bottom all consist of water-worn particles of rock. If we could trace back the history of these deposits we should find that each grain of sand and each pebble of gravel once formed part of the solid rocks of the land. It will be pointed out in Lesson XVIII. that these materials are not produced on the sea-bottom, but are brought thither from the land. The deeper water of the sea has too little motion to be able to grind down rocks into sand and mud. It is along the margin of the land, where waves can act upon the rocks, that this kind of waste goes on. Every stream, too, is busy carrying and pushing mud, sand, and gravel to the sea. Thus, whether the materials have come from the surf-beaten shore, or from the hills and valleys of the land, they are strewn over the sea-bottom, and mingle with the plants and animals which inhabit these depths. The sea is indeed the vast receptacle into which all the materials derived from the crumbling of the surface of the land must ultimately find their way. Wherever there is perceptible movement of the bottom water, sand, and even fine gravel, may be pushed along. But, throughout most of the great sea-basins, the water which lies upon the bottom, though not quite motionless, is too sluggish in its flow to disturb the fine sediment which must be the only kind of deposit that can be slowly accumulated there.

8. Much light has recently been thrown upon the condition of the deep-sea bottom by the soundings of the *Challenger* expedition. In comparatively shallow water near land, whether islands or continents, variously coloured deposits of mud and sand occur, evidently in great part formed from the finer sediment which has been worn away from the surface of the land and carried out to sea. But away from land, in the wide and deep depressions of the ocean-basins, other characteristic deposits are laid down.

The most universal of those are fine red and grey clays, in which minute grains of sand, and fragments of the volcanic rock called *pumice*, may be recognised. Remains of minute organisms (*Globigerina* and *Radiolaria*) occur in them, sometimes abundantly. Most of the bottom of the Pacific and Atlantic basins, at depths of more than two thousand fathoms, as far as yet explored, is covered with these clays. Other wide-spread oceanic formations, called ooze, are derived from the accumulation of the remains of minute plants or animals. In the Southern Ocean, towards the icy Antarctic land, a pale, straw-coloured deposit has been dredged up from depths of nearly two thousand fathoms. It consists chiefly of the remains of minute siliceous plants called *diatoms*. Another kind of siliceous deposit, formed from the minute animal forms called *Radiolaria*, covers a considerable space in the intertropical regions of the Pacific.

9. But the best known of these deep-sea formations is the ooze which was first brought to light when the earliest soundings were made for the telegraphic cable between Britain and America. The floor of the Atlantic for thousands of square miles is covered with a fine, soft, chalky deposit. A sounding-lead dropped upon the bottom sinks deeply into this ooze. When the dredge goes down, it plunges into this light, fleecy mud, and comes up full of it. The upper layer is a yellowish, creamy, somewhat sticky substance, full of entire and broken shells of such minute animal organisms (*Globigerina*) as are shown in Fig. 14, with fragments of sponges and other forms. Underneath this surface layer the ooze shows the shells in a more crumbling state, and no doubt in the deeper parts of the deposit, which the dredge cannot reach, the substance may consist only of a kind of fine chalk, formed almost entirely of the mouldered remains of these lowly forms of life. Living sponges,

star-fishes, crinoids, and various shells have been brought up by the dredge with the ooze. It would seem that while this widely-extended deposit may contain the remains of such forms of life as inhabit deep water, it is derived

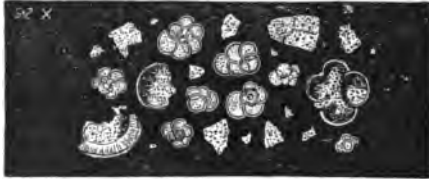


FIG. 14.—Some of the Ooze of the Atlantic floor, magnified.

mainly from the *Globigerina*, which lives in immense numbers in the surface waters, and which, when dead, sinks through the depths to the bottom.

LESSON XV.—THE TEMPERATURE OF THE SEA.

1. Since the temperature of the air and of the land varies in different latitudes, we may expect to find that of the sea to show corresponding variation. There are indeed climates in the sea as well as on the land. The temperature of the surface water of the sea varies with the seasons, and with the direction of the great ocean currents. Thus, on the west side of Britain this superficial sea-temperature is about 49° Fahr., and over a great part of the North Atlantic it ranges between 44° and 54° . Further south, in the equatorial part of that ocean, it rises to from 80° to 83.5° . In the North Pacific Ocean the mean surface temperature is about 70° , and in the South Pacific about 67.7° . In Polar seas, away from the influence of warm currents, the water does not rise much above the temperature of melting ice. On the

other hand, in land-locked parts of the ocean within the tropics, where the water is confined, and exposed to great solar heat, the temperature sometimes rises to that of a warm bath. Thus, in the Red Sea it has been noted at 94° , and in different parts of the Indian Ocean at from 88° to 91° .

2. The importance of the temperature of the sea is seen in the influence which it exerts upon the climate of the maritime tracts of land. The temperature of these regions depends greatly upon that of the winds which come from the sea. The air resting upon the ocean surface acquires the temperature of that surface and carries it over the land. Save for this effect upon the air the sea would have comparatively little influence upon the climates of the land. But by warming or cooling the air over tracts thousands of square miles in extent it becomes the great distributor and regulator of temperature. This has been partly explained in Lesson XI., and will be further treated of in Lesson XVIII.

3. It is evident, however, that if we take the readings of the thermometer immersed merely in the surface water, we may obtain no indication whatever of the real temperature of the great mass of the sea. In the air, for instance, observations with the thermometer at the sea-level would still leave us in the dark as to the temperature at different heights above that level. But by taking observations at intervals up to the mountain-tops, the gradual lowering of the temperature referred to in Lesson IX., Article 20, has been determined. In like manner, by means of a specially arranged kind of thermometer, the temperature of the sea at any depth can be ascertained. The instrument is protected in such a way as to be unaffected by the pressure of the water. It is let down with a cord, allowed to remain long enough at each depth to acquire the temperature of the surrounding water, and then pulled up.

4. In this way the temperature of the Atlantic and Pacific Oceans has been determined at intervals from the surface down to the deepest depressions of the bottom. Temperature-soundings have likewise been made in the Indian Ocean. As the result of such observations all over the world it has been discovered that in general the surface-water is the warmest part of the sea, and that the water gets colder towards the bottom. In the open sea indeed, communicating freely with the Polar regions, as in the North Atlantic Ocean, the deep bottom-water has everywhere been found to have very nearly the same temperature as that of the cold Polar seas. Sometimes the thermometer has come up with readings as low as 29.5° ; that is, two and a half degrees colder than the freezing point of fresh water.

5. Soundings taken by the ship *Challenger*, which was sent out by the British Government in the year 1872 to ascertain the depths, temperature, and other characters of the great sea-basins, showed everywhere a progressive lowering of the temperature from the surface downwards, with variations in the rate of decrease. In the North Atlantic, water having a temperature above 40° forms the upper layer of the ocean to a depth of from 750 to 1000 fathoms. Under that warmer sheet the temperature slowly sinks, until in the deep abysses, more than 3,000 fathoms down, it reaches 34.4° . Hence the main mass of this ocean consists of chill water. In the South Atlantic the warm upper layer is only about 300 fathoms thick. Below it, the temperature of the water falls from 40° to 32.9° , so that the body of chill water is colder and comes nearer to the surface than in the North Atlantic. In the equatorial part of the ocean, while the surface temperature is 76° — 80° , the cold water under 40° is found at a depth of only 300 fathoms. The bottom layer of water is at less than 35° for a thickness of 600 fathoms,

while some of its deeper cavities have the icy temperature of 32.4° .

6. In the North Pacific Ocean a general bottom temperature below 35° Fahr. has been observed. Sometimes the body of water of this low temperature exceeds 2,000 fathoms in depth. In the South Pacific Ocean, beyond Cape Otway, Australia, the bottom temperature sinks to 32.5° , while further south (Lat. $53^{\circ} 55' S.$), the body of icy water below 35° comes within less than 100 fathoms from the surface. Below a depth of 1,000 fathoms the water is as cold as ice (32°) down to the bottom, where the thermometer marked 31° at a depth of 1,950 fathoms.

7. This low temperature of all but the upper parts of the sea shows that as the deep water could not be cold unless it came from cold latitudes, there must be a continual transference of water from the cold Polar tracts towards the equator. The heavy cold water sinks down, and creeps on below the upper warmer layers. These, again, move towards the poles to supply its place. When the cold water reaches the equatorial regions, it rises towards the surface, where, being gradually warmed, it moves away again as a surface layer towards the poles. Hence these soundings not only reveal the various temperatures of the sea, but bring before us evidence of vast slow movements, by which the waters of the sea are constantly mingled and kept from stagnation.

8. But the temperature soundings have sometimes brought out in another way proof of the slow circulation of the water of the sea. In the North Atlantic, between Scotland and the Faroe Islands, below a uniform surface temperature of 50° — 53° Fahr., two parallel belts of water have been observed lying side by side and having nearly the same depth. In one of these the bottom temperature was found to be 42.7° , in the other 29.6° . Two very different sea-climates were thus ascertained to lie close

together, one with the mild character of the latitude, the other with the chilly temperature of the icy Arctic sea. The cold belt can be nothing else than a slowly moving stream from the Polar Ocean. In the sand brought up from the bottom of that belt, particles of volcanic rocks were observed, which may have come from the volcanic island of Jan Mayen, lying far to the north-east.

LESSON XVI.—THE ICE OF THE SEA.

1. When sailing across the North Atlantic between ports in Europe and in the United States, or when traversing the South Pacific to round the southern promontory of the American or African continent, vessels often encounter floating masses of solid ice, termed **Icebergs**. These are of all sizes up to huge mountain-like islands rising several hundred feet out of the water. Ice being lighter than water, an iceberg projects above the sea-level, yet has about eight times more of its bulk under water than above. Hence, when the mean height of one of these floating masses is 300 feet above the surface of the sea, its bottom must be some 2,400 feet below. The bergs vary infinitely in shape as well as in size. Sometimes, as in the Antarctic Ocean, they take the form of vast square blocks with vertical walls; sometimes they bristle into peaks and pinnacles, with deep cliffs and gullies between. At a distance they look like snow-covered islands. Seen closer, they gleam with all the intensity of colour—white, green, and blue—so characteristic of the ice of glaciers upon the land. For the most part, nothing but ice in different forms is seen upon their surface. Now and then a block of stone or some dark earthy rubbish may be

noticed. As they drift onwards into warmer latitudes they melt away both under water and above. Cascades of water tumble down their thawing slopes and fall into the waves below. It often happens that as the melting goes on under water, the centre of gravity of an iceberg is altered, so that the mass shifts its position. or, becoming top-heavy, turns completely over.



FIG. 15.—Iceberg at Sea.

2. Such are the drifting icebergs which in summer cross the navigation track across the Atlantic between Newfoundland and Britain. They form a serious source of danger to vessels, for they may be encountered at any hour of the voyage. Many a ship has struck against one in the dark and gone to the bottom. They chill the air around them, and thus often give rise to dense fogs. A

sudden fall of temperature is regarded by sailors as an indication of their approach to an iceberg, and a warning to be on the outlook.

3. But icebergs are not formed in the open sea, where they float until they are entirely melted. To reach their birthplace we must travel to the cold regions which surround each pole. In North Greenland, for example, the country is covered with ice, which creeps slowly down the slopes and gathers in the valleys into wide and deep masses, which not only come down to the sea-level, but actually push their way out to sea in vast walls of ice rising 300 or 400 feet above the water, and stretching sometimes for sixty miles along the coast. From time to time portions of these great tongues of ice break off and float out into the open sea. These are icebergs. They consist, therefore, not of frozen sea-water, but of true land-ice. Each berg which may be met with floating and melting far away in mid-ocean had its origin among the snows of the Polar lands. (Lesson XXVIII.)

4. Since so large a proportion of an iceberg is under water, we may expect to find that these floating masses are comparatively little influenced by the winds and waves which touch only their upper parts. The larger bergs move with the ocean-currents or drifts in which so much of their substance is immersed. Hence they may sometimes be seen moving steadily and even rapidly along, right in the face of a strong gale.

5. So long as the icebergs sail over deep water, they move freely about as the currents or winds may drive them. But when they get into shallow enough water to allow their bottoms to grate along the sea-floor, they tear up the mud or sand there, until at last they are stranded. The coast of Labrador is often fringed with such stranded bergs, some so small as to be driven to the beach, others so large as to run aground while still a good way from



FIG. 16.—The Birthplace of the Arctic Icebergs.

the shore. Chill fogs consequently hang over that desolate region all through the summer.

6. Besides the icebergs, two other kinds of ice are found on the sea in high latitudes, both derived from the freezing of the water of the sea itself. As voyagers advance into the narrow seas about the pole, they encounter great sheets of ice, which, at first in irregular, scattered fragments, become larger in size and more continuous in extent, until at last the whole expanse of the sea, as far as the eye can reach, is covered with ice. This is known as **Floe-ice** or **Field-ice**, and is due to the freezing of the surface-water of the sea during the intense frost of Polar winters. When the sea-water freezes, the salt is left behind by the ice, which, except for the little salt vesicles entangled in it, is fresh, and, when thawed, quite drinkable. In early summer, the ice-field, or sheet of ice, which, stretching outwards from the land, may be regarded as continuous for hundreds of miles during winter, begins to break up and float away. If the sea remained motionless it would, by the end of winter, be covered with a sheet of ice about eight feet thick. But owing to the pressure arising partly from the movements of tides and currents in the water, partly from changes of temperature and atmospheric pressure, including winds and storms, the ice-sheet is liable to continual disruption. Cracks are formed abundantly through it, and portions are squeezed up or broken into innumerable huge fragments, which are pushed over each other until the frozen sea becomes like a heap of ruins. And these movements are accompanied with sharp reports, as of cannon, and with loud growling noises, as if some monsters of the deep were engaged below in fierce and angry warfare. Some notion of the appearance of this broken-up ice may be formed from Fig. 17. This represents a scene in the

memorable Arctic voyage of the Austrian vessel *Tegethoff*, which, after almost incredible endurance, was even-



FIG. 17.—Scene among the disrupted ice of the frozen Arctic Sea.

tually abandoned by the courageous crew, who, taking to

boats or sledges drawn by dogs, pushed across the frozen sea and reached Novaya Zemlya.¹

7. In consequence of the breaking up of the vast sheet of ice, openings are formed in which the unfrozen water can be seen, until it is fast bound in ice by the cold. These water-channels or lanes, between the separated fragments, are the passages, by which vessels make their way through the ice-pack. But as the great sheets are carried against each other in the general movement, it sometimes happens that a ship is caught between them, and pushed up on the floe, or crushed so effectually that it goes to the bottom as soon as the fragments of ice separate again.

8. Except where piled in heaps by pressure, which breaks up its surface as well as its outer edges, floe-ice occurs in level sheets, the surface of which rises but little above the level of the sea. It never rivals the height and grandeur of true icebergs, though it covers a much wider space of sea. Nor does it travel so far from the regions of its birth. When the ice-field breaks up in summer the portions next the land may remain there, and of these indeed some parts may continue unmelted for generations. But other parts further from shore, and sometimes many hundreds of square miles in extent, are loosened and carried away by the sea-currents which drift from the pole. Vessels frozen into the ice-field have in this way been carried many hundreds of miles, until disengaged by the disruption and thawing of the ice. Some of the icebergs from Greenland, however, travel much further south, and may be met with now and then even as far south as lat. 37°, that is, about the same parallel as Richmond in Virginia, and Cape St. Vincent in the south of Spain.

¹ For a graphic account of this and other Arctic voyages, see *New Lands Within the Arctic Circle*, by Julius Payer.

9. When the sea freezes along the margin of the land, as it does in a remarkable way in North Greenland, it forms a cake of ice which, rising with the tide, is frozen to the shore. By degrees a shelf of ice, called the **Ice-foot**, rising from twenty to thirty feet above the general level of the floe, and having a width of 120 feet or more, forms along the coast and clings to it all winter. Immense quantities of earth and stones, dislodged from the coast



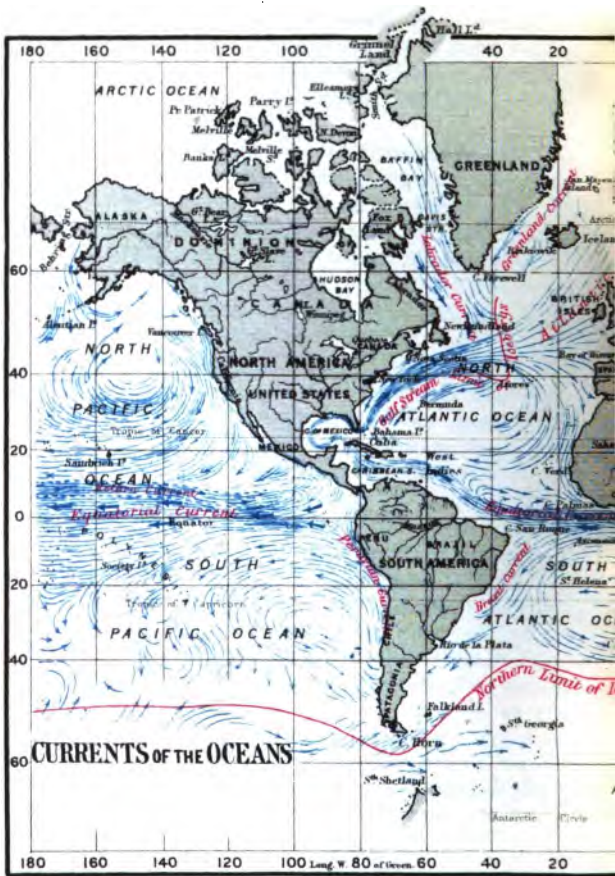
FIG 18.—The Ice-foot of Greenland.

cliffs by the severe frost of the Arctic winter, fall upon the ice-foot, so that its surface becomes in some places a field of rubbish which completely covers and conceals the ice underneath. When the summer storms come, this shore-terrace of ice is broken up, and large pieces of it, laden with the waste of the cliffs, are floated away out to sea,

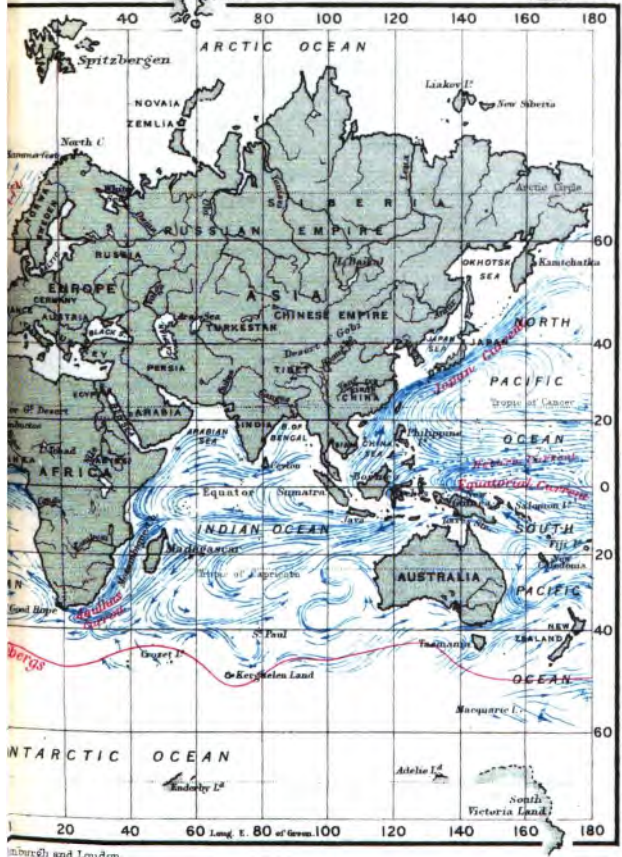
there to join the fleets of bergs and broken sheets of floe-ice. Some portions are driven ashore again, others are caught and frozen fast into the floe-ice of next winter, while others succeed in escaping into the more open sea, where they gradually melt and tumble their load of earth and stones on the sea-bottom.

10. Floe-ice and the ice-foot are formed by the freezing of the surface of the sea. There is yet another way in which some of the ice of the sea takes its origin, viz., by the freezing of the water lying on the sea-bottom. This is known as **Ground-ice**. It is probably formed only in inclosed and shallow seas and inlets, and is of little consequence compared with the thick and wide sheets of floe-ice. It is well known in the Baltic Sea. In still weather, before the surface of that sea is frozen, little thin cakes of ice may be observed floating about, sometimes with portions of sand or scattered pebbles imbedded in them. These are formed on the bottom, from which they break away and rise to the surface. They form a source of some danger, at least to small boats, for they sometimes appear suddenly in such numbers as to cover the sea, the surface of which is then apt to freeze too, so that the boats are in danger of being nipped between the detached ice-rafts, or of being beset and frozen into the united cake of ice. Sometimes large blocks of rock as well as masses of seaweed are borne away from the sea-floor and carried to the surface by the ascending sheets of ground-ice. In the rivers of cold countries, as for instance, in the St. Lawrence, similar ground-ice is formed in winter, sometimes round iron chains or anchors, which, if the inclosing ice is thick enough, may actually be lifted by it.

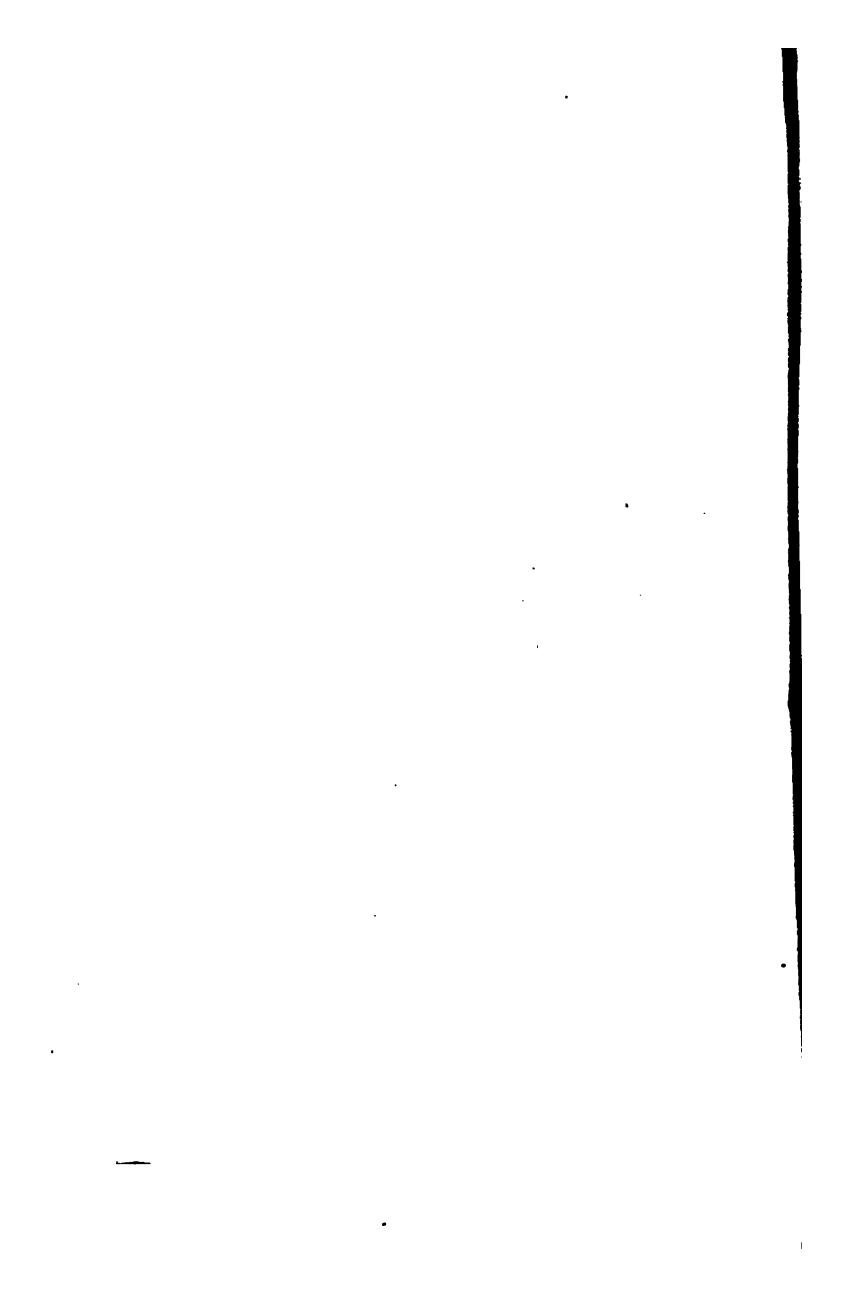




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Land



Edinburgh and London.



LESSON XVII.—THE MOVEMENTS OF THE SEA.

1. The restlessness of its water is one of the features of the sea which first impress the onlooker. Even in a calm day, when not a leaf appears to be stirring on the earth, and the clouds seem perfectly motionless in the sky, the sea may be seen heaving, or curling into faint ripples, or rolling in broad undulations which break into foam when they reach the shore. On the other hand, in stormy weather the unrest of the air is fully equalled by the tumultuous uproar of the sea, when clouds of driving rain and salt sea-spray are swept along by the tempest until it almost seems as if sea and sky were commingled.

2. We do not at first perceive that the movements of the sea are not all by any means so fitful and capricious as they appear to be, but that, on the contrary, they are regulated by easily understood laws, and can be foretold and, if need be, provided against. Nor do we at once realize to what different causes these movements are due.

3. Let us suppose that, placing ourselves in some favourable position for observing the motions of the sea, we carefully register the facts which come under our notice. The margin of any of the great oceans would be suitable for the purpose. Such a coast-line as that of the west of Ireland, for example, forms an excellent field of observation, allowing opportunities of watching some of the most characteristic movements of the water of a great ocean basin.

4. The first point we remark is the obvious close dependence between the motion of the air and that of the water. As a rule, when the air is still the water is smooth, though on some days we should find that even with no wind, a good deal of agitation might be observ-

able in the sea, and large waves might roll heavily to shore. In such cases, however, we may often learn that a gale has been blowing somewhere out at sea, raising the waves, of which the final undulations now reach the land. Every gradation of ripple, wave, and billow may be traced by us on the water, answering doubtless to gentler or stronger movements of the air. In this respect the surface of the sea shows the same sensitiveness as that of any lake or pond on the land, but its movements are more marked, because it is so much larger, and allows the accumulated effects of the continued pressure of the wind to make themselves far more apparent.

5. But we could not long attend to the aspects of the sea without observing traces of other movements independent of the daily or hourly variations in the state of the atmosphere. The Irish coast, which we have taken by way of illustration, affords now and then, like that of other parts of north-western Europe, scattered leaves and fruits which have been washed ashore from the wide ocean, and which differ greatly from the vegetable productions of any European country. These are in reality West Indian plants. They have been drifted across the Atlantic, and hence they indicate the existence of some general drift or current which sets across the breadth of that ocean from west to east or north-east. In like manner bottles containing pieces of writing which have been thrown overboard from ships at sea have been picked up on the same shores, after safely traversing hundreds or even thousands of miles of sea.

6. But by far the most constant and regular movement is one which, if we have never witnessed it before, cannot but fill our minds with wonder. Twice every day we find the edge of the sea to rise and fall or to advance and retreat upon the land. On a line of steep coast-cliff the

surface of the sea may be observed slowly to sink during about six hours, and then for a corresponding interval slowly to rise again. On a flat or gently shelving shore this vertical movement manifests itself in another way. The fall of the water lays bare wide stretches of rocks, sand, or mud. During the retreat of the sea we notice that each ripple or wave does not come quite so far as those which preceded it. Step by step the waves seem to be drawn back, until in some large bays many square miles of flat ground are laid bare, across which we might walk. During the advance of the water, on the other hand, each succeeding wave comes a little further up the shore than its predecessor, until all the uncovered flats are once more under the sea, and boats can sail over places which two or three hours before we may have crossed on foot. This rise and fall of the sea is known as the Tides.

7. A little observation enables us to make quite sure that these regular movements have no connection with those of the air. Even when a strong gale is blowing from the land, and driving the surface of the sea into foaming waves, it does not arrest the advance or retreat of the tide. The waves continue to creep up or down the shore even in the face of the wind.

8. Again, we cannot but remark the wonderful regularity of the movements. Each advance of the water occupies rather more than six hours, and each retreat requires the same interval, so that the time of high or low water may be foretold with accuracy. From an early period of human history, and long before the real cause of the tides was understood, it was observed that there is a constant relation between them and the state of the moon. In the course of the observations which we have supposed ourselves to make, we find that the time of high water corresponds with the passing of the moon across the

meridian, and that the highest and lowest tides occur at the time of new and full moon.

9. Putting now the sum of these observations together, we have ascertained the existence of three distinct kinds of movement in the water of the sea. First, wind-waves ; second, surface-drifts or currents ; and third, tides. From what was said in Lesson XV. it is clear that a fourth form of motion must be added to these, viz. that general creep or movement of the cold polar water along the bed of the sea towards the equator, and the flow of the upper waters towards the poles. We must consider each of these movements a little further.

10. (i.) **Wind-waves.** These, like the ripples we send along the surface of a trough of water by blowing briskly at one end, are due to the pressure of the wind. Viewed from the top of a high cliff, or from the mast of a ship at sea, they are often seen to succeed each other in long parallel lines, travelling regularly and rapidly across the surface of the sea. Though the waves move forward, the water itself in the open deep sea shares but slowly in this movement. It rises and falls with a slight oscillation as each wave passes, though, if the wind continues to blow for a time, the surface water is gradually pushed onward. A field of corn or tall grass may often be observed to be thrown into waves as a smart breeze sweeps over it, yet each stalk and blade retains its place, and merely bends up and down and to and fro with a kind of motion like that of the particles of water in a wave.

11. When the sea has been thrown into violent agitation by a storm, the waves do not at once subside when the storm is over, nor do they remain confined to that region only through which the storm has passed. So sensitive is the great body of oceanic water that the waves are propagated far beyond the area of the storm in vast undulations, or *ground-swell*, as this heaving is called.

In the deep water of an open ocean the only trace of this movement is seen in the broad undulations which, like a great pulse, keep the surface regularly rising and falling. A ship sailing over this surface alternately mounts on the top of a swell and sinks into the trough between two of them. But as the sea gets shallower towards land, the up-and-down movement passes into a true onward rush of the water. The upper part of the undulation travelling faster than the lower parts, which are impeded by the friction of the bottom, begins to assume the aspect of a wave, curling over as it advances like a huge wall of green water, to burst into foam against the land.

12. The motions of waves and ground-swell must be insensible in the deeper parts of the sea. In reference to the great body of the sea, they are merely surface agitations, probably seldom extending sensibly more than a few hundred feet downwards. So that when a hurricance is raising the surface of an ocean into the most violent commotion, we must think of the deep abysses below as dark, silent, and calm. And this is a matter of some importance when we consider the various offices of the sea, as will be done in the following Lesson.

13. Since the longer the space of deep sea over which the wind has to blow, the larger are the waves, we may expect to meet with the grandest waves in the great oceans. In the accounts of storms at sea it is common to meet with such expressions as that "the waves were mountains high." The size of waves is apt, however, to be much exaggerated. Thus, a series of measurements, taken during a voyage across the North Atlantic, gave forty-three feet as the extreme height of the waves in stormy weather. In the narrow British seas, which are in a great measure barred off from the Atlantic, the largest waves are less than half the maximum size which they attain in that ocean.

14. No one can watch one of these huge waves as it nears the land without being impressed by the great force with which it beats upon the coast. The scene represented in the frontispiece of this volume may give some idea of the grandeur of a stormy sea. Each long undulation which rolls in from the outer sea mounts higher as it approaches the shore. At every step it becomes a more and more marked ridge of water, which begins to curl into a green-crested wave with a steep concave front towards the land. By degrees the upper advancing crest topples over, and the immense mass of water plunges with all its weight upon the rocks or sand of the coast-line. When the broken water has spent itself in rushing up the beach, it runs rapidly down again under the next advancing wave. In its recoil it drags back the sand and the gravel with a loud rattle or hoarse roar, sometimes audible for many miles, as the stones of the shingle are ground over each other.

15. The force with which such *breakers*, as they are termed, fall upon the land has been measured and calculated. An undulation, or "roller," of the ground-swell twenty feet high has been computed to fall with a pressure of about a ton on every square foot of surface exposed to its reach. In summer the average force of the Atlantic breakers on the west coast of Britain amounts to about 611 lbs. on the square foot; in winter the force is more than three times as great. On some occasions a pressure of as much as three tons and a half has been recorded. Some of the effects of the action of waves upon the land will be referred to in the next Lesson.

16. (ii.) **Currents.** Since the sea responds so sensitively to every movement of the air, we may expect to find that as there is a regular system of circulation in the atmosphere, so there must be a corresponding system in the ocean. That this is really true has now been **made**

clear by observations in all parts of the world. The sea has been found to be in continual circulation. Such evidence as that given in Article 5 makes the reality and extent of the surface-movement readily apparent. Those superficial parts of the sea which have a regular, continuous, onward motion, are termed "currents." They are called superficial because seldom more than 500 feet deep, which is but a small fraction of the depth of the total liquid mass over which they move.

17. The current-system of the sea presents in its leading features great simplicity. It arises from the action of the great atmospheric currents, which, steadily blowing on the surface of the sea, drive its waters before them. So that if we clearly follow the course of the circulation of the air, we need have little difficulty in understanding that of the sea. The general features of the circulation of the oceans are shown in Plate VIII.

18. The trade winds set the surface-water of the sea in motion, so that it acquires a gradual tendency towards the equator. On the north side it comes from the north-east, on the south side from the south-east. Uniting along the equatorial belt, it necessarily assumes a westerly direction, and crosses both the Atlantic and Pacific oceans as the *Equatorial Current*. In the former ocean it appears not to exceed about 300 feet in depth, and to move with a surface rate of not more than eighteen miles a day.

19. Were there no land to interfere with its flow, this Equatorial Current would pass round the globe as a continuous stream of warm water. But owing to the position of the continents across its path it is broken up and made to diverge in different directions, each of the independent branches receiving a distinct name. Thus in the Atlantic Ocean, setting out from the African coast, it crosses to the American side, where, striking against the projection of Cape St. Roque, it splits into two. The smaller branch

turns southward as the *Brasil Current*, skirts the coast as far as the mouth of the La Plata, whence it bends eastward, once more crosses the Atlantic towards the Cape of Good Hope, and turns northward along the west coast of Africa, until it is drawn again into the great equatorial current. The larger branch sweeps round the northern coast-line of South America in the Caribbean Sea and the Gulf of Mexico, whence it issues through the Florida Strait as the warm and rapid ocean-river well known by the name of the *Gulf Stream*. After emerging from the narrows, with a maximum surface temperature of 80° Fahr. and a rate of 70 to 120 miles a day, this current runs parallel with the coast of the United States, but separated from it by a cold current, which descends from the Arctic seas. It gradually lessens in speed, and begins to spread out and to turn north-eastward. One portion, getting thinner and cooler, but yet retaining a higher temperature than that which is proper to the latitudes, and joining the general surface drift of the Atlantic water towards the north-east, extends to the coasts of Britain and Norway, and is said even to be traceable beyond Spitzbergen; the other turns southward between the Azores and the coast of Spain, bends round the north-west shores of Africa, and then getting once more within the influence of the trade winds, is again sent across the Atlantic.

20. Within the wide circuit embraced by the second or south-eastern branch of the Gulf Stream lies a vast area of comparatively still-water, over the surface of much of which dense masses of sea-weed grow. This tract is known by the name of the *Sargasso Sea*.

21. In the Pacific Ocean the equatorial current has more room to develop itself. Starting from the western coast of the American continent, it streams westward across the whole breadth of that vast ocean, encountering

on its way only the obstruction of scattered islands and submarine banks. It reaches the eastern margin of Asia where that continent is flanked by the islands of the Malay Archipelago. Hemmed in there, it divides into two main branches, one of which turns northward as the warm, rapid, and river-like *Japan Current*, and sweeps along the east coast of Asia as the Gulf Stream does along that of North America. The other and larger branch forces its way into the Indian Ocean, and joins the westerly equatorial current there.

22. These movements of the equatorial waters could not be carried on without drawing the water of the northern and southern hemispheres into the general circulation. Thus in the Pacific Ocean, which is open to the south, there is a general northerly set of the cold southern water. A broad current or drift pours into the hollow of the western coast of South America and passing northwards joins the equatorial movement. Part of it, however, strikes the promontory of Cape Horn, round which it sweeps as a well-marked current. Since the Pacific basin is almost closed at its northern margin, no great current of warm water can pass northwards through Behring's Strait, and on the other hand no considerable body of cold water can escape from the Arctic Sea, although a cold current does issue from the Strait and pass down the west side of North America. In the Atlantic basin, however, where the opening into the polar region is much wider, and where the prolongation of the Gulf Stream bears its warmth far within the Arctic circle, there is an opportunity for the cold water of the Northern Ocean to make its way southward in return currents. One of these flows down the east side of Greenland, another descends from Davis Strait and pursues its course southward along the American coast, interposing between the land and the warm Gulf Stream, and partly passing actually under that current.

Icebergs have been carried across the Gulf Stream as far as lat. 36° S., which would seem to prove that the cold Arctic water continues to flow steadily with a sensible motion for a long way southward. The recent soundings of the *Challenger* show that the Gulf Stream between New York and Bermuda, with a temperature of 70° Fahr. and upwards, does not exceed about seventy fathoms in depth, and that below a depth of 600 fathoms the temperature of the water is always lower than 40° Fahr. Between the Gulf Stream and New York this cold water rises to within 300 fathoms from the surface, while further north, along the shores of Nova Scotia it comes quite to the surface, forming there the surface Arctic current out of Davis Strait.

23. But besides these surface-currents, which are always more or less readily perceptible, the waters of the ocean basins are affected likewise by a slower motion, which extends to the deepest abysses and from the equator to the poles. The existence of this general "creep," or slow diffusion, has been made clear by the temperature soundings described in Lesson XV. The very remarkable fact has been proved that even under the equator the main mass of the sea is cold, except a comparatively shallow layer at the surface warmed by the sun, and that at the bottom the temperature is as low as in the Antarctic and Arctic Seas. Were there no transference of cold polar water towards the equator, the temperature there ought to be comparatively high. But the presence of cold water even within 300 fathoms of the surface shows that there must be a deep and general movement from the polar regions to the equator, under the surface-currents, which, especially in the North Atlantic, carry the warm equatorial water into the cold polar tracts. The cause of this movement of the cold water is not yet well understood.

24. (iii.) **Tides.** One of the most remarkable and regular of the motions of the sea was alluded to in Art. 6, as that alternate rise and fall of the water called the tides, the coincidence of which with the position of the moon could not but be noticed from the earliest times. To understand this motion, we must bear in mind that while all the members of the solar system mutually attract each other, there are two which specially influence our earth—the sun by reason of its vast size, and the moon on account of its nearness. Each of these two luminaries exercises a strong attractive force upon our planet, the tendency of which is to pull out the side of the earth which is opposite to it. The solid part of the globe resists the strain, but the liquid ocean, unable to do so, is drawn outwards so as to be heaped up on that side where the attraction is exerted.

25. On the other side, too, where the distance from the attracting body is greatest, and the force of attraction therefore least, the water is not drawn away but remains behind, forming another vast swelling, the summit of which is exactly opposite to that on the near side of the earth. Were there no land to interfere with these motions, the distortion produced by this cause would give the form of an ellipsoid to the liquid mass covering our planet, its longer axis pointing to the moon. Were the earth and the other members of the solar system at rest, the two swellings of the ocean surface would remain stationary. But owing to rotation, these protuberant masses of water are forced to travel rapidly round the globe.

26. Let us fix our attention on that side of the earth which happens at the time to be facing the moon. Owing to the earth's rotation combined with the moon's own motion, our satellite appears to be revolving round the earth in about twenty-five hours. The outward bulging of the ocean-surface, which is caused by the moon's attraction,

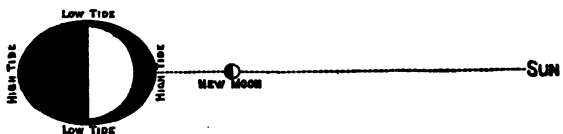
must therefore follow the moon, and run completely round the earth in about twenty-five hours, and the swelling on the opposite side of the earth must take the same course. Each of these two uprisings is felt, at every place which it passes, as a broad undulation or wave which appears once a day, or more exactly in every 24 hours 50 minutes, that being the length of the interval between each appearance of the moon on the meridian. Thus in that interval there are two times of *high-water* or *flood*, and two times of *low-water* or *ebb*.

27. It is not, however, the moon alone by which the water of the sea is affected in this way. The sun, too, draws the ocean-surface outward ; but this influence, which has about one-third of the force of the moon's, is combined with the latter in the production of the broad tidal undulation. When the sun and moon are in the same line, which happens at new and full moon, their combined attraction must evidently most powerfully draw out the surface of the sea, while at the quarters of the moon their respective influences partially neutralize each other.

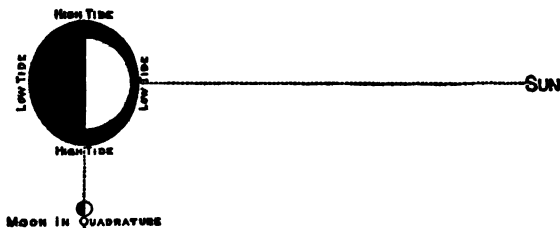
The accompanying figure (Fig. 19) shows how this effect is accomplished. The highest rise and, of course, the lowest fall of the tidal undulation must occur at new and full moon. These are called *spring tides*. The least rise and fall take place at the time of the moon's first and last quarters. These are known as *neap-tides*.

28. Since the movements of the tidal undulation depend upon those of the earth, sun, and moon, they can be measured and predicted a long while beforehand. Observations have likewise been made in all parts of the world regarding the times of high-water and the height of the tide, so that the length of time taken by the tide-wave to travel from point to point has been carefully determined. The rate of motion depends upon the depth of water and

the absence or presence of land, being most rapid where the water is deepest and has no obstructions in its course. In the equatorial parts of the Atlantic Ocean it exceeds five hundred geographical miles in the hour. It is estimated to take fourteen or fifteen hours to come from the



SPRING-TIDES



NEAP-TIDES

FIG. 19.—Diagram illustrating the origin of the tides.

south of Africa to the south-west of Europe, where it gets into comparatively shallow water and requires to struggle through many narrow passages. Striking the south-west of Ireland, it branches out, one part turning north along the west coast of Ireland and Scotland, then bending south along the east coast, while the other part turns to the east up the English Channel. Now the former of these branches takes nineteen hours to circle round the British Islands as far as the mouth of the Thames, where it meets

and mingles with the wave which started twelve hours later, and worked its way in about seven hours through the English Channel.

29. Though the tidal undulation is called a wave, and travels with great speed across the deep-ocean basins, the water itself does not partake of this onward motion. Each particle of water merely oscillates as the rapid pulsation passes, like the stalks of corn in a field, as already cited (Art. 10). But when the undulation enters a narrow and shallow sea, it becomes jammed between the converging coasts and experiences increasing friction against the bottom. Its rate of motion consequently slackens, but in proportion as this takes place the undulation gathers height and force, and becomes a true current.

30. In the deep ocean-basins the passing of the broad tidal undulation probably produces no sensible effect upon the bottom. The surface of the water merely rises gently for about six hours, and then gently falls again, the total height of the undulation from low-water to high-water not being more than a few inches. Thus, in the middle of the Pacific Ocean the rise is sometimes less than a foot; in the Atlantic, round St. Helena, about three feet. But where the undulation meets with the resistance of converging masses of land and a shallowing bottom, it is heaped up, sometimes, as in the Bay of Fundy, to a height of seventy feet, and rushes along as a great wave or as a surging and foaming ocean-river. Thus in the Bristol Channel, which opens towards the west, the tide rolls up a rapidly-narrowing channel, and during spring-tides attains a height of forty feet in the estuary of the Severn. The advance of the tide is shown by a wave nine feet high, which rushes in front, succeeded by smaller ones, until the full tide has filled the estuary. This wave, or *bore* as it is called, forms a marked feature of the flowing tide in many bays and estuaries which open out towards the quarter whence

the tidal undulation comes. Thus on the west coast of Europe it occurs in the Elbe, Weser, Seine, Dordogne, and Garonne. At the mouth of the Seine the tidal wave enters with a speed of fifteen to twenty feet per second and a height of six-and-a-half to ten feet. The bore on the Hooghly River rushes up with such force as to do great damage to shipping unprepared for its approach.

31. When the converging shores are not closed at the end, as in the case of an estuary, but open out again into a wider sea, the tidal undulation takes the form of a rapid, river-like current, or *race*. Heaped up on the one side, it attains a higher level than on the other side of the narrow passage, though which therefore it rushes with great force and speed. The British Islands, lying, as they do, on the margin of the Atlantic Ocean, present many examples, of which the strait called the Pentland Firth, between the Orkney Islands and the north of Scotland, may be taken as one of the best. Standing at high or low water on one of the headlands overlooking that narrow passage, you look down upon a strip of blue sea about six or eight miles broad opening westward into the wide Atlantic and eastward into the North Sea. As soon as the tide begins to flow or to ebb you observe this smooth belt of water to become more and more troubled, until, when the motion of the tide is at its height, it sweeps past at a rate of eleven miles in the hour, boiling up and foaming along like some vast river. Here and there, where sunken rocks lie in the path of the current, the tumult of the water is strongest, while, should a high wind be blowing against the tide, huge waves and sheets of foam are tossed high into the air, and the whole surface of the firth becomes white with breakers. No small vessel can then attempt to force its way through the strait.

32. In other narrow passages between islands where the tide is thrown from side to side against sunken rocks,

or where two opposing currents meet each other, the water forms *whirlpools*. Such are found in the well-known Corryvreckan between the islands of Jura and Scarba on the west coast of Scotland, where the current runs alternately east and west at a speed of eleven miles in the hour. The famous Maelstrom, at the southern end of the Lofodden Islands on the Norwegian coast, is another illustration.

33. The strip of sand, gravel or mud which is alternately covered and laid bare by the rise and fall of the tidal undulation is called the beach. When the tide is at the full, the sea reaches to the upper limit of the beach or high-water mark, when the tide is at the ebb the sea does not come further than the lower limit of the beach

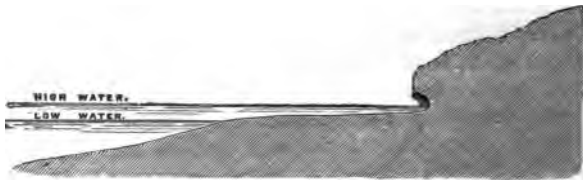


FIG. 20.—Diagram showing the relation of the beach to the lines of high and low water.

or low-water mark. The distance between these two limits must evidently depend partly upon the height of the tides, and partly upon the slope of the beach. It is of course greatest where the largest rise and fall of the tides is combined with the gentlest inclination of the shore.

LESSON XVIII.—THE OFFICES OF THE SEA.

1. In the foregoing Lessons the more important purposes which the sea serves in the general life of the earth have been described. Let us here consider them in a brief summary.

2. (i.) **The sea supplies most of the moisture of the atmosphere.** Were the air overlying the land dependent for its moisture solely upon the evaporation from the land underneath, rain and dew would almost cease, and neither vegetable nor animal life could flourish. The clouds which overspread the sky and discharge their rain-showers upon the ground, the springs which feed the brooks, the innumerable streams which go to swell the bulk of the great rivers, the snows which overspread the higher mountains, the dews which refresh the surface of the earth even in tracts where no rain falls, all come directly or indirectly from the sea. In spite of the enormous volume of fresh water which is continually poured into the sea by innumerable streams, no change of the sea-level is perceptible. That water is evaporated again, rises into the air and is borne along by the winds, till it is condensed and discharged upon the land.

3. The influence of the sea upon the moisture of the air is well shown by the much greater rainfall of maritime than of inland tracts (Lesson X. art. 29). The air lying over the sea is probably for the most part not far from the point of saturation, so that when it moves away landwards it is ready to part with some of its vapour as soon as it meets with a mass of air or of land colder than itself. Near the coast of India the evaporation from the sea is said to amount to about three-quarters of an inch in the twenty-four hours, or nearly twenty-three feet in the year. In the Red Sea a layer of eight

feet of water is believed to be annually driven off by evaporation. A rough estimate gives a depth of fifteen feet as annually evaporated from the trade-wind region of the oceans.

4. In the torrid zone, where the evaporation is greatest, much of the moisture raised from the sea undoubtedly falls back into it in those heavy and continual rains which characterise that belt of the earth's surface. Some of it goes to supply the heavy rainfall of such regions as the Khasi Hills (Lesson X., art. 31), the Himalaya, and the high lands of Abyssinia. Very little of this equatorial vapour, however, can pass over to the temperate regions on either side, because the air which contains it, after discharging heavy showers in the zone of constant precipitation, ascends into the high regions of the atmosphere as a cold and comparatively dry current, travelling away from the equator, and reaching the surface of the earth again, ready rather to take up moisture than to part with it.

5. In Europe the rains are supplied chiefly by the moist winds which drink up the vapour of the warm surface of the Atlantic; west, and especially south-west, winds are wet, east and north-east winds are dry. The vast cauldron of the Indian Ocean furnishes the rains of the south of Asia and the east of Africa. When the south-west monsoon blows, it bears the vapour of that ocean to India and China, and discharges it in deluges of rain. The large rivers of western Africa are supplied by the monsoons, which bear the vapour of the Atlantic into the interior of that continent. In North America the mountainous western sea-board derives its moisture from the moist winds which blow from the south-west across the Pacific Ocean. In South America the south-east trade-winds carry the vapour of the South Atlantic across the continent, until they discharge the last of it on the slopes

of the Andes, and pass over to the Pacific coast as dry, rainless winds.

6. (ii.) **The sea regulates the distribution of temperature.** Its currents carry the heat of the tropics to warm the regions lying towards the pole; while, on the other hand, they bring the cold of the poles to temper the heat of the lower latitudes.

7. This general influence of the sea, well illustrated by the charts of isothermal lines, has been referred to in Lesson IX., art. 15, and the cold and warm currents of the North Atlantic Ocean were given as examples. The heated water of the Gulf Stream turns north-eastward and crosses the Atlantic, spreading out over a larger surface and losing heat as it advances, yet distinctly traceable by its warmth even into the Arctic seas. The average winter temperature of the sea round the British Islands is considerably higher than that of the land. But as already remarked (Lesson XV. art. 2) it is not directly by contact with the land that the sea tempers the climate. It warms the air overlying it, which then passes on and carries the warmth across the land. And here we see the advantage to the west of Europe in that spreading out of the Gulf Stream water which has been referred to. Were this current to pass across the ocean with the same breadth and speed which it has when it issues from the Florida Strait, even though it should retain its highest temperature, it would be too narrow to produce much effect on the climate of Western Europe. But by widening out over a far larger area and moving much more slowly, it exposes a much greater surface to the south-westerly breezes, which, warmed and moistened by it, pass over to Europe.

8. The extent of this amelioration of the climate in Western Europe may best be seen by comparing the winter temperature of places on the same latitude.

Hammerfest, the most northerly seaport in the world, is open even in January, while to the west, on the same parallel of latitude, the east coast of Greenland is covered with snow and ice, and the sea remains frozen in vast floes even during summer. Still further south the harbours of Glasgow and Liverpool are not only never frozen, but enjoy a mean annual temperature of 47° — 51° Fahr. On the corresponding latitudes in North America the coast-line of Labrador is cumbered with ice all the year. Even at St. John's, Newfoundland, which is two degrees further south than Liverpool, the harbour has been closed with ice in the month of June.

9: These contrasts, which are probably the most striking of the kind anywhere to be found upon the globe, do not depend wholly upon the influence of the warm Atlantic current in raising the temperature of Western Europe. They are partly caused also by the influence of the cold Arctic current already described as descending from Davis Strait, keeping close to the American coast and depressing the temperature there. Even as far south as lat. 44° , that is, on the same parallel as the south of France and north of Italy, the water which washes the shores of Nova Scotia is not more than five or six degrees above the temperature of thawing ice.

10. In the southern hemisphere, where sea so largely predominates, the climates are distributed with considerable regularity according to latitude. In the northern hemisphere, however, where the equalizing influence of the sea is opposed by the existence of large masses of land alternately heated and chilled, no such regularity exists. And yet even there the tempering action of the sea makes itself sensible along the borders of the land, where neither are the winters so intensely cold nor the summers so hot as they are inland. Hence a continental climate is one of extremes—great heat in

summer, severe cold in winter ; an insular climate is more equable—a distinction well brought out by comparing the summer and winter temperatures of Ireland with those of regions on the same latitudes in the midst of the continent, such as the heart of Russia. (See Lesson XXXI.)

11. (iii.) The sea wears away its shores, and thus tends to reduce the area of the land. No one can watch the action of the waves (Lesson XVII. art. 14) without being well able to realize how powerful it must be in the destruction of even the most rocky coast-line. The mere weight of so many tons of water as are suddenly thrown by a huge wave against the land during a storm suffices to loosen and detach fragments of rock. In a subsequent Lesson we shall trace how even the hardest rocks are apt to be split up by the action of the atmosphere. The waves, therefore, are aided in their work of derpolation by the cracks and lines of weakness which the atmosphere has prepared for them.

12. But the fragments of stone removed in this way become powerful implements in the further destruction of the coast-line. They are caught up and hurled forward by the waves, battering down the cliffs, and at the same time being further broken up themselves. Ground to and fro by the advance and recoil of the waves, they assume the smoothed and rounded forms so familiar in the gravel and shingle of the sea-shore. But even then their demolition continues, for they are still rolled backwards and forwards, until they become at last mere sand and mud, in which condition they may be swept away out to sea, and laid down in the quiet depths there.

13. There is yet another way in which the sea loosens and demolishes the rocks on its margin. Those who live at a seaport on any coast exposed to storms must often have noticed that after a violent gale portions of the

solid masonry of the pier-walls and breakwaters have been started from their places. Even with the utmost care in the building and repair of these works, it is hardly possible to avoid injury of this kind ; and unless constant vigilance is used in cementing every crack and opening in the masonry, the destruction of the pier or bulwark may be rapid and complete. Now it is not by the mere weight of water thrown against the building, nor by the impact of any blocks of stone hurled forward by the waves, that the tones are started out of a piece of well-built masonry. When a large wave falls upon such a surface, the air in every cranny of the wall is driven inward. When the wave recoils again a great suction takes place behind it, and the compressed air rushes out of the masonry. Moreover, where the water is driven in with immense force so as to fill crevices and openings within the masonry, it exerts upon their walls the same pressure as the wave of which it is for the moment a part, on the principle of the hydraulic press. After a time some weak part of the work is discovered ; one or more stones are loosened and then pulled out at the recoil of some large wave, so that a breach is made, which, if the storm continues, may be rapidly widened. The same kind of influence is exerted upon cliffs. Every mass of solid rock is traversed more or less by crevices which afford scope for the combined action of compressed air and waves. When the sea has drilled a cave at the base of a cliff, the action of the air alternately compressed and released as the waves rush in and out, loosens the rock at the end and roof of the passage, so that if the cliff be not too high, or the rock too solid, an opening is actually made between the end of the cave and the ground behind the top of the cliff. Through this opening, called a *blow-hole*, the sea when in storms sends clouds of spray. Sometimes the mouth of the opening lies a

short distance from the edge of the sea. It is singular to find on a hill-side, or in the middle of a field or moor, a huge cauldron-like hole, with the sea ebbing and flowing at the bottom.

14. Since even cliffs of the most solid rocks are worn down by the incessant attacks of the waves, we may readily believe that where the material is of less durability the rate of destruction must be still more rapid. On some parts of the eastern shores of England, where the waves beat against cliffs of crumbling clay, the rate of waste



FIG. 21.—View of the sea-cliffs south of the river Tyne (magnesian limestone worn away by the waves, and isolated fragments left). See also Fig. 76.

occasionally reaches even to as much as the loss of a yard of land in the year. The sites of some of the former ports and towns of Yorkshire now lie beneath the restless waters of the North Sea a mile or more from the present shore.

15. We must not suppose, however, that the sea is everywhere cutting away the margin of the land. It does so chiefly on those coasts which lie exposed to prevalent

storms and where the form of the shore allows the waves to act with effect. But on every coast-line there are sheltered places where the waves either do not encroach at all or at least do so very slowly. For example, if the sea is shallow for a long distance outward, with a flat, sandy beach and a low coast-line, the largest waves may roll in and spend their force merely in rushing along the flat beach, and may do no damage to the shore. But in many places the sea, instead of removing material, casts it up on the land. The finer sediment swept away from one part of the coast and carried off by the currents is sometimes thrown ashore again at no great distance. When that takes place, the land may gain there as much, or nearly as much, as it loses from that part of its margin whence the sand and mud are derived. In this way the flat shores of Lincolnshire encroach upon the sea, because they continually receive some of the material removed from the coast of Yorkshire. In many parts of the world, as along the coast of North America from the borders of Mexico to those of Virginia, and on both sides of the Madras Presidency, so much mud and sand are brought down by the rivers from the land that bars are formed, by which the land behind them is protected from the waves.

16. (iv.) The sea receives and preserves the materials out of which new land will in course of time be formed.—We shall see in a later Lesson how the surface of the land undergoes continual decay. Even its hardest rocks crumble down, and their remains are swept out by brooks and rivers to the sea. In the quiet depths of the ocean-basins the sand and mud derived from the decay of the surface of the land are laid down. There they remain undisturbed, slowly accumulating and spreading, until some future movement of the earth's crust shall raise them above the level of the sea.

17. If you look into the rocks of which the present dry land is formed, you will find many of them to contain abundant remains of coral, shells, fishes, and other marine creatures. These animals lived in the sea, they were inclosed in the sediments of the sea-bottom and preserved there ; these sediments, slowly consolidated into rock, have been upraised so as to form the land on which we live. What has happened so frequently in the past will probably often take place in the future.

18. Besides the deposits formed of the fine sand and mud carried away from the land, vast spaces of the bottom of the ocean are likewise covered with a slowly-gathering deposit derived from the remains of minute forms of life as described in Lesson XIV. So far as we can tell, some of the rocks of the land, the chalk of England and France for example, have been laid down in much the same way, by the gradual accumulation of the tiny shells of foraminifera and other inhabitants of the sea.

19. It is evident that while the surface of the land is subject to continual decay and removal, that portion of the solid earth which lies below the comparatively shallow depth (a few hundred feet at most) to which the influence of waves extends, is preserved. By far the largest part of the destruction which the sea works is done between tide marks. (See Fig. 21.) Waves, as we have seen, do not stir the deeper parts of the ocean, which are therefore left in undisturbed repose. When we reflect how large a proportion of the earth's surface is covered with sea, we perceive that although the waves are always beating against the shores, the destruction which they effect takes place only on some parts of the mere margin of the land, while the vast spaces under the sea are not only preserved from waste, but are continually receiving fresh deposits. So that if we consider the action of the sea as a whole, we find it to be far more preservative than destructive.

CHAPTER IV.

THE LAND.

LESSON XIX.—CONTINENTS AND ISLANDS.

1. From the two envelopes of air and sea in which our earth is inclosed we now pass to the consideration of what lies beneath or within them—the solid mass of the planet. One of the first points to be noticed here, is that while both air and sea admit of being penetrated and examined all over the world, only a very small proportion of the solid framework of the globe rises out of the sea to form land, and of that comparatively insignificant part little more than the mere surface or skin is known to us. We cannot at will pierce the land as we can sound the sea. Evidently, therefore, all that can be discovered by actual observation regarding the constitution of the solid globe must be derived from what may be seen at or near the surface of the land.

2. What then, to judge from all that can be learnt from its visible parts, is this solid globe on which we live? We are familiar with the earth and soil of the surface, and we have seen the rocks beneath it—in one country clays and limestones, in another sand and gravel, in another sandstones and shales, in still another granites and other crystalline formations. How have these various materials arisen? Can we learn anything regarding their

origin, and do they in any way throw light upon the history of the earth? Why, for example, should the land rise into lofty ridges at one place and sink into low plains in another; and have these ridges and plains been always the same as now?

3. No matter in what part of the world we may live such questions as these must often arise in our minds. The more prominent the features of the land, the more of course do these questions press for an answer. In a mountainous country, for example, or one from which a distant range of mountains can be seen, the high peaks and ridges into which the surface of the land has been thrown form so striking a contrast to the level plains below that we cannot help at times asking ourselves how and when these vast elevations took their rise. We watch the clouds on the far summits as they gather into thunderstorms, or dissolve into showers of rain among the valleys, or powder with fresh snow the white shoulders and crests of the mountains. There seems to be a constant turmoil in the air of these heights. What effect must it have upon the forms of the mountains? Do these ever recurring tempests, these frequent rains and snows, which fill the torrents and feed the great rivers, leave the outlines of the mountains unchanged?

4. But even where the landscape is least marked by any striking feature an observant eye cannot fail to notice much that is suggestive of inquiry. No scene for instance can be more monotonous than that presented by the wide plains at the mouths of large rivers. And yet when the mind is once opened to perceive the interest of such subjects as these, can the nature of the soil of these flat lands escape notice, and fail to be connected with the mud brought down by the river? Whence does the mud come? How long has the river been busy in strewing it over these plains?

5. Such are some of the questions which are now to come before us in the Lessons contained in this chapter. In entering upon them we shall do well to remember that as far as possible we should try to find illustrations of them in the district where we may chance to live. We shall not be able to do so in all, nor perhaps even in most cases. Yet the search for examples will always be of the utmost use in making us better acquainted with the physical geography of our country, and in accustoming us to the observant use of our eyes. The features of a mountain-chain, a plain, a table-land, a watershed, a spring, a brook, a river, a lake, and of every other part of the land, should be studied on the ground, wherever we can find examples of them. In all such cases, too, we should not rest content until we have put to the proof the statements regarding these features in the Lessons, and seen how far they are borne out or can be extended by what we see in nature for ourselves.

6. As the land consists of those portions of the solid globe which happen to project above the level of the sea, its distribution shows the position of the large wrinkles into which the surface of the planet has been folded, while its form and its materials must evidently furnish our chief sources of information regarding the composition and the history of the general mass of the earth. First, the distribution of the land over the globe may be noted; then its horizontal outlines or coast-lines where it meets the sea; next, its vertical outlines or relief, as it rises up into the air.

7. **Distribution of the Land.**—The total area of dry land on the surface of the globe has been computed to amount to fifty-two millions of square miles. It was stated in Lesson V. that by much the larger part of this land lies in the northern hemisphere. Within the Arctic Circle (Plate I. fig. 5) it forms an almost

continuous ring round the north polar regions, whence it extends southwards in long irregular masses which taper away into points (Lesson V. Art. 10). Each of these masses of land or continents may be looked upon as consisting of a northern and southern division. In the western hemisphere they are united by a narrow neck of land or isthmus. In the eastern hemisphere Europe and Africa are separated by the narrow water-channel or inland sea called the Mediterranean. Europe and Asia form indeed one united and continuous continent which is prolonged southwards by Africa on the one side and by a vast archipelago, or clusters and chains of islands, into Australia on the other. In a general view the land may be regarded as forming three pairs of continents : the most regular being North and South America, and the least regular being Asia and its island prolongation into Tasmania.

8. Looking now at the position of these continental masses we may observe that there is among them a general tendency to run in a north-west and south-east direction. This is specially noticeable in the case of America. It is seen, too, in the bend of the islands from the south-eastern margin of Asia to Tasmania and New Zealand. Even in the European and African pair of continents it may be traced, if we connect Greenland, by Iceland, the Faroe and British Islands, with the main mass of Europe, for we can then see how an uneven and broken ridge of land extends from the polar circle south-eastwards to the Cape of Good Hope. Hence the shape of the Atlantic ocean is determined as that of a deep, long, narrow channel or trough between this ridge of land on the east and the American ridge on the west. Owing to the expansion of Asia to the north-east, the American and Asiatic continents almost meet at Behring Strait, thus inclosing the vast Pacific basin on the north,

9. But in order to have a clear notion of the main heights and hollows on the surface of the solid globe we cannot merely consider the trend of the continents. Besides these chief masses of land innumerable smaller patches project above the surface of the oceans in the form of scattered groups or clusters of islands. As these are only the tops of submarine ridges or mountains their position serves to indicate the direction of the elevations of the surface of the earth which lie under the sea.

10. In the Pacific Ocean a scattered series of islands begins to the west of the coast of Chili, and stretches by the Society and Marquesas Archipelagos north-westwards through the Sandwich group, and thence by the Bonin Islands towards Japan. This long line sends off a well-marked branch along the equator, taking in the Marshalls and Caroline Islands, and trending north to the southern end of Japan. A third branch strikes off from the Marshalls Islands to the Friendly Islands, and thence to New Zealand. A parallel chain includes the Phillipine and Molucca groups, Papua, and many small islets which dot the ocean as far as New Caledonia and the Fiji Islands.

11. When these lines are traced on the map the vast basin of the Pacific is seen to be traversed by one main ridge stretching from South America to Japan, with several diverging branches. It is, of course, the mountain tops of these ridges which form the islands. As a rule these islands do not rise much above the sea-level; the loftiest summits are in Hawaii, where one reaches to 13,760 feet. The highest point of Tahiti is about 7,000 feet. Recent soundings across the Pacific show an average depth of between 2,000 and 3,000 fathoms, so that some points on the Pacific ridge rise as much as 30,000 feet above the general level of the Pacific floor, that is, rather more than the height of the highest mountain above the sea.

12. These ridges, however, it must be remembered, are not continuous submarine mountain-chains. They are cut through by broad and deep hollows, which are even sometimes much deeper than the general level of the floor of the ocean. Thus the deepest sounding yet obtained in any part of the sea (4,475 fathoms or 26,850 feet) was taken in the Caroline Archipelago.

13. In the Atlantic basin (Lesson XII. Art. 15) scattered islands afford similar indications of the general direction of the ridges upon the solid globe. North of the equator a submarine bank crosses the ocean from Cape Verde to Newfoundland; its highest summits rise above the sea as the Cape Verde and Canary Islands, Madeira, and the Azores; other elevations which do not reach up to the sea-level form submarine banks, like the Great Bank which lies off Newfoundland. South of the equator a north-westerly line of submerged ridge is indicated by the islands of St. Paul, Ascension, and St. Helena, while another seems to lie further south with the island of Tristan da Cunha marking its highest point. Again, from the southern promontory of the African continent a long ridge must stretch under the sea in a south-easterly direction into the frozen lands of the southern pole. Its direction is marked by the Prince Edward, Crozet, Kerguelen, and Heard Islands.

14. The Indian Ocean presents the same kind of evidence that the solid earth has an uneven surface below the sea. A long line of islands and submerged banks extends from Madagascar and Bourbon northwards to the equator. Another similar line starts from opposite the Malabar coast and runs through the chain of the Laccadive and Maldive islands for seven or eight degrees south of the equator.

15. To be able to follow these descriptions you should take a good map of the world or a school-globe and trace

the several lines upon it. In this way you will best obtain a general idea of the position of the chief ridges and hollows on the surface of the earth, and see in how curious a manner that surface has been wrinkled. Turning now more particularly to the land itself, you observe it to be defined by two limits, one where it meets the sea—that is, its coast-lines, the other where it is surmounted by the air, giving its vertical relief or contour.

16. Coast-lines of the Continents.—Though the limit between land and sea is always sharply drawn, the map of the world shows how varied and irregular it is. But even the largest and most detailed map can show only the more prominent curvings and windings of the coast-line. When such a map is carried along the edge of the land and compared with the actual boundaries of the sea, many little projections and recessions of the coast are seen which have not, and from their small size could not, find a place upon the map. Comparatively seldom does the coast continue in a straight line for more than a brief space. It curves to and fro, jutting out into *headlands*, *capes*, or *promontories*, and retiring into wide *gulfs* and *bays*, or into narrower *firths*, *creeks*, and *inlets*.

17. Besides these abundant changes in its horizontal extension or plan, the coast-line of the land presents continual variety in its vertical elevation or relief. Sometimes, for example, the land rises abruptly out of the sea in steep and lofty cliffs against which the waves are ever beating. Elsewhere we find the shore shelving gently into the water. It has been observed that where the coast-line is precipitous the water is usually deep (Fig. 22), the land descending beneath the sea with the same abruptness as it rises out of it, while on the other hand a low shore commonly indicates shallow water (Fig. 23). This relation between the form of the ground and the depth of the water is so general that in navigation a vessel will usually

be kept well away from a low shore, though brought without hesitation close under a high headland.

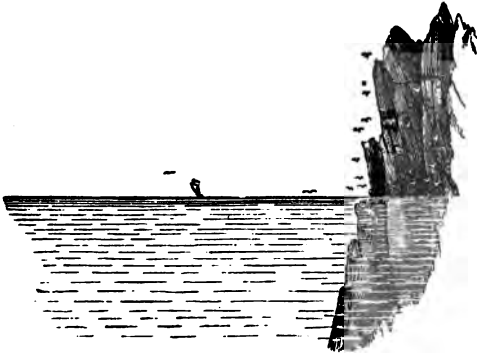


FIG. 22.—Steep shore descending abruptly into deep water.

18. If you happen to live near the sea it may be worth your while to examine the coast-line of your own neighbourhood. You will find as a rule that the projecting and higher



FIG. 23.—Low shore shelving into shallow water.

parts consist of harder materials, the receding and lower parts of softer materials. This is true also of the larger irregularities of the coast all over the world. Wherever you see upon the map a bold headland standing well out to sea you may infer that it consists of some hard rock which

rises up in cliffs and can offer a stout resistance to the waves. Wherever, on the other hand, you observe the land to retire into wide unindented bays you may suppose it to be low and to be formed of soft clays or other materials which have yielded more easily to the encroachments of the sea. In the great majority of cases your inference would be correct.

19. Compare now the coast-lines of the three pairs of continents and you will not fail to observe a remarkable contrast between those of the northern and southern divisions. The northern continents are marked by coast-lines so indented that the oceans penetrate far inland in many bays, creeks, inlets, firths, and inland seas. The southern continents, on the other hand, are distinguished by long monotonous stretches of shore, unbroken by bay or creek. The following calculation has been made of the relative proportion of coast-line to extent of surface among continents.

| | | | | | | | |
|---|-----|---|---------------------------------|-----|-----------------------|---|---|
| FIRST PAIR— | | | | | | | |
| N. America | has | 1 | geograph. mile of coast-line to | 265 | sq. miles of surface. | | |
| S. America | „ | 1 | „ „ „ „ | 434 | „ | „ | „ |
| SECOND PAIR— | | | | | | | |
| Europe | „ | 1 | „ „ „ „ | 143 | „ | „ | „ |
| Africa | „ | 1 | „ „ „ „ | 895 | „ | „ | „ |
| THIRD PAIR— | | | | | | | |
| Asia includ- ing the Islands ...) | „ | 1 | „ „ „ „ | 469 | „ | „ | „ |
| Australia | „ | 1 | „ „ „ „ | 332 | „ | „ | „ |

20. It will be seen that the contrast is most marked between Europe and Africa. The former continent has six times more coast-line in proportion to its superficial extent than Africa has. There can be little doubt that this contrast has had a powerful influence on the civilisation of the two continents. In the one case abundant bays and arms of the sea have offered ready facilities for discovery, conquest, and commerce. Nation has been

brought into contact with nation ; the arts of peace and war have spread into every country ; no region however remote has lain wholly beyond the reach of communication with the rest ; and hence all the communities of Europe, whether slowly or rapidly, have been carried forward by the general progress of mankind. But with this abundant intercourse and steady advance contrast the condition of Africa, with its vast expanse of hitherto inaccessible and unopened territories, peopled by races having no intercourse with other tribes save their immediate and usually hostile neighbours, and remaining from time immemorial in the same stationary condition of barbarism. Doubtless the noble rivers and lakes of Africa will eventually become highways of civilisation into the very heart of the continent, and be made to serve the purpose which in old times was accomplished in Europe by the far-reaching arms of the sea.

21. General Relief of the Continents.—The horizontal outlines or coasts of the land are greatly surpassed in variety and interest by its vertical outlines or relief. It is easy to show the former upon a map, but much more difficult adequately to express the latter. Hence even the best maps convey a most imperfect idea of the external aspect of the land. Leaving details for the next Lesson we may notice at present some general features which are illustrated in all parts of the globe.

22. Even the loftiest peak of land rises but a small way above the average level of the earth's surface. Think, for instance, of the proportion between the height of even the highest summit of the Himalaya Mountains and the total bulk of the earth. The top of Mount Everest stands 29,000 feet, or $5\frac{1}{2}$ miles, above the level of the sea, an enormous height when compared with the little undulations of the plains or even with the heights of many hilly countries. And yet it is only $\frac{1}{1438}$ th of the polar

diameter of the earth ; so that on a globe of ten feet in diameter the highest mountain in the world would be represented by a little projection only one-twelfth of an inch in length.

23. If, then, the most stupendous mountains are really so small, it is evident that the whole mass of dry land must form but an insignificant part of the entire bulk of the earth. For by far the largest part of the land is not mountainous. Calculations have been made regarding the probable average height of the land, if the mountains could be levelled down and the valleys filled up so as to reduce the whole to one general level. This, exclusive of Africa and Australia, has been estimated to be about 925 feet. The continents, however, differ much from each other in this respect. To Europe, for example, the great geographer Humboldt assigned an average elevation of 676 feet, while that of Asia he computed to reach 1,132 feet, and that of Africa has been put down at 1,800 feet. From these estimates it is found that even vast mountain-chains do not form such a large proportion of the mass of the land as they might be supposed to do. It has been calculated, for instance, that if the whole of the Alps could be ground down and spread over Europe, the surface of that continent would not be raised more than about 21 feet.

24. In no respect does the land stand out in more marked contrast with the sea than in the endless variety of its relief. While the surface of the oceans is one vast level, that of the land exhibits every contrast of form from the jagged peaks and precipices of the mountains down to the level flats of the meadows and plains. We are apt to see no order in this variety. Mountain and valley seem to succeed each other, and to change their shapes at random. And yet an attentive study of them shows that they have not been stuck down on the earth's

surface at mere hazard, but have an intelligible meaning and history, and a traceable connection with each other.

25. First of all it is to be noticed, that each of the continents is traversed by a line or axis from which the ground slopes on either side to the sea. This axis does not necessarily coincide with the highest parts of the land : these may lie now to the one side of it, now to the other ; but it is the line of average highest elevation, as is shown by the way in which the rivers flow from each side of it. Nor is the axis placed down the centre of the continent ; more usually it is much on one side. In America, for instance, it lies close to the Pacific sea-board. In Europe it runs from Cape Finisterre through the chain of the Pyrenees, the Alps, the Carpathians, and the Caucasus to the shores of the Caspian Sea.

26. The position of the axis determines the general slope on either side. When it runs along the centre of a continent the angle of slope on either side will be nearly the same. When it lies close to one side the angle must be higher on that side than on the other. Each continent or country, with an axis lying far from the true centre of the region, has therefore a short and steep slope on one side and a long and gentle slope on the other. South America presents the most remarkable example of this feature. The axis, with an elevation of perhaps 8,000 or 10,000 feet, runs down the line of the Andes at a distance of only from 50 to 100 miles from the Pacific, but 2,000 miles from the Atlantic Ocean. On a much smaller scale the same character is shown by Scandinavia, where the axis lies close to the Atlantic, and where consequently the ground descends rapidly into that ocean from the snow-fields of Norway, but slopes gently eastward across Sweden into the Gulf of Bothnia. In the British Islands, too, the western slope is short and steep in the northern

half of Scotland and in Wales, while the eastern slope is long and gentle.

27. But observe how the axial line runs through the great continents. On the whole it is so placed that the steep and short slope shall face the larger ocean. Look again at America, with its vast line of heights towering steeply out of the Pacific on the one hand, and inclining gently for hundreds of miles into the Atlantic on the other. On the opposite side of the Pacific the lofty highlands of Asia rise as a great mountainous girdle which stretches from Kamtschatka to Arabia, and is thence prolonged through Africa to the Cape. The slope to the Pacific and Indian Oceans from that range of high ground is comparatively short, while on the opposite side to the Arctic and Atlantic Oceans it is gentle and long. It would seem that on the average the gentler slope of a continent is four or five times longer than the steep one.

28. Moreover, you may notice that subordinate to the main axis a minor axis or ridge is apt to occur on the other side of the continent fronting the smaller ocean. This likewise may be seen most conspicuously in America. The line of the Alleghany and White Mountains lies near the Atlantic border in North America, and the Sierras of Brazil rise close to the sea in South America. These dominant lines cannot be mere accidents; they must be due to the same causes which have ridged up the whole of the solid land.

29. Again, it is evident that when the continents are flanked with high ground on either side they must contain wide plains, or at least vast spaces of lower ground, in their centre. America illustrates this aspect by the great plains of Canada, of the Mississippi, Orinoco, Amazon, and La Plata. The greater part of Europe is a low plain bounded by the chain of the Alps, Carpathians,

and Caucasus on the south, the Urals on the east, and the high grounds of Scandinavia and Britain on the west.

30. In most cases the upheaval of the continents has been accomplished in such a way that these central plains or lower grounds slope seawards. The water which falls upon them, therefore, is not retained there, but flows off to the ocean. In the crumpling and erosion of the surface of the earth, however, depressions have been formed upon each of the continents below the level of the surrounding country. Into these the water flows, and fills them up till they overflow and discharge their surplus by means of streams. Such are the numerous water-basins known as lakes. But in some regions the water cannot find an outlet or any escape except by evaporation. On the high central platform of Asia a desert region of this kind, full of salt-lakes, lies inclosed between the Hindu Kush and Thian Shan Mountains, and stretches through Turkestan and Mongolia for a distance of about 2,000 miles. In North America a much smaller basin, shut in among the ridges of the Rocky Mountains, receives its drainage into lakes without outlets, of which the Great Salt Lake of Utah is the most important. On the south-east of Europe the depression of the Caspian descends far beneath the level of the sea; even the surface of that inland sheet of salt water being about eighty-four feet below the general level of the oceans. The surface of the Dead Sea, which fills a small detached hollow shut in by high grounds, is 1,298 feet below the level of the Mediterranean.

We have now traced the position and broad general outlines of the great ridges on the surface of the earth. The following Lesson deals with some of the more prominent features of the surface of the land.

LESSON XX.—THE RELIEF OF THE LAND—MOUNTAINS,
PLAINS, AND VALLEYS.

1. From what was stated in last Lesson it may be inferred that although the surface of the land is very uneven, rising sometimes, as in the Himalaya, five-and-a-half miles above, and sometimes, as in the hollow of the Caspian (3,000 feet deep), sinking nearly three quarters of a mile below, the sea-level, its inequalities are not altogether disposed at random. And this is really the case. Each continent has its own system of heights and hollows, and these are grouped in relation with the general axis.

2. The main features of the continents are the lines of mountain chains. It is along these lines that the elevation of the land has been greatest. They are, as it were, the crests of the great waves into which the surface of the solid globe has been thrown. They govern the trend of the chief valleys, they determine the position of the plains, they regulate the climate, the winds, rains and rivers, of the land. They ought first to be considered, therefore, in any description of the general aspect of a country: we may then pass to the valleys and plains.

3. **Mountains.**—The term 'mountain' is rather vaguely used to describe any large and lofty elevation rising conspicuously above the region which surrounds it. Sometimes a single conical mountain towers above a plain or shoots up from the sea. Solitary cones of this kind are usually volcanoes, such as Etna, Vesuvius, and the Peak of Teneriffe. More frequently a mountain is merely a more prominent and elevated part of a long lofty ridge, and is joined at the sides to other similar mountains, into which that ridge is divided by the valleys which cross it. Such a connected series of mountains is called a

mountain-range. It often happens that two or more such lines of mountain run parallel with each other as one vast and continuous mountain-chain, or mountain-system. Let us consider for a little the aspect of some leading mountain-chain as an illustration of this feature of the surface of the land.

4. The Alps of Europe have been so long familiar to mankind that their name has passed into use as an appellation for the main mountain-system of any country. To one who approaches it from the north, that noble chain of mountains only reveals its characters step by step. Crossing the plains of France or Germany into the district of the Jura, one finds the ground begin to rise and sink in long ridges and valleys, which follow each other in parallel lines, like the undulations on the surface of a great ocean. The smooth valleys are green and fertile ; the ridges, for



FIG. 24.—Section of one side of a Continent.

the most part also smooth, bear their slopes of rich pasture and shaggy wood, but often show along their crests long narrow valleys or amphitheatres bounded by walls of rock. Here and there deep transverse gorges cross the ridges and carry the drainage out into the plains beyond. Threading his way through these transverse valleys, the traveller finds the ridges on either hand to grow higher, and the hollows between to become steeper and deeper, until he reaches one of the last and loftiest of these outer elevations. From its top he can see on the one side, as in a map, the succession of wave-like folds of hill and valley across which he has journeyed ; on the other, beyond a broad plain which lies at his feet, his eye may take in the whole

panorama of the Alps—a vast array of mountains crowding behind each other along the sky-line, their higher slopes and crests white with snow, and hardly at times to be distinguished from the white clouds which rest upon them.

5. On a nearer view this great line of heights is found to be flanked with minor ridges, rising into loftier and bolder masses than the Jura, as they approach the main mass of the mountains, but still retaining their covering of forest, save on the steeper or more rugged slopes, where the naked rock forms broken crags or steep bare precipices. Beyond these outer bulwarks rise the central ranges, with flat meadows and corn-fields or deep still lakes at their base, dark pine-forests and green hollows of pasture down their sides, but rearing their shoulders above the limit at which tree or grass can live, and raising their heads far up into the region of eternal snow.

6. It is not until we have gained some commanding point of view above the snow-line that we can form any proper conception of the vastness of the scale on which these mountain-ridges have been piled up. The point which we have reached after some hours of laborious climbing seemed the very summit of a mountain when we stood in the valley far below. And yet we now find a vast sweep of much higher ground all around. Peak rises behind peak, crest above crest, with infinite variety of outline, and with a wild grandeur which often suggests the tossing and foaming breakers of a stormy ocean. Over all the scene, if the air be calm, there broods a stillness which makes the majesty of the mountains yet more impressive. No hum of bee or twitter of bird is heard so high. No brook or waterfall exists up among these snowy heights. The usual sounds of the lower ground have disappeared. Now and then a muttering like distant thunder may be caught, as some loosened

mass of snow or ice falls with a crash into the valleys ; or the wind brings up from below in fitful gusts the murmur of the streams which wander down the distant valleys.

7. It is seldom that these high summits are free from clouds during at least some part of the day. They rise so far into the air that often for weeks together they are wholly wrapped in mist. The traveller who climbs them finds sometimes that the higher points rise above the cloud-belt into the clear upper sunshine, and he may even now and then look down upon a thunder-storm raging along the lower crests of the mountains and sending torrents of rain upon the valleys, while he has a clear blue sky above him.

8. Another feature which strongly impresses the mind is the contrast between the climate and vegetation of the plains and those of the mountains. Among the Swiss valleys, for instance, corn-fields, gardens, orchards, and vineyards cover the lower grounds, while thick woods climb the slopes. As one ascends, the plants of the valleys gradually disappear, others of species peculiar to the higher and colder air of the mountains take their place ; the pine-trees, after a time, grow fewer, and at last cease, until after further ascent the line of permanent snow is reached, at a height of about 9,000 feet above the sea (Fig. 78). Beyond that line the climate has quite an Arctic character, with hardly any traces of life, either vegetable or animal. Within the torrid zone the mountain-contrasts are still more marked. The sultry plains of the Ganges, for example, show a luxuriant tropical vegetation ; on the middle slopes of the Himalaya the climates and the vegetation resemble those characteristic of the temperate zone, while the higher parts are like polar latitudes in their temperature and partly in their plants.

9. Every mountain chain is traversed by a system of valleys. Those which run along the line of the chain and

separate its ridges, or ranges, are called longitudinal; those which run across the ridges and divide them into independent mountains are termed transverse. Thus the Swiss Alps contain the two parallel lines of the Bernese Oberland and the Pennine range, separated by the longitudinal valleys of the Rhone and Rhine. The former of these ranges is cut through by many transverse valleys and passes, such as the Rawyl, Gemmi, and Grimsel, between which rise massive snowy mountains. The latter is likewise deeply trenched by valleys and passes, separating some of the monarchs of the Swiss mountains—Mont Blanc, the Matterhorn, Monte Rosa, and the St. Gothard.

10. The direction of the great mountain-chains of the globe has evidently a close connection with that of the continents. Of this connection America, which in many respects is a typical continent, furnishes an admirable illustration. The long, continuous line of mountain-chain which extends from the southern spur of the Andes to the most northerly swell of the Rocky Mountains, a distance of some 9,000 English miles, coincides with the general trend of the continent, and forms the axis or back-bone of that vast tract of land. This main line of elevation in the New World runs in a general north-west and south-east direction. In the Old World a vast mountain system, with many branches, stretches from the north-east of Asia across the centre of that continent and the south of Europe to the north-western headlands of Spain—a distance of about 12,000 miles. Here the general trend is on the whole east and west. These are the two great mountain ridges of the world.

11. We shall return to the structure and origin of mountains in a later Lesson (Lesson XXIX.) after the nature and arrangement of the materials of the solid earth have been considered. Meanwhile we have to

consider some of the other leading forms of outline which the land presents.

12. Plains.—The chief flat lands of the globe lie between parallel mountain ridges ; to a smaller extent they occur as narrow belts, or fringes, flanking the seaward slopes of mountains or of elevated tracts. Take again the American continent by way of illustration. In North America a vast plain stretches between the Rocky Mountains on the west, and the Alleghany and White Mountains on the east from the Gulf of Mexico up to the Arctic Ocean. Of course this wide expanse of land is not a mere dead level. It abounds with ridges of low hills and lines of valley, but neither its heights nor its hollows are sufficiently marked to destroy its character as a plain. The enormous territory watered by the Mississippi and its tributary rivers forms the southern and fertile half of this plain, and is so flat that vessels can sail into it for a distance of about 4,000 miles from the sea. For thousands of square miles its gently undulating surface is covered with grass. These tracts west of the Mississippi are known as prairies or savannahs.

13. More than half the surface of Europe is a vast plain bounded by the heights of Scandinavia, Scotland, and Wales on the north-west, by the ranges of the Urals on the east, and by the chains of the Pyrenees, Alps, Carpathians and Caucasus on the south. From the west of England the plain sweeps eastward across the north of France and Germany, spreads out over most of Russia, and sinking into the depression of the Caspian Sea, passes into Asia, where it bends northward between the Ural and Central Asian mountains, covering there a belt of territory 1,000 to 1,500 miles broad, and nearly 4,000 miles long. In the "Black Lands" of the Russian plain, the soil is a dark rich fertile loam yielding abundant corn-crops. Further south, where it sinks below the sea-level towards the

Caspian, it becomes full of salt and brackish lakes, and assumes a bare and barren aspect. Round the Caspian Sea and northward into Asia the clayey or sandy soil is covered with grass in spring, but is dried up and withered by the scorching summer, and frozen by the severe winter. These tracts are the Steppes of Russia. Within 400 or 500 miles of the Arctic Ocean, trees and grass disappear from their surface and are replaced by moss; the soil is frozen to a considerable depth and only thawed for a few inches down in the height of summer. These northern wastes are the Tundras of Siberia.

14. Some plains are so placed as to remain mere bare sandy wastes, without vegetation and incapable of ever being inhabited. The north of Africa, for example, presents to us the widest barren region or Desert on the surface of the globe, that of the Sahara. This is a long broad plain, rising here and there into hills and tablelands; but at its eastern end actually sinking to a depth of from 100 to 150 feet below the sea-level. Where springs come here and there to the surface, little spots of vegetation called "*Oases*" make their appearance, but with these exceptions the vast plain is a region of dry barren shifting sand, where no rain falls, and where the sun's heat is more intensely felt than anywhere else in the world.

15. In the Sahara Desert, as well in many other plains both in the Old World and in the New, traces of shells are found belonging to species which must have lived in the sea. It can be shown that they have not been transported from the sea to where they now occur, but have lived and died there. Hence they serve to prove that these wide flat or undulating tracts of country must at one time have lain below the waters of the sea. This is an interesting fact which will be afterwards again referred to, in its bearings upon the origin of the present contour of the surface of the land.

16. The tracts of comparatively level land are not always low plains. Sometimes they lie at a considerable elevation. When their height exceeds about 1,000 feet above the level of the sea they are usually called **table-lands, or plateaux**. They undulate like plains into heights and hollows, are sometimes deeply trenched by valleys and river-gorges, and here and there rise into lofty ranges of hills.

17. A table-land may be inclosed between parallel ranges of mountains. Of this arrangement the most stupendous example is the vast plateau of Central Asia, 10,000 feet above the sea, with its mountain barriers of the Thian Shan on the north and the Kuen Lun and Himalaya on the south. Another remarkable instance is the "Great Basin" of North America between the Colorado and Columbia valleys—a table-land about 5,000 feet high inclosed between branches of the Rocky Mountains.

18. Other table-lands rise out of the sea or from low plains bordering the sea. The African continent, for example, is one vast table-land with but a narrow fringe of low plain skirting its shores. Spain and Portugal form another example of a table-land rising steeply from the sea. This well-marked peninsula consists of a plateau about 2,600 feet high, crossed by five hill-ranges (in Spanish *Sierras*) and trenched more or less deeply by river-valleys—the Douro, Ebro, Tagus, Guadalquiver, and others.

19. Here and there a table-land has been so extensively worn down into valleys that its former character becomes almost lost. Scandinavia is an illustration of this result. The central parts of the table-land still remain—a vast undulating plateau between 4,000 and 5,000 feet high, but the outer portions are so traversed by lines of valley as to be cut into narrow ridges. The Highlands of Scotland have been still more thoroughly trenched by valleys until

the original table-land can scarcely be any longer traced (See Lesson XXIX.).

20. Having noted now the leading forms of surface by which the land is marked, we are led to ask whether any explanation can be given of them. How have the vast mountain-chains been formed? Have they always existed, and have their crests and pinnacles, their precipices and ravines and valleys, always been as we see them now? An answer to these questions can be given with considerable confidence and fulness, but to be able to understand it we must attend with some care to two aspects of the physical geography of our globe. 1st. The nature of the earth's interior and the reaction of the interior upon the surface; and 2nd, the circulation of water upon the land and the other external processes by which the surface of the land is affected. The next Lessons will be devoted to these two subjects, and we shall then return to the question as to the origin of the present features of the land.

LESSON XXI.—THE COMPOSITION OF THE EARTH.

1. Of what materials does the solid land consist? How are these materials arranged so as to form the vast mass of the globe? Man cannot pierce beyond the mere outer skin of the planet; can he then ever hope to learn the probable constitution of its interior? These are some of the questions to which we must now turn. In endeavouring to find answers for them let us begin with what is nearest and most familiar.

2. No matter in what part of the world we may live, the uppermost layer or covering of the land is almost always one of vegetation, in one place grass, in another forest, in another thickets of jungle or brushwood. Here and

there, indeed, where the surface is formed of shifting sand no plants may be able to take root, or in rough, and especially mountainous ground, ledges and crags of bare rock shoot up into the air. But even in such places as these we may often detect some struggling weeds or shrubs which are striving to bind the loose sand together, or to find a lodgement among the crannies of the rock.

3. Stripping off the layer of vegetation we see below it the layer of soil on which the plants grow. They send their roots down into this layer and extract from it some of the materials out of which their framework is built up. It varies greatly in colour and composition, being sometimes a stiff, grey clay, sometimes a soft, black, crumbling loam, sometimes a brown or yellow sand, and sometimes so stony as to be little else than a sheet of gravel. In all cases, however, it will be found to consist of small particles or larger fragments, which appear broken or worn. These have been evidently derived from some solid rocks. Look at them with a magnifying-glass and you will notice this broken and derivative character still more clearly. Roots of plants descending through the soil open it up and permit rain and air to pass into it, while the common earth-worm likewise aids in this work, and brings up lower portions of the soil to the surface.

4. The soil contains more or less organic matter, besides the grains of sand, earth, or clay. You drive this off by burning the soil, and in so doing you deprive the soil of its fertility. For the organic matter is derived from the mouldering remains of plants and animals, and is made use of by living plants. When a farmer has cultivated his fields for a long time, each successive crop has taken so much organic matter out of the ground that in the end the soil becomes so exhausted that it needs to have fresh organic matter put into it in the form of manure before it will continue to yield good crops. In some countries

where the soil is deep and fertile a long time may elapse before such exhaustion takes place.

5. The soil varies greatly in depth. A common thickness is three or four feet, but on sheets of rock it may not be so much as an inch, while in rich plains it may be several yards. The next layer below it is called the **subsoil**. It consists of the same materials as the soil, but less finely broken up and with little organic matter. Indeed the soil is simply the upper part of the subsoil

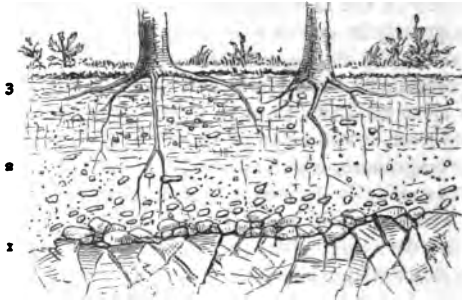


FIG. 25.—Section to show the formation of (3) soil and (2) subsoil from rotting away of (1) rock underneath.

which has crumbled away and been mixed up with the decayed remains of plants and animals. Usually only the longer roots of the larger plants, such as forest-trees, reach down into the subsoil. But as the soil is washed off by rain from the surface of the ground, the upper parts of the subsoil, coming in consequence nearer to the surface, are more exposed to air, rain, plant-roots, and earth-worms, so that they gradually pass into soil.

6. Beneath the subsoil lies the **rock**, clay, or other substance, from the breaking up and decay of which th

subsoil is formed. The character of the soil thus essentially depends upon the nature of the rock underneath. We shall trace in subsequent Lessons how universally and strikingly the surface of the land is crumbling down, even when it consists of the most solid rocks, such as granite, limestone, or sandstone. For the present we may bear in mind that the covering of soil so generally distributed over the land, and on which the verdure and fertility of every country depend, has been formed by the gradual decay of the rocks and of the remains of the successive generations of plants and animals which have lived upon the surface.

7. Beneath the decaying soil and subsoil, then, we arrive at the undecomposed rock. In most countries of the world a great variety of rocks occurs. To describe these and trace their origin and history forms the subject of the science of Geology. But a little consideration here of some of their more important characters will enable us to understand better how the materials of the land have been put together, and what is the probable constitution of the interior of the planet.

8. In the first place, then, the most cursory observation suffices to show that most of the rocks of the land have, like the soil, been formed out of worn fragments or particles of older rocks. Sandstone, for instance, out of which so much of the solid framework of the plains, hills, and mountains has been built, consists of mere sand compacted into solid stone. Shale and slate are only so much hardened clay or mud. Conglomerate or puddingstone is nothing but a mass of solidified gravel. In all these cases the materials of which the rocks are composed have been worn away from still older rocks, and have been rolled to and fro in water, as gravel, sand, and mud are produced at the present day.

9. Rocks formed of these water-worn materials have

been accumulated to depths of many thousand feet. They now frequently form lofty ridges of mountains. They spread also over the lower grounds ; most of the wide plains of the world have rocks of this kind below them.

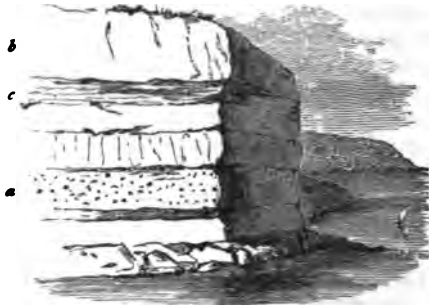


FIG. 26.—Bedded arrangement of rocks. (a) conglomerate ; (b) sandstone ; (c) shale.

One of their most obvious features is their arrangement in layers, beds, or strata, which vary in thickness from less than an inch to several feet or yards, and are piled regularly above each other. Hence a cliff of these rocks has a striped or banded look (Fig. 26). They are known as bedded or stratified rocks.

10. Together with these there often occur other rocks formed wholly or in great measure of the remains of plants or animals. For instance, amid a mass of sandstone and hardened clay you will occasionally come upon traces of the fronds of ferns (Fig. 27), and the seed-cones, leaves, stems, and roots of other plants. These vegetable remains sometimes have been so crowded together that they form black or brown seams of coal, and in that condition supply much of the fuel which man consumes at the present day.

11. Again, you will still more frequently meet with rocks composed mainly or entirely of broken shells, corals, and other animal remains. Many limestones, for example, have this composition. In the accompanying figure (Fig. 28) a piece of limestone is shown which is made up of the crowded joints of the encrinite or stone-lily—a marine animal of which there are some small modern kinds still living. Chalk in like manner is formed of the broken

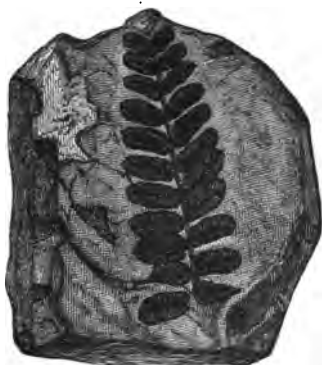


FIG. 27.—Piece of shale containing portion of a fossil fern.

remains of minute forms of marine animal life as represented in Fig. 29, which you may compare with Fig. 14, illustrating the modern ooze of the Atlantic sea-bottom.

12. Such rocks as limestone and chalk cover thousands of square miles in almost every country, and not only underlie parts of the plains, but, like the sandstones and shales, rise up even into lofty mountains. The Alps, Himalaya, Rocky Mountains, and Andes, for example, are actually built up in great measure of limestones and other rocks, formed mainly or in part of animal remains.

13. It appears, then, that most of the land, as far as we can see it, is made of rocks, which are either hardened



FIG. 28.—Piece of limestone, showing how the stone is made up of animal remains.

gravel, sand, and clay, or have been formed out of the broken and compressed remains of once living plants and

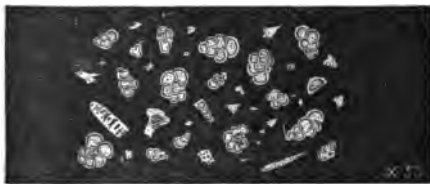


FIG. 29.—Some of the grains of a piece of chalk.

animals. Whence we conclude that what is now land

must have been under water, and, since the limestones and other strata contain chiefly marine forms of life, that this water must have been the sea. It follows, therefore, that the materials out of which the land has mainly been formed were laid down in the sea, and have been by some means raised up into land.

14. In the second place, besides these various stratified rocks, which have been chiefly derived from the wearing down of older rocks than themselves, others may be seen in which no bedded or stratified arrangement exists. These are not made up of fragments derived from



FIG. 30.—A piece of granite, showing the composition of a Crystalline Rock.

previously-formed rocks, but, for the most part, are crystalline, that is, are made up of crystals, either felted together or imbedded in a glass or base. Granite, porphyry, and basalt are examples of such crystalline rocks. Instead of spreading out in vast sheets over the continents as the stratified series does, the crystalline rocks occur in lines or points, and rise through the stratified rocks sometimes in huge disruptive masses, sometimes in veins and variously-shaped intrusions. They are particularly to be noticed along the central portions of mountain-chains, where they may be seen coming up through the very oldest of the stratified rocks. They likewise appear in the

eruptions of volcanoes—the molten rock known as lava being one form of the crystalline rocks.

15. There can be little doubt that the rocks of this second or crystalline series have come from below, and have been thrust in a melted state among the other rocks, or have been poured out at the surface as lava. Hence we must infer that beneath the outer layer of stratified or derivative rocks, though it may be many thousands of feet in thickness, there must be an inner layer or mass of crystalline material which has here and there been squeezed through the stratified rocks, as in the axis of mountain-chains, or has communicated with the surface by means of the openings of volcanoes.

16. So far our conclusions are founded on what can be actually seen. They may be regarded, therefore, in the light of established truths. We have traced the materials of the solid earth from the thin outer layer of surface-soil down through the thick piles of stratified sandstones, clays, limestones, and other rocks into the still lower granite, lava, and other crystalline masses. We have seen that by some means the materials of the land have been raised up out of the sea, and that portions of the underlying crystalline rocks have actually been pushed up into the very heart of the mountains. Can we descend any further and trace still deeper layers in the structure of our planet? Not directly; no distinct lower portions have come through the crystalline rocks. Nevertheless, enough evidence may be gathered to indicate with some probability what lies still further down.

17. Most of the rocks at the earth's surface weigh from twice to thrice as much as water. In other words, their specific gravity is from 2.0 to 3.0, that of pure water being reckoned as 1.0. Experiments made with the pendulum and the plumb-line on the earth's attraction, indicate that the weight of the planet is about twice as

much as that of the rocks at the surface, or, more shortly, the density of the earth is about 5.5. You must not at once conclude, however, that the interior of the earth consists of materials twice as heavy as those forming the land.

18. Owing to the increase of gravity towards the centre of the planet the density of every substance, if no counteracting force should come into play, would become greater in proportion as it approached the centre. Air, for instance, would be as heavy as water at a depth of thirty-four miles, and water would be as heavy as mercury at a depth of 362 miles. Hence, we might expect that the earth should be vastly heavier as a whole than even the heaviest of the ordinary rocks at its surface. That it is not so can only be satisfactorily accounted for by the existence of some force within the planet which opposes this increase of density. The only force known to us by which this can be done is heat. We should infer, therefore, even were there no positive evidence on the subject, that the internal parts of the earth must have an enormously high temperature whereby their materials are so expanded that the density of the whole planet is kept from being so great as the progressive increase of density from the surface downwards would lead us to expect.

19. The deepest mines do not reach one mile in depth, or $\frac{1}{362}$ of the distance from the surface to the centre of the earth. Man can never indeed hope to penetrate far into the interior of the planet on which he lives. There are yet several kinds of evidence open to him which go to demonstrate this high internal heat. These are : 1st, mines, wells, and borings ; 2nd, hot-springs ; and 3rd, volcanoes.

20. (i.) It has long been known that the air of deep mines is warmer than that above ground, and that, as a rule, the deeper the mine the warmer is the air

in it. A coal-mine sunk near Manchester, for example, to the great depth of 2,151 feet has a constant temperature at the bottom of 75° Fahr. while the average temperature near the surface is only 51° Fahr. The water which rises from deep borings is warm. Thus from the well which was sunk to a depth of 1,798 feet at Grenelle, near Paris, water issues with a temperature of 81·7° Fahr. As the results of observations made all over the globe it appears that an increase of temperature always occurs, and that though it varies in amount in different kinds of rock, its average rate is about 1° Fahr. for every sixty feet of descent.

21. If this rate should continue it is plain that at a comparatively short depth even the most refractory substance would be at its melting-point. Water at 12,000 feet would be as hot as boiling water at the surface. At a depth of about twenty-four miles it would be as hot as melting gold.

22. (ii.) In almost all countries of the world springs of warm water gush out from the ground. In volcanic districts the water is often even at the boiling-point, and remains so permanently. Nay, in some cases, as in one of the hot springs of Iceland, the water is actually raised to a temperature of 261° Fahr., or 49 degrees above the common boiling-point, so that when it approaches the surface of the ground it flashes into steam with a loud roar, and throws out jets of boiling water and steam.

23. Warm springs often occur at a distance from any active volcano. Those of Bath, for example, in the lowlands of the south-west of England are more than 1,000 miles from the burning mountains of Iceland on the one hand and more than 1,100 from Vesuvius on the other. At Buxton in Derbyshire the springs have a temperature of 82°. In Germany the wells of Wiesbaden throw out water at a temperature of 158°, and those of Carlsbad at

167°, while in the north-west of Spain some of the springs have a temperature as high as even 192°. All these places are remote from an active volcano, but it may often be noticed that hot springs rise along mountain-chains or at least on lines where the rocks have been intensely crumpled and where they may have been greatly heated during the crumpling. Here and there, as will be pointed out in the next Lesson, the remains of volcanoes occur which are now quite cold and silent, and from which no eruption has taken place within the memory of man. Yet some of these extinct volcanoes are still associated with hot springs. In the ancient volcanic district of Central France, for example, numerous springs occur, sometimes with a temperature of as much as 174°.

24. By far the most remarkable kinds of hot springs, as showing the great heat of at least some parts of the earth's interior, are those which throw out their contents with violence. They occur in volcanic districts and are called geysers, that being the local name borne by those of Iceland, which were first described. The part of Iceland where they chiefly occur is a space about two miles square in a low, wide valley among volcanic rocks. In this limited tract the ground is pierced with abundant openings from which jets of steam and hot water escape. These openings, said to be nearly 100 in number, are usually surrounded with basin-shaped rims or mounds of white, grey, and variously coloured incrustations of silica, deposited from solution in the hot water. They vary in size from mere tiny bowls, a few inches across, up to huge cauldrons, like the Great Geyser, which rises fifteen feet above the ground and measures fifty-six feet in diameter. In the centre of the basin of the Great Geyser a pipe or funnel, eight feet wide, descends into the earth. From this opening boiling water is constantly ascending into the basin and flowing over the lowest part of its lip into the plain. At intervals

of a few hours a loud rumbling is heard to proceed from the funnel, the water in the basin begins to boil up, and jets of it together with clouds of steam are spurted up for a few feet. After these intermittent eruptions have continued for about a day they are succeeded by a much

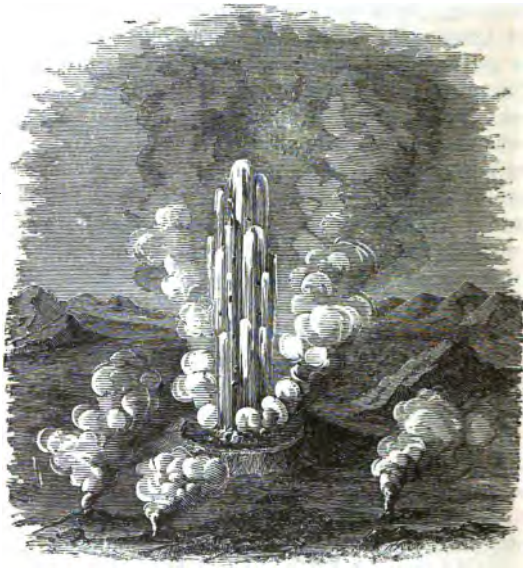


FIG. 31.—View of the geysers of Iceland. Great Geyser in eruption.

more violent explosion. The ground trembles slightly as the rumbling sound increases in violence until a huge column of boiling water 150 or 200 feet high is thrown up into the air, with loud explosions and clouds of steam. In this way the basin and pipe are emptied of water, but they

gradually fill again, the rumblings and jets of water and steam begin anew, until next day another grand outburst empties the geyser.

25. The Yellowstone National Park, in the Territories of the United States, contains several hundreds of geysers scattered over a wide tract of country. Many of these are extinct, their rims and basins of siliceous incrustations standing out above the ground to mark where the boiling water once rose. Some exceed the Great Geyser of Iceland in size and in the height and volume of water and steam which they discharge when in eruption.

26. Another interesting series of geysers occurs in New Zealand, where also active and extinct volcanoes are met with. The spouting hot springs of Orakeikorako gush forth in numbers on either side of a river-valley, while at Tetarata the water is so largely charged with silica that it has formed a series of terraces and basins down the side of the precipitous bank from near the top of which it takes its rise.

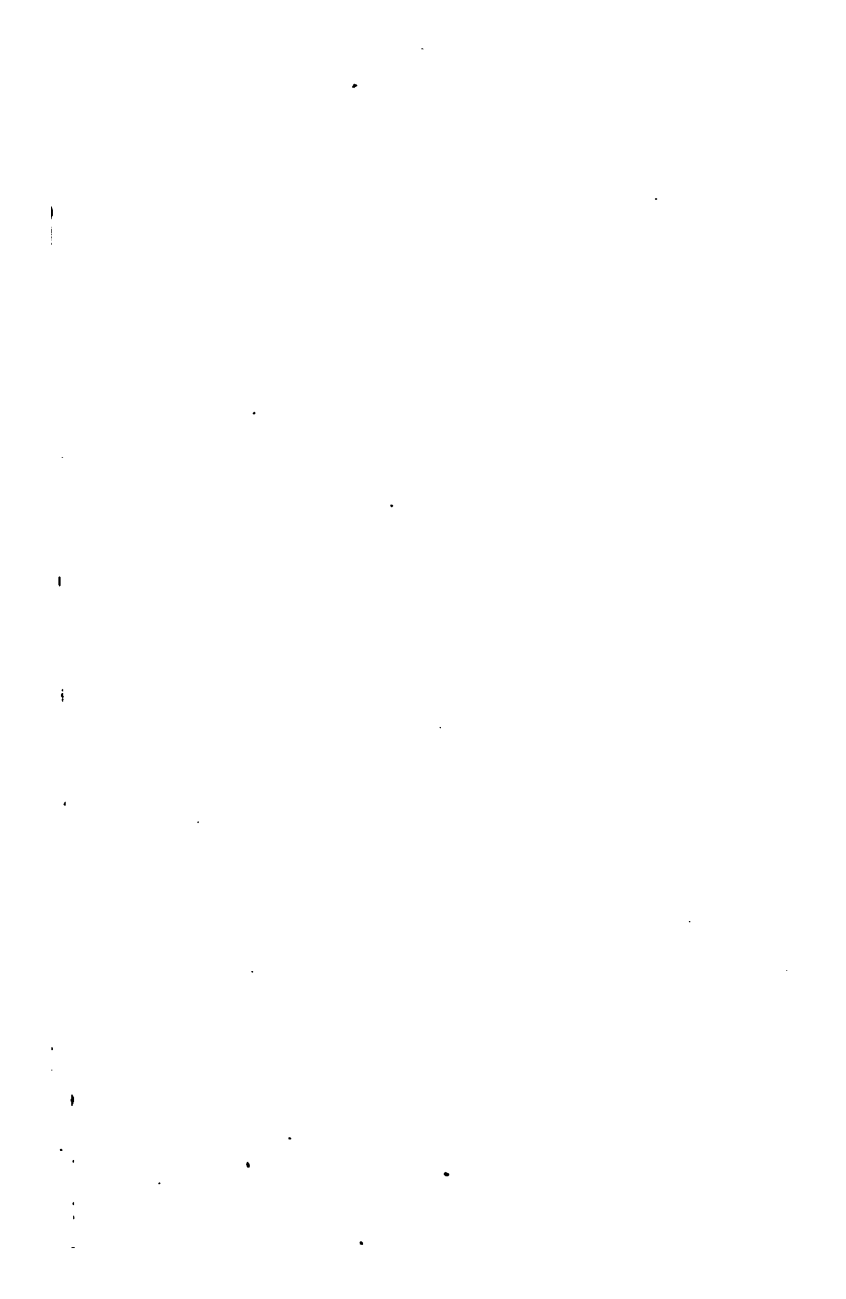
27. (iii.) **Volcanoes** form so interesting a part of the earth that they will be described in some detail in the following Lesson. Let us note here that they are openings from which steam and other hot vapours as well as streams of melted rock are discharged, that they occur in many parts of the world, and that they prove that within the earth masses of rock exist at a white heat.

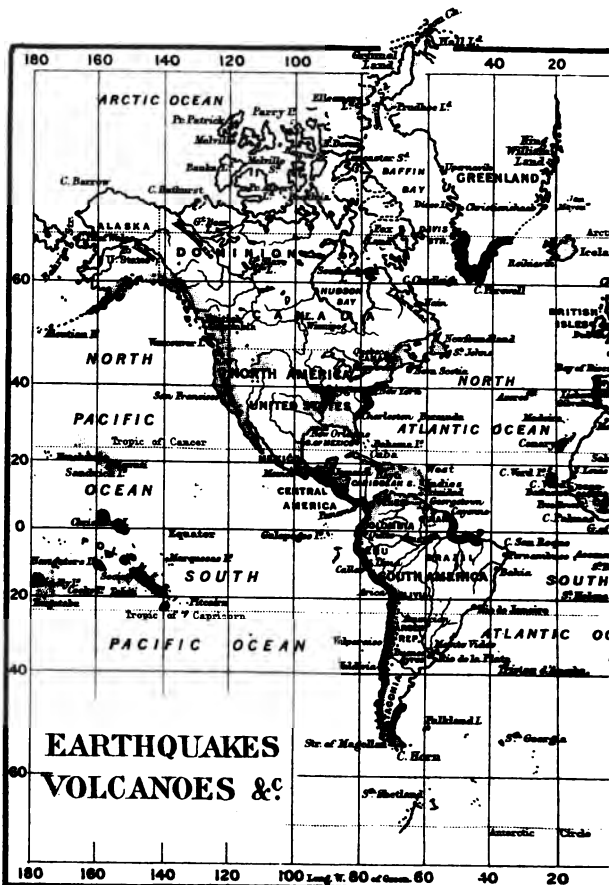
28. It is plain, then, from such evidence that the earth cannot but have an enormously high internal temperature. Much dispute has arisen as to whether the interior is liquid or solid. Some have maintained that the earth consists of a ball of molten material with an exterior crust which has been variously estimated at from twenty to a thousand miles in thickness. Others have insisted that a planet so constituted could not rotate as the earth does, and that our globe must be solid to the centre.

29. If the heat goes on increasing at the same rate as it is observed to do in the comparatively shallow borings which man can make into the earth, the ordinary melting-point of even the most refractory substance would, as we have seen (Art. 21), be met with at a comparatively short depth beneath the surface. But it does not follow, therefore, that the materials of the interior should be actually in a state of fusion. We know that pressure has the effect of raising the melting-point of such substances as those which form most rocks, that is, it keeps them from melting until a still greater heat is applied. The pressure at great depths within the earth must be enormous. Hence below a depth of a few miles, where the temperature reaches the ordinary melting-point of most rocks at the surface, each successive layer of the earth's substance may not be actually fused but may be just at the melting-point proper to its depth. The whole globe might thus be solid, but the least diminution of pressure at any point would allow the parts so relieved to melt at once. This, so far as can be judged, seems to be the most probable condition of the interior of the earth.

30. The internal heat being so great, there seems no reason why the innermost parts of the planet may not consist of metal, such as iron or gold. And indeed there is some reason to infer that they are really metallic. Cracks have been made abundantly through the rocks forming the land, and in many of these metallic ores occur which are believed to have not improbably come up from a metallic region below. Researches into the constitution of the sun and of the other planets have tended to confirm this view of the metallic composition of the central parts of the earth.

31. Summing up what has been said in this and previous Lessons regarding the constitution of the globe, we find that beneath the outer envelopes of the atmosphere and the



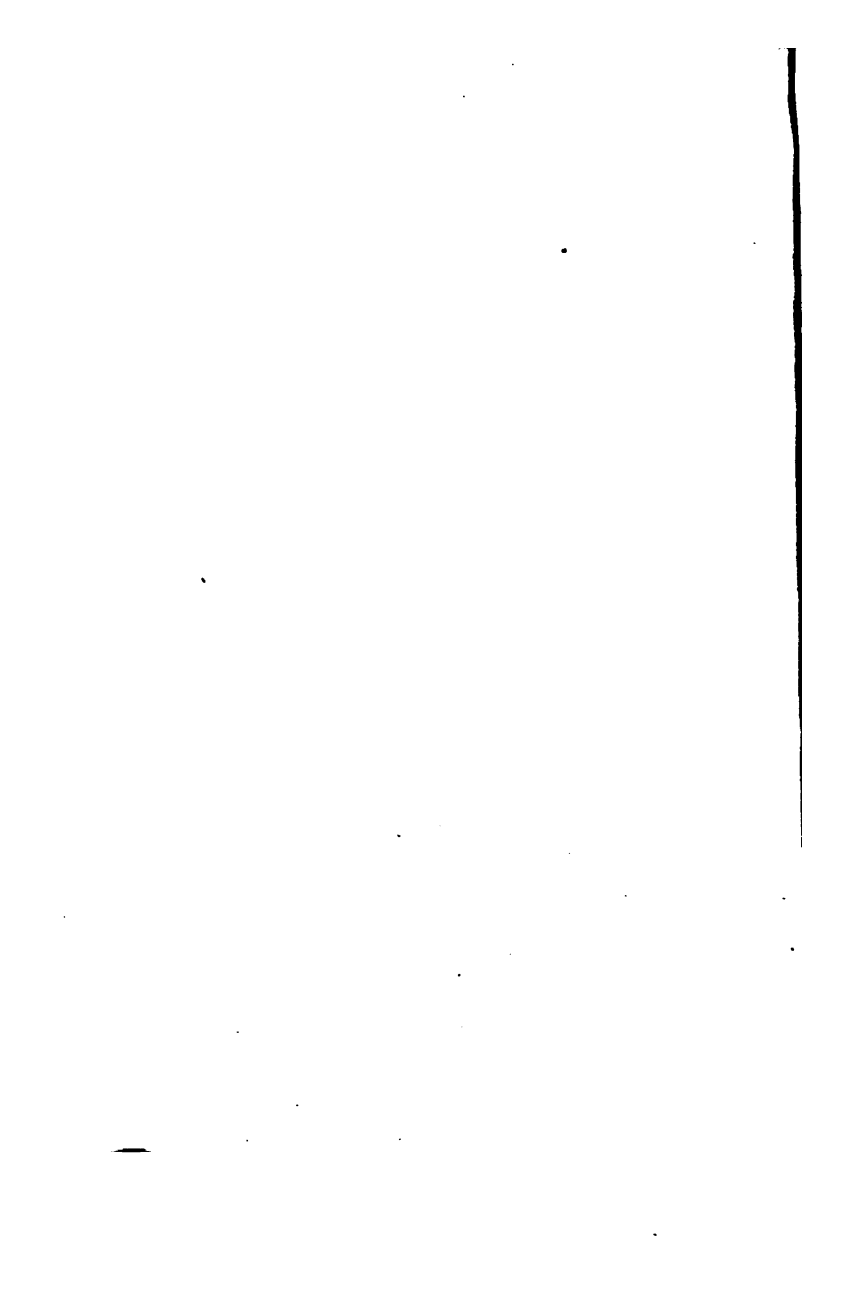


W. & A. K. Johnston

- Regions subject to earthquakes; The intensity
- Active, dormant, or recently extinct Volcanoes.
- Regions now rising or which have been upheaved
- Regions now sinking.



Frequency of the shock being marked by the depth of shading.
in comparatively recent times.



sea, the solid earth has an upper layer of loose, crumbled materials—gravel, sand, and mud on the sea-floor, soil upon the land—under which lies a thick series of stratified and derivative rocks ; that still further down masses of crystalline rocks, which have in many places been forced through the overlying series, descend to an unknown depth, and that within them there may be a metallic central mass. We see, moreover, that since the heat rapidly increases as we descend, the innermost parts of the globe must have a temperature far more than sufficient to melt every known substance and even perhaps reduce it to the state of vapour, but that the pressure of gravity may be great enough to retain the general mass of the globe in a solid state, except at such places as supply the streams of molten lava which flow from volcanoes.

LESSON XXII.—VOLCANOES.

1. Reference has been made in previous Lessons to volcanoes and some of their characters. We are now to study them more in detail. The word volcano is derived from Vulcanus, the name of the Roman god of fire, who was supposed to have his subterranean forges at the roots of the mountain Etna. It is applied to any conical mound, hill, or mountain, formed of materials which have been erupted from beneath the surface. When it is active it emits gases, steam, water, mud, dust, stones, or molten rock from its summit, or from fissures or other openings on its sides. The term volcanic action is used to describe all the kind of work done by a volcano. A volcano may be dormant when it remains a long time without giving any of the usual signs of volcanic action ; and it is said to be extinct when, though its external form may be

retained, it has never been known to be active and does not seem to have been so for a great many centuries.

2. The size of a volcano varies from a mere little mound a few yards in diameter, like some of the mud-volcanoes around the Caspian Sea, up to a giant mountain like Cotopaxi, which rises among the Andes to a height of 18,887 feet above the sea, its upper 4,000 feet forming a smooth snow-covered cone with an orifice at the top whence hot ashes and stones are scattered far and wide over the surrounding country.

3. At the top of a volcano lies a basin-shaped hollow called the crater, from the bottom of which the pipe or shaft descends, whereby the volcanic products are brought to the surface. Vast showers of fine dust and stones are frequently thrown out from most volcanic craters. These materials falling down the slopes of the cone gradually increase its diameter and height. In like manner streams of molten rock called lava, issuing either from the lowest part of the lip of the crater or from some fissure or orifice on the side of the hill, pour down the slope and harden there, thus still further augmenting the bulk of the volcano.

4. As a volcano increases in size, and cracks are formed in weak parts of the cone, smaller cones are piled up on its flanks by the emission of dust, stones, and lava from these fissures. A large volcanic mountain, like Etna, or the Peak of Teneriffe, is thus sometimes loaded with small volcanoes which often reach a height of five or six hundred feet. In the drawing (Fig. 32) a plan is shown of one of these great volcanoes with its groups of minor cones.

5. At the beginning of a volcanic eruption rumblings are heard like the muttering of distant thunder, while the ground is felt to tremble slightly. These noises and tremors increase in intensity, successive loud explosions take place in the pipe of the volcano, and at last clouds

of fine dust and steam are hurled with prodigious force far up into the air. The steam rapidly condenses into rain, which falls in torrents down the outer slopes of the mountain. The fine dust is sometimes given out in such quantities as to darken the sky for many miles around. In the famous eruption of Vesuvius, which in the year 79 destroyed the Roman cities, Herculaneum, Pompeii, and

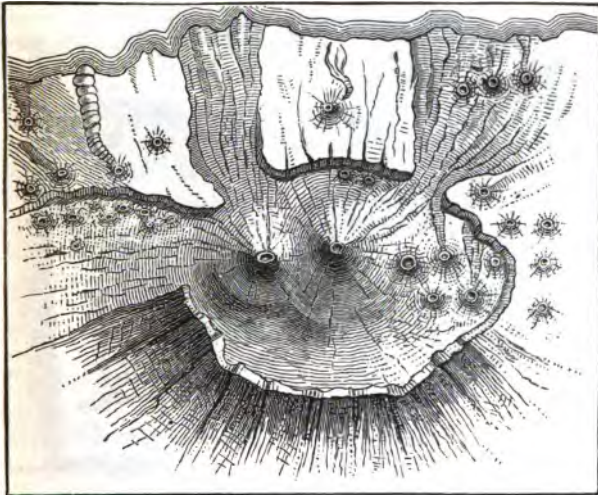


FIG. 32.—Plan of the Peak of Teneriffe, showing the large crater, partly effaced, and smaller craters with lava currents issuing from them

Stabiæ, the air was as dark as midnight for twelve or fifteen miles round, and a thick deposit of fine ashes and stones fell on this whole district. Some notion of how completely this eruption of ashes covered the country may be obtained from the accompanying drawing, which

represents one of the streets of Pompeii as it now appears, after the volcanic deposit under which the city lay buried for sixteen centuries has been cleared off. At the further end of the street where the excavations have stopped, the dark layer on the top of the walls, and on which the pine-tree is growing, represents the thickness of the general



FIG 33.—View of a Street in Pompeii.

covering of volcanic ashes. The finer dust, carried up into a strong upper current of air, is sometimes transported for hundreds of miles before sinking to the ground. Illustrations of this kind of transport were cited in Lesson XI. to prove the existence of these upper currents in the atmosphere.

6. Besides the light dust vast numbers of white-hot blocks and smaller fragments of stone are ejected from the crater. Many of these strike against each other as they rise and fall, producing at night a wonderful spectacle as their sparks and flashes light up the darkness. Of the force with which the stones are occasionally discharged some notion may be formed from the feats of the volcano Cotopaxi, which is said to have thrown a block, estimated to weigh 200 tons, to a distance of nine miles. The volcano of Antuco in Chili has sent stones flying to a distance of thirty-six miles.

7. It is not difficult to understand why in the earlier stages of a violent eruption so vast an amount of fragmentary materials should be ejected. During the successive explosions the walls of the crater and the accumulated lava and other volcanic ejections which had gradually choked up the funnel are rent and finally blown into fragments. The water which, under enormous pressure, has been retained in the liquid form at a temperature far above the boiling-point, and in that condition has, perhaps, saturated the molten and solid rocks in the pipe of the volcano, instantly flashes into steam when in its upward ascent it can overcome the influence of the superincumbent pressure. So violent is the conversion of such red-hot or even white-hot water into steam that the molten lava is thrown out in huge jets and actually blown into the finest dust. It is the liberation of successive portions of this highly-heated water into the gaseous state that produces the explosions which form so magnificent a part in a great volcanic eruption. At each explosion a vast ball of steam shoots up into the air, at once condensing into white clouds and rolling over in huge folds (Fig. 34), which either dissolve and float away in the higher atmosphere or are further condensed and fall as rain.

8. We can perceive, too, why it is that the upper part of a cone during the early part of a tremendous eruption should sometimes disappear. It is shattered by the explosions and blown up into fragments, which fall back either into the crater or down the outer slopes of the mountain. Mount Vesuvius has supplied some excellent

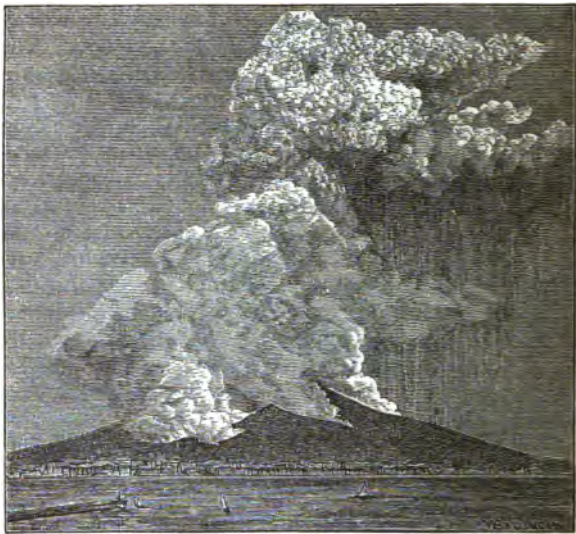


FIG. 34.—View of Vesuvius as seen from Naples during the eruption of 1872.

illustrations of this kind of destruction. Prior to the first century of the Christian era that mountain was a dormant volcano, from which no explosions had ever been known to come, but which had an enormous crater on its summit, overgrown then with brushwood and wild vines, like the

crater of Astroni and some of the other extinct volcanoes near Naples (see Fig. 5). Suddenly in the year 79, when the great eruption took place which destroyed Pompeii, the south-western side of the cone was blown away and a new cone of much smaller dimensions was formed inside the circuit of the former crater. In the drawing of Vesuvius (Fig. 35) as seen from the sea, the semicircle left of the old crater appears behind the modern diminished cone.

9. It is no doubt by the enormous expansive force of

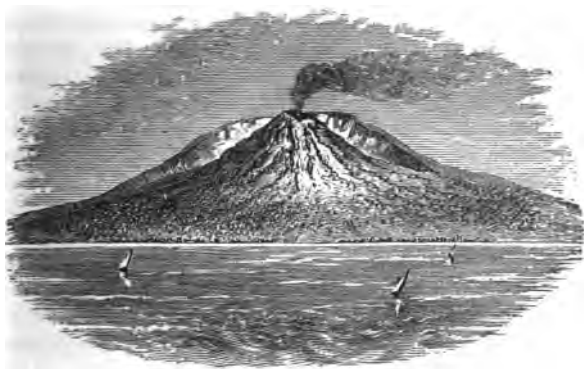


FIG. 35.—Mount Vesuvius as seen from the sea, with the remaining part of the old crater of Somma behind.

the imprisoned water and steam, that the molten lava is forced up the pipe of the volcano. After the first explosions, lava is seen to flow either from the top or from one or more points on the side of the cone. Should the sides of the mountain be solid enough to resist the enormous pressure of the ascending column of lava the latter will, of course, find no escape until it fills up the crater to

the level of the lowest part of its rim, over which it will pour down the mountain. More usually, however, there are weak parts, such as rents, caused by the previous explosions, through which the lava finds egress to the outer slopes of the cone.

10. Few sights in nature are more terrible than that of a lava torrent as it issues and pours down the side of a mountain. At first it glows with a white light, and flows freely like melted iron. It rapidly grows red and darkens; at the same time its surface hardens into a black or brown crust which breaks up into rough cinder-like pieces, beneath which, as you may see at the rents, the main mass still remains red-hot. A short way from the point of emission the lava-stream looks like a river of rugged blocks of slag grinding against and over each other with a harsh metallic sound, and revealing every here and there a glimpse of the fiery flood underneath over which they are floating. Clouds of steam and hot vapours rise from all parts of the moving mass. Its rate of march varies much according to the slope of the ground, distance from the point of exit, and other causes. Thus, in the year 1805, a current of lava ran down the first three miles of the slope of Vesuvius in four minutes, yet took three hours to reach its furthest point, which was only six miles. In 1840 one of the more liquid lavas of Mauna Loa, in the Sandwich Islands, was observed to run eighteen miles in two hours.

11. Standing in front of a slowly advancing lava-current, you see what seems a huge mound of rubbish, somewhat resembling one of the piles of debris at iron-works, or a new railway embankment which still shows its raw earth and stones. You observe, however, that the blocks are tumbling over each other and that the mound is slowly advancing. A breeze coming from the lava is felt to be hot and stifling, for it carries with it some of

the steam and acid vapours which are coming off abundantly from the molten rock. Trees, walls, gardens, and vineyards are successively buried under the burning flood. A house or range of houses may form a temporary barrier ; but eventually these buildings are pushed over like mere houses of card and engulfed (Fig. 36), or the lava is piled

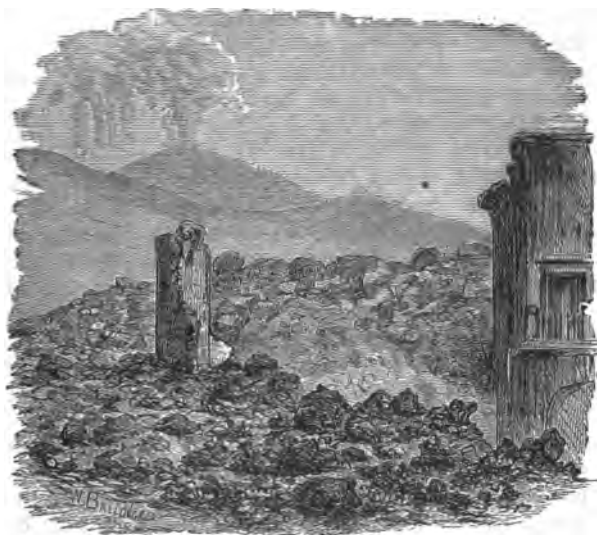


FIG. 36.—Houses surrounded and partly demolished by the lava of Vesuvius, 1872.

up against and passes round them, meeting again below so as to leave their highest points, perhaps, projecting from the middle of a rugged sea of lava. It sometimes happens that the still liquid rock inside bursts through the hardened crust at the side or in front and pours down in a new

direction. This is always one of the risks to be considered in such populous and cultivated tracts as those along the slopes of Vesuvius.

12. The distance to which a torrent of lava may flow can never be foretold, even when it is seen pouring down from a mountain at a great velocity. At Vesuvius many of the lava-streams have reached the sea ; others have not travelled further than a few hundred yards from their point of emission. The most tremendous floods of lava ever known to proceed from a volcano were those of Skaptar Jokul in Iceland, in the years 1783-5. Vast torrents of molten rock were poured over the country, filling up river-courses, ravines, and lakes, and completely destroying the surface for many hundreds of square miles. The lava reached to an extreme distance of forty-five miles in one direction and of fifty miles in another. In some level places, where the lava spread out, the stream reached a breadth of fifteen miles and a thickness of 100 feet, but it accumulated in narrow valleys sometimes to a depth of 600 feet. It has been computed that the total mass of lava poured forth during that series of eruptions would form a mountain equal in bulk to Mont Blanc.

13. The hardened crust on a lava-stream forms an excellent non-conducting layer between the still molten mass underneath and the air. You can walk on that crust, though through the fissures the red, glowing lava may be seen lying only a few inches below. It is evident that the innerparts of such a current, even at a considerable distance from the point of emission may retain a very high temperature long after the current itself has come to rest. Thus the lava of the eruption of Vesuvius in 1858 continued to give out steam and hot vapours in 1873, and still retained so much heat that one could not keep one's hand in some of the fissures for more than a few seconds,

although the surface of the current was everywhere quite cold.

14. When the lava begins to flow freely from a volcano the violence of the eruption usually abates. The showers of ashes gradually cease, or at least do not extend beyond the crater and its immediate vicinity. The explosions and subterranean rumblings disappear, and except for the cloud which is always hanging over the summit of the mountain and which marks how constantly and abundantly steam still continues to rise from the crater, nothing at last remains to indicate from a distance that volcanic action is not extinguished, but merely quiescent for a time. Yet a visit to the summit of the cone at such an interval would show that hot vapours and gases still keep streaming out from the summit and sides of the mountain. Occasionally after an eruption of Vesuvius it is observed that a large destruction takes place among hares and birds on the hill-sides, owing to the plentiful emanation of the poisonous carbonic acid gas. Many centuries after a volcanic district has ceased to be subject to any eruptions this gas continues to rise, sometimes in large quantities, either bubbling through the water of springs or coming out of crevices in the ground.

15. The number of active volcanoes on the surface of the globe is computed to be about 173. These are not distributed at random, but follow certain lines, as shown in the map on Plate IX. It will be observed that these lines generally coincide with some of those ridges on the surface of the earth which have been already referred to in Lesson XIX. Almost all volcanoes are thus ranged close to one or other of the great ocean basins. The continental and island barriers which encircle the Pacific Ocean form one vast ring of active volcanoes. Beginning on the east side we find in the giant chain of the Andes a series of active volcanoes, some of them the loftiest on the

globe, running along the western margin of South America. It is continued through Guatemala and Mexico into North America, and stretches through the Aleutian and Kurile Islands, forming a close-set fringe to the Pacific Ocean on the north. The line is prolonged down Japan, the Formosa and Phillipine Islands to the Malay Archipelago where it divides into two branches. One of these turns south-eastward by New Guinea and the New Hebrides into New Zealand. The chain is prolonged, but with wide gaps, across the Pacific basin in the volcanoes of the Friendly, Society, Marquesas, and Easter Islands, towards the coast-line of South America, thus completing the vast volcanic ring. Even along the centre of that basin on the submarine ridge alluded to in Lesson XIX., Art. 11, magnificent active volcanoes appear in the Sandwich Islands, some of those in Hawaii exceeding 13,000 feet in height.

16. Returning to the Malay Archipelago we observe that the second branch of the volcanic line turns north-westward through Java and Sumatra, where more active and dormant volcanoes are crowded into a shorter space than anywhere else in the world. It is prolonged through the islands off the west coast of Burmah. After a wide interval we catch it again in the west of Asia, on the southern borders of the Caspian Sea, whence it may be traced at the Greek Islands, Vesuvius, and the Italian Islands, to the Azores, Canaries, and Cape Verde Islands.

17. Besides these main lines, however, scattered volcanoes occur on other distant islands or on the edges of the continents. In the North Atlantic and Arctic Oceans lie the active craters of Iceland and Jan Mayen. On the west side of the basin of the Indian Ocean we find small volcanoes on the Red Sea and in the Isle of Bourbon. The line of volcanoes in Terra del Fuego at the southern apex of America seems to be continued in the chain of

the South Shetland Islands, and to extend even into the supposed Antarctic continent, where, amid the vast snow-fields of that region, Sir James Ross in 1841 discovered an active volcanic cone 12,369 feet high.

18. But besides these active volcanoes there is a still larger number which are either dormant or extinct. It would seem that few large tracts of land exist where evidence may not be obtained of former volcanic action. Lava-streams and consolidated beds of volcanic dust may be found in almost all countries. Sometimes, indeed, as in Central France (Fig. 37) the cones are still as fresh as if they had been thrown up only recently ; and yet no record remains that they have ever been in eruption within the times of human history. Hence if to the present long list of active volcanoes we add those which are now extinct, we shall find the whole surface of the land to be studded over with points of volcanic eruption.

19. When we consider that each of these points marks an orifice at which highly-heated materials have been poured forth from the interior, and that they are so widely distributed over the earth's surface, we see how important is their evidence as to the high internal temperature of the earth (Lesson XXI., Art. 27).

20. It was pointed out in the last Lesson (Art. 29), that while the interior of the earth is probably solid as a whole, each portion of it, beyond a depth of a few miles, is probably at the melting-point and ready to pass into a liquid condition when any diminution of the pressure takes place. The ridges on the surface of the earth, formed, as they have been, during the contraction and consequent general subsidence of the outer parts of the planet, have doubtless by their uprise relieved the pressure upon the parts underneath them. This relief has probably allowed portions of the interior to pass into the state of fusion. You observe how the active volcanoes of the globe are mostly arranged

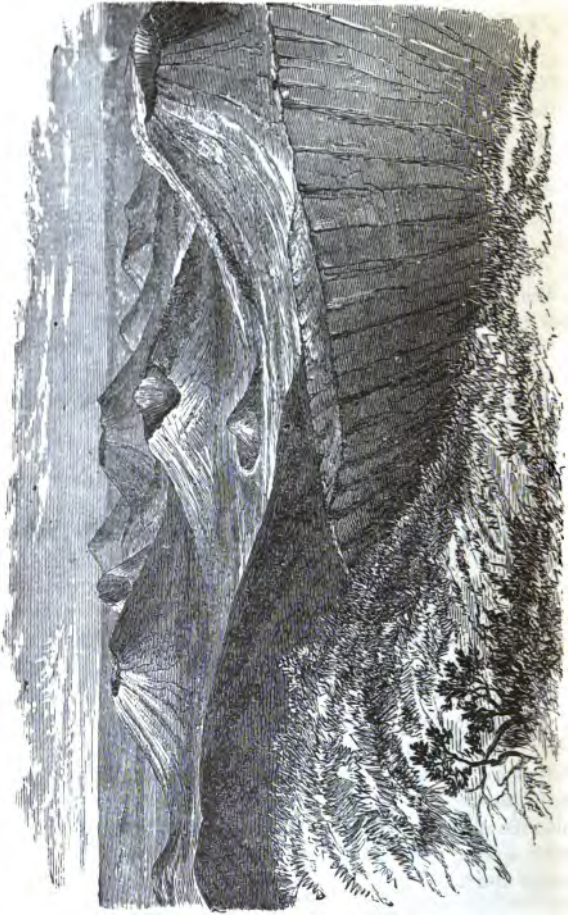


FIG. 37.—View of the extinct volcanoes of Central France, taken from the Puy de Pariou.

along such lines of elevation, whether on a continent as in South America, or in chains of islands as in the western and northern sides of the North Pacific Ocean. This arrangement can hardly be accident. It helps to connect the elevation of the land and the phenomena of volcanoes by showing how we should expect large spaces of melted rock to lie under those very regions where active volcanoes occur. Volcanoes do not pierce every mountain-chain, however, though in some cases they can be shown to have once existed but to have been long extinct.

21. The reservoirs of molten rock may exist underneath without giving rise to actual volcanic explosions so long as no passage is opened to the surface. When water derived from rain, rivers, lakes, or the sea, filters through the upper rocks and eventually reaches these deep, intensely hot regions it is raised to a very high temperature, indeed to the white heat of the molten rock with which it comes in contact. When portions of this superheated water succeed in effecting their escape, their removal relieves the pressure on the mingled mass of molten rock and water below, which is then forced upwards, and, where possible, finds an exit to the surface. It is these subterranean movements which give rise to the explosions, ashes, stones, and lava-streams of an active volcano.

LESSON XXIII.—MOVEMENTS OF THE LAND.

1. In the preceding Lesson some account was given of the way in which volcanic action affects the surface of the globe. Along certain lines or at certain points where communication has been opened between the surface-waters and the intensely hot materials below, showers of dust and stones and streams of melted rock have been emitted in such quantity as to form huge mountains like Etna,

the Peak of Teneriffe, and Cotopaxi. But other notable effects show the influence of the internal condition of the globe upon its surface. The solid earth is subject to movements either sudden and violent, or slow and imperceptible. It is sometimes convulsed and rent open, sometimes one tract is gradually raised up above the sea-level, while another is step by step depressed. We shall consider these movements under the three divisions of—1st, Earthquakes; 2nd, Upheaval; 3rd, Subsidence.

2. (i.) **Earthquakes.**—These, as the word denotes, are tremblings or concussions of the ground. They vary in intensity from mere slight tremors, like that caused by a loaded waggon moving along a street, up to such violent catastrophes as those in which the ground is thrown into undulations whereby rocks are loosened, trees shaken out of the soil, and villages and cities levelled to the ground.

3. It is chiefly the terrible devastation of life and property which is usually dwelt upon as the main feature of importance in earthquakes. In a moment, without the least warning of any kind, a rumbling like that of distant thunder or the firing of cannon is heard or felt, and before men have had time to ask each other what the tremor can be, the earthquake comes upon them. The ground rises and falls under their feet; houses rock to and fro until they are rent from top to bottom, or fall with a crash into ruins; the earth here and there opens and closes again. In a few seconds a whole city is demolished and hundreds or thousands of its inhabitants are dead or dying. The Lisbon earthquake in 1755, for instance, destroyed 60,000 human beings. The Calabrian earthquake a generation later proved fatal to not less than 40,000. Nor is the destruction confined to one limited district. Frequently it spreads over thousands of square miles, carrying death and havoc far and wide through different provinces and states.

4. When, however, we regard earthquakes as part of the machinery of nature, it is not so much their influence upon human life and property as their permanent effects upon the surface of the land which claim our notice. An earthquake is not felt simultaneously over the whole region affected by it. It begins at one side or end, and travels rapidly to the other, or it is experienced first and most violently over a certain central space, from which it spreads in all directions, getting feebler as it diverges. The sound of its approach, like that of distant thunder, artillery, or rumbling waggons, may precede by a few seconds the advent of the earthquake itself. When the shock comes the ground is felt to be alternately raised and depressed, with more or less violence, as if an undulation like the ground-swell of the sea were passing beneath. Several shocks in succession, but commonly less destructive than the first, may follow within a few seconds.

5. The earthquake is really of the nature of a wave passing through the substance of the solid earth. Just as the masts of ships at rest in a harbour are seen to rock to and fro as the swell of the sea rolls under them, so trees and other tall objects have been observed to sway backwards and forwards during an earthquake.

6. On steep banks and cliffs, large masses of clay or rock are sometimes disengaged and sent rolling down into the water-course or valley below. The streams are thus choked up, temporary lakes are formed, until, by the bursting of the barrier of fallen debris, the water rushes out with great force and carries another kind of devastation down the valley.

7. On many occasions the ground has been seen to be rent open. The fissures thus formed sometimes swallow up trees, houses, or other objects on the surface and close up again. Sometimes they remain open and are subse-

quently deepened and widened by running water into ravines and valleys.

8. Again, the ground has been found on some occasions to have been permanently raised above or depressed below its former level. Thus in the year 1835, during a violent earthquake, which convulsed the coast of Chili in South America, the ground at the island of Santa Maria was suddenly upheaved from eight to ten feet above high-water mark, so that the gaping sea-shells still adhering to the rock were exposed to the air, where they soon began to putrefy. The valley of the Mississippi was the scene of a succession of earthquakes from the end of the year 1811 to 1813, at the close of which time the ground was in many places left with huge yawning fissures, and in certain districts sank down so as to be converted into lakes, some of which are fifty miles in circumference. Part of this submerged land is called "the Sunk Country." And even now, large trees of walnut, white oak, mulberry, and cypress may be seen ten or twenty feet or more below the water, and a canoe may be paddled among their submerged branches.

9. When the earthquake shock takes its rise under the sea and travels thence towards the land, the greatest amount of destruction takes place. Not only does the solid land rock to and fro, but the waters of the sea are shaken and sent with terrible force against the margin of the land. From the point of origin of the shock a long, low undulation is propagated over the surface of the sea in all directions. When it reaches shallow water, its front, like that of the bore of a tidal wave in an estuary (Lesson XVII. Art. 30), becomes steep and advances with great rapidity. To those on shore the approach of this wave is seen to be preceded by a retreat of the sea from the beach. Spaces never bare before are now exposed to view as the water retires, but only for a few seconds. The wave, gathering up the retreating water, rushes

forward with a front sometimes sixty feet high, not only covering the beach but even sweeping far in upon the land.

10. This sea-wave does not travel so fast as the earthquake-shock. Hence it arrives at a coast some time later and completes the destruction. This was the case in the famous earthquake of 1755, by which Lisbon was destroyed, and again in 1868, when Peru and Ecuador were devastated by a disastrous earthquake. On the 13th of August of the latter year a great sea-wave inundated and overthrew Arica, the principal port in the south of Peru. In a few minutes every vessel in the harbour was either ashore, wrecked, or floating bottom upwards. A man-of-war was swept inland for a quarter of a mile. Another vessel disappeared, and no vestige of it was ever seen again.

11. The true cause of earthquakes is not yet well understood. Probably they take their rise from more causes than one. Thus they may sometimes be due to the giving way of the roofs of the cavities which no doubt exist in the interior of the earth, especially in volcanic countries, or to the sudden fracture of rocks under great strain, or to the sudden generation or escape of steam. Whatever be the cause of their origin, some sudden blow within the interior of the earth will account for the phenomena of earthquakes. The point directly above that from which the blow is given suffers the severest shock, and the intensity diminishes as the waves travel outwards from this centre, like the rings propagated from the point on the surface of a piece of still water which is struck by a stone.

12. If, after an earthquake has passed, the point can be ascertained where the shock was vertical, and if, by further observations upon the direction of the rents in walls and other evidence (Fig. 38), it can be determined

at what angle the wave emerged at the surface at one or more places beyond the centre, an indication may be obtained as to the probable depth at which the shock took its rise. Let *v* in Fig. 39 be the point at which the shock was felt to have come up vertically, and *A* a village or town where from the prevalent direction of the rents in the buildings the shock appears to have come up



FIG. 38.—House rent by earthquake (Mallet). The arrow shows the direction from which the earthquake-wave must have reached the surface.

obliquely. If we prolong the path indicated by these rents until it meets the vertical, we obtain at *F* the probable focus of disturbance. Calculations of this kind have been made by Mr. Mallet, who concludes that, on the whole, earthquakes are not deep-seated, but probably never arise at a greater depth than thirty geographical miles. As implied in the diagram, the successive waves are closer and stronger at the focus, whence they spread

in all directions through the solid earth, becoming less and less violent as they recede.

13. Earthquakes, as represented in Plate IX., are most frequent in volcanic districts, though not by any means confined to them. The great earthquake region of the Old World stretches from the Azores along the basin of the Mediterranean into the heart of Asia. In the New World the western border of the Continent suffers most from earthquakes, especially from Guatemala southwards through Ecuador, Peru, and Chili.

14. The area convulsed by a single earthquake is often great. It has been computed in the case of the Lisbon earthquake to have stretched over a region four times



FIG. 39.—Diagram-section to illustrate the propagation of an earthquake-wave and the mode of calculating the depth of its origin.

greater in extent than Europe. The earthquake of the 13th August, 1868, was felt in Peru for a distance of 2,000 miles.

15. (ii.) **Upheaval.**—After an earthquake has ceased the ground is sometimes found to have been raised above or depressed beneath its previous level. But besides these sudden and violent movements others have affected different parts of the earth in such a slow and quiet manner, as not to be perceptible at the time, but to be capable of detection by observation of their effects at the margin of the land.

16. Old harbours and sea-walls are sometimes found to stand now at some height above even the highest tide.

Islands once separated by a water-channel from the land are now joined to it. Caves evidently hollowed out by the sea may be seen far above the reach of the waves. Barnacles and sea-shells are found still adhering to rocks at heights of several hundred feet above the level where they lived and died. Terraces of sand and gravel, quite like the present beach, and containing sea-shells, occur at different heights above the sea (Figs. 40 and 41). These



FIG. 40.—View of an old sea-terrace or raised beach, with sea-worn caves on its inner margin.

terraces are old beaches, each of which marks a former level of the sea before the land rose to its present height. Many such terraces skirt the shores of Great Britain. They form a marked feature of the coast in the north of Norway (Fig. 41). On the western margin of South America they occur in great perfection, reaching sometimes

a height of 1,300 feet above the sea, where sea-shells, still in position, attest the amount of uprise.

17. By evidence of these various kinds it has been ascertained that many long tracts of coast-line are slowly rising from the sea. Thus the shores of Sweden at Stockholm and northwards appear to be upheaved at a rate varying from six or ten inches to two-and-a-half feet in a hundred years. Further north the island of



FIG. 41.—Raised sea-terraces of the Alten Fjord, Norway.

Spitzbergen is fringed with raised beaches up to a height of 147 feet. The coast-line of northern Russia and Siberia for hundreds of miles has been recently elevated out of the sea, as is shown by the raised beaches with marine shells. It would appear, indeed (see Lesson XII. 21, and XXVI. 21) that the tundras and steppes, which stretch from the Arctic Ocean to the Sea of Aral and the Caspian and Black Sea, are only an upraised

sea-bottom. At no distant date, therefore, the sea must have come far south, dividing northern, central, and south-western Asia from Europe. The shores of the Mediterranean afford local illustrations of uprise, while the great sandy desert of the Sahara, containing here and there scattered sea-shells, up to height of 900 feet, is a case of the recent elevation of a wide tract of sea-bottom.

18. (iii.) Subsidence.—While in some regions the slow, imperceptible movement of the ground is an upward one, in others it is downward. This may be shown by different kinds of evidence. Thus some seaport towns, as in the south of Sweden, contain under their present streets traces of older structures which are now below the sea-level. Stumps of trees, with their roots still imbedded in the soil on which they grew, are to be seen actually under the water of the sea, as at different parts of the coast of Scotland and the south-west of England.

19. But the most striking evidence of the subsidence of large portions of the earth's surface is supplied by the Coral Reefs of the Pacific and Indian Oceans. These are reefs or submarine banks of a kind of limestone, formed entirely by the growth of coral polypes. Sometimes they skirt the margin of the land at a distance of one or two miles, with a shallow-water channel or lagoon inside. The outside of the reef is its highest part, and there the large strong kinds of coral live which delight in the dash and play of the waves; the finer and branched kinds prefer the stiller water of the lagoon. These are called Fringing reefs. They skirt the east coast of Africa and occur also in the West Indies, and fronting the shores of Brazil. The accompanying figure (Fig. 42) shows a section of an island where the land (L) is skirted by a fringing reef (R).

20. A second kind of reef is called a Barrier reef. It

may skirt a continent, as in the north-east of Australia, where the great reef fronts the land for 1,000 miles ; or it may encircle an island, as at Tahiti. It rises from such deep water outside that its slope to the sea must be very steep. Inside it is separated from the land by a deep



FIG. 42.—Section of an island (L) with a Fringing coral reef (R).

lagoon channel. This structure is shown in Fig. 43. There may sometimes be more than one island inside a barrier reef.

21. The third form of reef has received the name of Atoll or coral island. It is a ring of coral rising out of the middle of a deep ocean, and having a breadth of about a quarter of a mile. The water outside is so profound close to the island that the outer slope of the reef down to the bottom of the ocean must be a submarine



FIG. 43.—Section of an island (L) with a Barrier reef (R).

precipice. Inside lies a lagoon of comparatively shallow water, full of the more delicate branching kinds of coral. The structure of an atoll is represented in Fig. 44, and its general aspect in Fig. 45. No land exists inside. The waves break off fragments from the outer edge of the reef

and pile these up above the ordinary limits to which the sea reaches. By degrees a narrow ring of land is formed on the reef. Seeds are carried to it and take root, and at last a human population arrives and finds shelter and the means of subsistence.

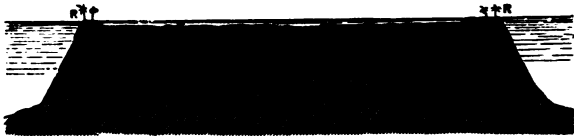


FIG. 44.—Section of an Atoll or coral island (R R) built over submerged land (L).

It has been ascertained that the reef-building corals cannot live in deep water, that indeed they do not go down further than about twenty fathoms. Consequently they cannot have grown up from the bottom of those



FIG. 45.—View of an Atoll or coral island.

profound abysses in which the atolls occur. Yet there can be no doubt that the coral rock formed out of their skeletons does actually continue down to the bottom. When they began to build they cannot have been more

than about twenty fathoms from the surface. The only inference to be drawn is that the bottom has been gradually sinking, while the little coral-builders have kept pace with the subsidence and maintained their reef at the sea-level.

23. The three sections just given (Figs. 42, 43, and 44) illustrate the stages in this subsidence. First we have the fringing reef, with its shallow lagoon and not very deep water outside. Then this shore reef passes into the barrier reef, with its deeper lagoon channel and much deeper sea outside. The land inside becomes less in height and extent as it settles down beneath the sea-level. At last it sinks out of sight altogether, and the barrier reef remains as an atoll.

In many instances, however, it would appear that the coral began to grow round the edges of tall volcanic peaks rising out of mid-ocean. In such cases it may not be necessary to suppose subsidence of the bottom. For the effect of the waves in cutting away the edges of the coral might combine with the growth of the coral itself to keep the outer margin of the reef tolerably steep, and even to grow out from the land upon a slope of coral debris.¹

24. From the evidence furnished by coral reefs it has been inferred that vast tracts of the sea-floor are sinking. Between Madagascar and the southern parts of Hindostan the Indian Ocean presents a series of atolls, some of which, like the Chagos Bank, are actually submerged. In the Central Pacific Ocean a still more extensive series stretches from the Caroline Islands to the further end of the Low Archipelago—a distance of fully one hundred degrees of longitude.

¹ This is the inference drawn from their recent observations by the naturalists of the "Challenger" Expedition, as I am informed by Mr. John Murray, one of their number.

LESSON XXIV.—THE WATERS OF THE LAND.

Springs and Underground Rivers.

1. In Lesson X. we traced some of the characters of a vast circulation of water between air, sea, and land. We found that from every water-surface on the globe invisible vapour is ascending into the air, where it is condensed into clouds, and whence it is returned to the surface of the earth again in rain, dew, snow, hail, or sleet.

2. The more we reflect upon the part which water plays in the general plan of our planet, the more importance shall we assign to it. The clouds form and melt and form again. Day by day, or season by season, the rain-showers reappear to moisten the parched soil. The brook and waterfall are ever rushing downwards, and yet they are continually fed with renewed supplies from above. The river still bears its broad breast of waters to the sea, as it has done ever since the earliest tribes of men settled upon its banks. The sea, though receiving the surplus drainage of all the continents, is not thereby raised in level, but yields to the air those abundant vapours which are borne back to the land, and, there condensed into running water, once more renew the downward journey to the sea. This continual coming and going of water may be looked upon as the pulsation of the very life-blood of our globe as a habitable planet.

3. Let it be remembered, too, that water enters largely into the composition of the bodies both of plants and animals. If its circulation were arrested our earth would cease to be the green populous planet which it is at present. Clouds, springs, rivers would disappear. Scorched by the fierce heat of the sun during the day,

and frozen by the intense cold of radiation at night, the land would become lifeless and silent.

4. As was shown in Lesson X., the moisture of the air returns to the land either in the liquid form; as water, or in the solid form, as ice. We have now to follow its further progress over the land under each of these conditions. In the liquid state the moisture chiefly appears as Rain; in the solid state as Snow.

5. When rain reaches the surface of the land, part of it sinks into the soil and the rest flows off into brooks and rivers, by which it is carried back to the sea. It is convenient to follow first the course of that portion of the rain which disappears underground. At first sight we might naturally suppose that it is permanently absorbed into the earth. But if it were withdrawn from the surface-circulation its abstraction would evidently tell upon the amount of water flowing over the land. Rivers and lakes would shrink in size or dry up altogether. Yet we do not find this to be the case. There must be some way in which the water that sinks into the ground is restored to the surface again. This takes place by means of **springs**, which are the outflow of the subterranean water from openings in the ground.

6. The intimate connection between ordinary springs and the rainfall is familiar to every one. We know that in a season of drought many springs and wells give a limited supply of water, or fail altogether, while as wet weather sets in they fill again. It is clear that they derive their supplies from rain-water which has percolated through the rocks beneath the surface. Such springs as have a deep-seated origin are less affected, or do not sensibly suffer at all from surface-changes, because they gather their stores from a wider area of subterranean drainage, where the effects of a scarcity of rain take longer to make themselves felt than is the case near the surface.

7. All rocks, even the hardest, are porous, and therefore pervious to water; the channels of brooks and rivers, the beds of lakes, and the floor of the sea are all more or less crack'd, so as to present openings for the descent of water. The rain-water, therefore, does not remain in the soil but sinks lower still and finds its way through the pores and joints of the rocks which lie below. Water is likewise supplied by lakes, rivers, and the sea, either oozing through the mass of the rock underneath, or availing itself of the open cracks, and pouring downwards into these, carrying with it sand and other impurities.

8. In sinking deep wells in some districts of France, leaves and other parts of plants have come up with the first gush of water from a depth of nearly 400 feet. These organic remains were comparatively fresh, and were supposed to have travelled in underground channels from hills 150 miles distant, and to have been three or four months on their subterranean journey. The same phenomenon has been observed in other places; sometimes even live fish have been brought up in borings from depths of 170 feet.

9. As the result of this constant percolation and descent of water from the surface, the rocks for some way down are in many places charged with moisture. Proofs of the almost constant presence of water are furnished in quarries, pits, and mines, in short, in nearly every place where any considerable cutting is made through rock. It is this underground water which forms one of the greatest obstacles in quarrying and mining operations. Before the introduction of steam machinery many coal-pits, after having been worked to a certain depth, had to be abandoned, from the impossibility of getting rid of the water. They were then said in the expressive language of the miners to be "drowned." The powerful pumping engines now to be seen everywhere in the coal-fields

point to the abundance of the water below ground, and to the labour and cost which are necessary for removing it.

10. Another and familiar illustration of the way in which water everywhere pervades the soil and rocks is to be seen in the sinking of wells. In most parts of the world these artificial cavities are dug out to serve as receptacles wherein the water which is soaking through the rocks may be collected. Wells may be successfully made even in places where it could hardly be supposed that water would be found. Thus on the borders of the African deserts, where little or no rain falls, and where therefore there can be but a scant supply of water from the surface, serviceable wells are dug. The French colonists of Algeria sink what are known as "Artesian wells" (Art. 18) on the northern margin of the great desert of Sahara, and the sandy tracts between Cairo and Suez yield water even so near the surface as at a depth of fifty feet. The existence of those fertile green patches called Oases in the midst of the barren deserts of Africa and Arabia is due to the rise of springs. Again in the valley of the Mahanadi and other Indian rivers, where in the dry season little or no rain falls, a hole dug out to the depth of thirty or forty feet is sure even then to be partly filled with water.

Hence the springs of a district do not always or necessarily obtain their water from the rainfall of the immediate neighbourhood. If that were the case there could hardly be perennial springs and wells in the African deserts, where rain is exceedingly rare.

11. To what depth the water will descend must depend greatly upon the nature and condition of the rocks at each locality. Very deep mines are often without water. When the Alps were pierced in making the railway tunnel between France and Italy, the rocks at a depth of more than 5,000

feet below the summit of Mont Cenis were quite dry. We need not suppose, therefore, that the water generally sinks to a very great depth. But here and there it no doubt does find its way down even into the intensely hot regions whence lava-streams proceed. It is this subterranean water which, as we saw in Lesson XXII., issues abundantly in vast clouds of steam, and plays so important a part in the arrangements of active volcanoes. It is probable that in spite of the plentiful discharge of steam at these volcanic openings, a part of the water which descends so far may be permanently lost, by being decomposed, and forced to enter into chemical combination with parts of the melted rocks. If this be the case, then the earth must be losing a little of its water; slowly and insensibly, indeed, but yet if continuously, with the probable result of ultimately reducing our planet to the dry and sterile condition of the moon.

12. Though the water which falls upon the land is distributed over the surface as rain, it does not reappear at the surface oozing everywhere from the soil. Sinking underground it finds its way along cracks and hollows of the rocks below, until it comes out again to the surface at certain points. Rain-water, though it falls over a whole district, at once runs off into brooks and larger streams, until it finally enters the sea. Somewhat in the same way the underground drainage is collected from many branching channels and brought out to the surface in springs.

13. A difficulty may sometimes be felt in understanding how the water having once sunk down can ever be driven up again. But we must remember that the springs which form its points of escape at the surface lie at a lower level than the ground from which the original supplies of rain have been drawn. If we try to realize to ourselves the manner in which the underground circulation is effected we see that it must be in one or other of two ways :

either by simple gravitation, as in what may be called Surface Springs, or by hydrostatic pressure, as in Deep-seated Springs.

14. (I.) In the case of Surface Springs the water which has been steadily flowing downward as well as onward in its underground course comes to a point where, owing to some depression of the land, it finds itself again at the surface. The subjoined woodcut will explain how such springs arise. A porous bed of rock (*b*), or one traversed with cracks or joints, lies nearest the surface, and allows the rain-water to soak through it down to a stiff impervious layer (*a*) by which the descent of the water is arrested. Unable therefore to sink further down-



FIG. 46.—Section across a valley to show how the simplest kinds of springs arise.

ward, the water flows along the surface of this lower bed. If a valley should happen to cut through these rocks, there will be a spring or line of springs (*s*) on the junction of the two rocks along the side of the valley. In the same way the rain which falls upon a mountain may sink underground and gush out in springs at the foot of the mountain. In springs of this kind the water merely descends in the ordinary way by gravitation, and issues where the surface happens to sink below the level of the subterranean water-channel. In heavy rain a good deal of water soaks through the soil into the nearest brooks without ever forming actual springs.

15. Wherever water accumulates among underground rocks, whether in the pores or in the open crevices,

so as to convert the rocks into subterranean reservoirs, it will rise in them up to the lowest levels at which it can find outlets to the surface, where it will appear in these surface- or land-springs. This water level must be reached before a well can catch a supply of water.

10. (II.) In Deep-seated springs, on the other hand, the water has in its journey sunk to a lower level than its point of escape, and has risen again by hydrostatic pressure, as in a syphon. Evidently the arms of a syphon may be as long as we choose to make them, yet while the one is longer than the other, and is supplied with water at the top, the water will continue to flow out from the top of the shorter arm. In like manner the subterranean

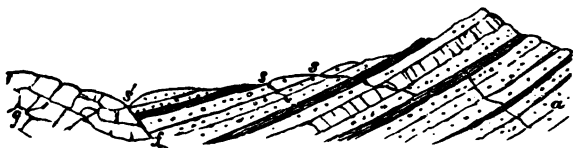


FIG. 47.—Section to show how deep-seated springs arise.

channel of a deep-seated spring may descend for hundreds of feet, yet the water which fills it will not cease to flow and to rise again to the surface. Having reached its greatest depth, often far below the level of the sea, the water accumulates there, saturating such porous rocks as lie in its way, until the pressure of the column of water behind it forces it up any fissure which may allow of its escape to the surface, and there it bubbles up as a spring. As in the accompanying figure (Fig. 47), the rain which falls on the high grounds, and is absorbed by the rocks and soil, flows down more or less permeable rocks which are arranged in different beds or *strata* (*a*). Taking advantage of the cracks which may opportunely present

themselves, as much of the water as these openings will admit rises through them to form springs, as at *s*, *s*. At *f* the strata are broken across by a great fracture or dislocation, which brings them against a hard massive rock (*g*). Here a considerable body of water may escape to the top at *s'*. Hence the direction of a great dislocation of the rocks is often traceable at the surface by a line of springs.

17. In nature the course which water takes under-

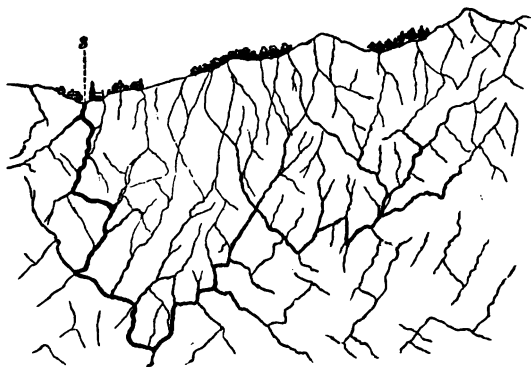


FIG. 48.—Section to show the intricate underground drainage which issues in a deep-seated spring. The numerous joints and cracks in the rocks lead the water at last into a main channel, by which it reascends to the surface as a spring at *s*.

ground is in reality for the most part much more complicated than is shown in this diagram. All rocks are abundantly traversed by divisional planes, called "joints;" they are likewise full of cracks, and they present many changes in texture. So that the water finds a most intricate network of passages through which it must make its way. It may many times come near to the surface and

then sink down again through other fissures, performing in this way a zigzag up-and-down journey before it finally issues in springs. A rough representation of this kind of circulation is given in Fig. 48.

18. On reflection we perceive that there must necessarily be many underground rocks which are permanently saturated with water. These, if we can reach them, will furnish an abundant and constant supply of water as pure as any that comes out in springs. Advantage is taken of this kind of knowledge to sink deep wells called "Artesian," from the old French province of Artois, where they have long been in use. The principle on which these wells are made is as follows :—When an impervious rock covers a porous one over a considerable district of country, the water

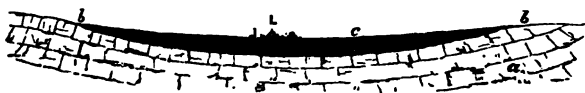


FIG. 49.—Section to the position of the water-bearing rocks below the clay at London.

which soaks into the lower bed will tend to accumulate there as in a great reservoir. If now a hole be drilled through the upper retentive bed, the water will rise in it at once, as it would do if a natural fissure existed there, and if the position of the bore be properly chosen, the water may ascend even above the surface of the ground. The water occasionally issues from these bore-holes with such force as to form a jet rising to a height of thirty feet or more above the soil. In the north of France the force of the jet has been made use of to drive mills. Many wells on this principle have been sunk in London and its neighbourhood. The annexed section (Fig. 49) across the district in which London lies shows roughly how the water which falls on the high grounds to the north and

south filters through the sands and gravels (*b*) at the top of the chalk (*a*), and is imprisoned by the retentive London clay (*c*).¹ When the wells are sunk down to this water-bearing zone, the water ascends in abundance. So many wells, however, have now been made that the level of the water in them has gradually sunk year by year, the consumption being thus somewhat more rapid than the supply to the underground reservoirs.

19. In trying to picture to ourselves the amount of water which is always circulating under ground, we ought not to measure it merely by what is seen coming out at the surface in well-marked springs. For in the first place we must bear in mind that the abundance of springs is really much greater than might at first sight appear. Much of the water which rises from under ground does not bubble up in the form of well-marked springs. When it reaches the surface it soaks through the soil or trickles over it in tiny runnels. On cultivated land such places are marked by greener patches of vegetation, or by tufts of rushes and a swampy soil. Sometimes, too, we may see them exposed even on ploughed fields during dry weather in spring time. From want of rain the bare soil gets dried and light in colour, but here and there it is diversified with dark brown patches. These point to places where the water is oozing out from below and soaking through the soil in spite of the drain-pipes of the farmer. In these and similar cases the water after coming up to the surface, sinks down again into the rocks and commences another underground journey.

20. In the second place we must remember that the natural flow of water from a higher to a lower level must carry much of the underground drainage to the rocks under the sea, so that when the water rises to the surface

¹ In the diagram the curve of the beds has necessarily been greatly exaggerated, there being in reality hardly a basin at all.

it will issue on the sea-floor instead of under the open air. Springs along the beach within tide-marks are sufficiently common, and many cases have been observed in the Mediterranean where powerful springs or even underground rivers rise to the surface of the sea at some distance from the shore, so that fresh water may actually be obtained at these places for the supply of ships. Vast numbers no doubt rise below low-water mark and remain unseen.

21. In some countries, where there are no rivers, and little or no rain, springs and artificial wells are the only sources from which a supply of water can be obtained by the inhabitants. But even in regions where both rain and rivers abound the usefulness of springs is hardly less apparent. Think for a moment what would happen if all the rain were immediately to run off the surface of the land without allowing any portion to sink under ground. Except where fed from melting snow the brooks and rivers would flow only after showers. Their channels would dry up in the absence of rain. It is the underground circulation which, by means of springs, keeps the watercourses of the land replenished, even in drought, and thus preserves the surface of the land fresh and green.

22. Rain is water nearly in a state of purity. In its descent it takes up a little air and some of the impurities which may be floating in the air (Lesson VI.). But these admixtures form only a minute proportion in rain-water, especially at a distance from the smoke and vapours of towns and manufactories. Let us, however, turn to the water of any spring, even the clearest and most sparkling, and we shall find that though it sank into the ground as pure or almost pure water, it reappears again with an impregnation of various substances, which can be made visible by boiling the water, when they remain behind as a film on the bottom of the vessel. They are chemically

dissolved in the water, and do not interfere with its clearness and brilliancy, nor in most cases do they impart to it any taste. They may be found in the water of every spring, but their quantity varies greatly. Sometimes they occur in such minute proportions as fifty parts in every million parts of water; in other cases they rise to as much as 32,700 parts in the million, which is about the same as the proportion of salts in some parts of the Pacific Ocean. (See Art. 30.)

23. Three questions naturally arise in reference to this remarkable impregnation of all spring-water: firstly, what are the substances present in the water? Secondly, how does the water obtain them? and, thirdly, what is the result of their constant removal by the water?

24. (i.) **Substances contained in spring-water.** If you take a glass of bright sparkling spring-water and let it stand for a time, you may observe minute bubbles adhering to the inner surface of the glass. These are air, or gas, which has been dissolved in the water. They rise by degrees and escape, and when they have done so the water ceases to have the same fresh brisk taste which it had when first drawn from the spring. In like manner, if the water is boiled it acquires an unpleasant insipid taste, not because anything has been put into it, but because the air or gas which it contained has been driven out. Hence it appears that at least one cause of the pleasantness of spring-water to the taste is the presence of gaseous materials.

25. Chemists have carefully examined these materials, and have found that they consist of oxygen, nitrogen, and carbonic acid, that is, the ordinary gases of the air, sometimes with the addition of sulphuretted hydrogen, which gives the rotten-egg taste to the waters in which it occurs. In many springs the proportion of dissolved gas is so small as to be hardly recognisable, even in a chemical

analysis, though in others it rises to ninety cubic centimètres in the litre.¹ In those springs which are marked by the abundance of the foreign substances dissolved in them the proportion of gas sometimes rises to upwards of 1600 cubic centimètres in the litre.

26. When the gaseous ingredients have been entirely driven off and the water is finally evaporated, a fine white film of some mineral substance is left behind. This residue is found to vary greatly in composition in different springs. It commonly consists of sulphates or carbonates of lime, soda, magnesia, and often includes chloride of sodium or common salt. When there is much residue the water is said to be *hard*. This is the case in most limestone districts : owing to the large proportion of lime in the water, more soap is needed in washing than in districts where the water is *soft*, that is, contains a small proportion of mineral matter. The softest water that can be had for domestic purposes is of course rain, for it contains little or no mineral admixture. Another proof of the presence of solid substances dissolved in spring water, is furnished by the crust deposited on the inside of kettles and boilers where hard water is used.

27. While all spring-water holds some mineral matter in solution, some springs contain so much that it affects the taste or becomes visible as the spring flows over the ground. These are known as Mineral Springs. The chief mineral substance in solution is often lime and this, when the water begins to evaporate as it trickles away from the spring, is deposited as a white crust upon stones, leaves, twigs, or any objects lying in its way. When the crust forms over moss it gathers round every little fibre and completely incloses it, so that though the plant dies the white crust remains and retains the outward form of the original moss. These *calcareous* springs, as they are

¹ There are 61.027052 cubic inches in a litre, and 16.386175 cubic centimètres in a cubic inch.

called, occur abundantly in limestone countries. In other springs, iron is the principal substance contained in the water, which has in consequence a strong inky taste. It often happens that as the water flows along, it deposits its iron as a yellow scum on the sides of its channel. Springs of this nature are known as *ferruginous* or *chalybeate*. Common salt is the characteristic of some springs, and when they contain a large proportion of it they receive the name of *brine-springs*. Sometimes they are nearly or quite saturated with salt, that is, they could not hold any more in solution. Many springs containing a considerable admixture of mineral ingredients are useful in certain diseases, either in the form of draughts or of baths. They are termed *Medicinal* springs. For example, the waters of Bath contain carbonate of lime, sulphate of soda, sulphate of lime, chloride of sodium, silica and carbonic acid ; those of Harrogate, carbonates of magnesia and lime, sulphate of magnesia, chlorides of sodium, magnesium, and lime, nitrogen, carbonic acid and sulphuretted hydrogen. The waters of Vichy are alkaline and acidulous from the quantity of soda, potash and carbonic acid which they contain. Those of Carlsbad abound in sulphates, those of Wiesbaden in chlorides.

28. In mountainous and snow-covered ground, spring-water may be met with scarcely above the freezing-point. From that extreme, every degree of temperature may be noted in different springs up to such boiling fountains as the Geysers. Evidently, the temperature of the spring must depend upon that of the rocks from which the water rises. If the water comes from melting snow and ice it will necessarily be cold. If it sinks deep within the earth, so as to come within the influence of the internal heat it will be warm. So that there seems good reason to regard the temperature as affording some

indication of the probable depth to which the water has descended and from which it has risen again.

29. Hot water has a greater power of dissolving most substances than cold water. Consequently, warm or thermal springs are often largely impregnated with mineral matter. Many medicinal wells are warm. Those of Vichy, for example, have a temperature of 111° Fahr., those of Chaudes Aigues, also in France, 178° Fahr., and those of Carlsbad, 165° Fahr. One of the substances particularly soluble in hot water is silica or flint, which has been already alluded to as occurring in the water of boiling springs, and as being deposited in the form of a hard crust round the orifices from which the water escapes, as in the Geysers of Iceland, the Yellowstone Park and New Zealand (Lesson XXI. Art. 24).

30. The proportion of mineral matter held in solution, differs widely in different springs. In common spring-water it may range from 50 to 400 or 500 parts in every million parts of water. But in districts where the water is "hard" the proportion may rise to 2,000 parts in every million. In mineral springs, the proportion is of course very much greater. Thus, in the Vichy waters, the solid contents are $\frac{5100}{1000000}$, those of Seidlitz contain $\frac{10000}{1000000}$; those of Püllna, in Bohemia, $\frac{2277}{1000000}$. The proportion in the Atlantic Ocean water, is about $\frac{25000}{1000000}$, while in the Dead Sea, it is $\frac{240000}{1000000}$.

31. When the discharge of water as well as the proportion of dissolved mineral matter is large, we see that the quantity of material taken out of the rocks below ground and carried up by springs, must be enormous. Thus, one of the hot-springs of Leuk, in Switzerland, having a temperature of 144° Fahr., is estimated to bring to the surface every year nearly nine millions of pounds of gypsum. If this mass of mineral could be all collected

it would form a square column, upwards of 650 feet high, and twenty-seven feet on each side. The brine spring of Neusalzwerk, near Minden, has been found to yield in one year enough of brine to form a cube of solid salt, measuring seventy-two feet. The wells of Bath, in like manner, are computed to yield annually enough of mineral matter to form a square pillar ten feet in the side, and eighty feet high.

32. (ii.) Origin of the Substances dissolved in Spring-water. We saw in Lesson VI. how rain in its descent takes up a little air together with some of the gases or other substances present in the air. One of these is carbonic acid. More of the same acid is obtained when the descending water reaches decomposing plant and animal remains in the soil. In some parts of the world, as in such old volcanic districts as Auvergne, in Central France, and the Eifel, in Rhenish Prussia, carbonic acid is given off in great quantities from subterranean sources, and is here and there brought up by springs, which like the Bad Tönnistein in the valley of Brohl, are so full of it that it escapes in copious bubbles when the water comes to the surface.

33. It is important to notice the sources whence underground water obtains its carbonic acid, because the presence of this acid gives the water great power in dissolving many mineral substances. How this solution takes place may be admirably illustrated above ground, by what happens at the arches of many bridges. If you stand beneath an arch, or the vaulted roof of a cellar, you may often observe that each line of mortar between the courses of masonry is marked by a sort of fringe of slender white stalks or pencils which hang downward. A little further observation will show you that at the point of each of these pendent stalks, a clear drop of water is hanging, which in time falls to the ground, and is slowly

replaced by another drop. Should the ground beneath the arch be undisturbed by passing footsteps, you will find that where the drops fall upon it, there is a white deposit like that of the roof above, growing up in little mounds, each of which is kept wet by the constant drip of the water. These white incrustations have of course appeared since the arch was built, and indeed they may be seen to increase in size from year to year. It is evident that they have come out of the masonry and that trickling water has had something to do with their formation.

34. Rain-water containing, as it does, carbonic acid, has the power of dissolving lime, and holding it in solution in the form of what is known as bicarbonate of lime. The mortar which binds the masonry together consists of lime. Being more porous usually than the stone or brick, which it cements, it allows a little of the rain which falls upon the bridge to sink through the seams and joints of the masonry. The water in its passage from the upper to the under surface of the archway, takes out a little lime, and carries it off in solution. As each drop appears upon the roof and hangs for a time before falling, it is diminished by evaporation, and losing part of its water and of its carbonic acid, cannot hold so much lime as at first. It is therefore compelled to deposit the surplus as a white film upon the roof. The drop then falls, and is succeeded by the next, which goes through the same stages, and thus the original ring of lime left round the edges of the first drop grows atom by atom into a long slender hollow tube or stalk like an icicle of stone. If it is undisturbed, it may lengthen until it actually reaches the floor, and its sides may be added to by further trickling water until it becomes a stout rod, or even a pillar seeming to support the roof. These hanging stalks or columns of lime are called *stalactites*.

35. But the drops do not leave all their lime behind

them on the roof. They carry a good deal of it with them when they fall to the ground. There they evaporate still further, and consequently deposit more lime as a white solid crust called *stalagmite*. This process of removal and re-deposit, seen so well in a small way in arches of masonry, occurs sometimes on a magnificent scale, in vast limestone caverns, as will be referred to further on in this Lesson.

36. If now in sinking through a few feet of masonry, rain-water can work such great changes, what may we not expect to take place beneath the ground where the water has to traverse vast masses of rock, and may have its proportion of gases, including carbonic acid, increased as it moves along! We should remember, too, how greatly augmented the solvent power of water must become, down in the hot interior, as is shown by the frequent large admixture of mineral ingredients in the water of thermal springs. There need be no wonder therefore that every spring contains more or less of these ingredients.

37. (iii.) Results of the removal of materials from below in spring-water. Let us now consider what must be the consequence of this universal percolation of water through the underground rocks, and of the removal and transport to the surface of so much of their solid substance. Evidently there are two points of view from which this question may be looked at: in the first, the effects on the surface, in the second, the effects below ground.

38. To realise in some degree the influence which the mineral substances brought up by springs have upon what takes place at the surface, we may reflect that the springs supply rivers and lakes, and thus, indirectly, the sea, with lime, iron, soda and other soluble ingredients. In fresh-water, as well as in the sea, there are large tribes of

animals which obtain the materials of their shells or skeletons from the lime or other minerals present in the water. If these mineral substances could be removed for a time, the result would be to destroy a large part of the denizens of our rivers, lakes, and seas. Man himself would suffer, not only by losing some of the most valuable of his supplies of food—molluscs (oysters &c.) crustaceans (lobsters, crabs &c.) and fishes, but from the mere absence in his drinking water of mineral substances, useful in keeping his body healthy. The failure of medicinal springs, too, would deprive him of an important element in the curative treatment of many diseases.

39. It is in its underground operations, however, and the effect of these upon the surface, that the removal of mineral matter by springs performs the most striking part of its work. Since every spring is busy continually dissolving and carrying up to the surface some of the solid substance of the earth's interior, and since the amount so carried by many springs is large enough to make even in a single year considerable hills if it were collected in the solid form, the result must be that cavities arise among the subterranean rocks. That this is really the case is best seen in limestone countries. Limestone is liable to be dissolved and removed by percolating rain-water in the same way as the mortar of an arch-way (art. 33). It is full of chinks and joints by which the water can find its way downward. Each of these passages is gradually widened by the solution and removal of the rock forming its sides. Hence in districts where limestone forms the uppermost rock, the ground is sometimes full of holes, which have been eaten out of the solid rock by trickling rain-water, and which lead down into numerous branching tunnels and chambers under ground.

40. One of the most remarkable examples of this kind of scenery is that of the Karst in Carniola on the flanks

of the Julian Alps. It is a tableland of limestone, so full of holes as to resemble a sponge. All the rain which falls upon it is at once swallowed up and disappears in underground channels, where as it rushes among the rocks it can be heard even from the surface. Some of the holes which open upon the surface lead downward for several hundred feet. Some turn aside and pass into tunnels in which the collected waters move along as large and rapid subterranean rivers, either gushing out like the Timao at the outer edge of the tableland, or actually passing for some distance beyond the shore, and finding an outlet below the sea. Here and there the labyrinths of the honeycombed rocks expand into a vast chamber with stalactites of snowy crystalline lime hanging from the roof or connecting it by massive pillars and partitions with the floor. Such is the famous grotto of Adelsburg near Trieste—a series of caverns and passages with a river rushing across them.

41. Still more extensive is the Mammoth Cave of Kentucky—a cavern about ten miles long, but with many ramifying passages which are said to have a united length of more than 200 miles. In the island of Antiparos a famous grotto lies 600 feet below ground, forming a spacious hall 300 feet wide and 240 feet high.

42. Partly from the solvent action of the descending water upon the sides of the chinks of the limestone, and partly from the falling in of the roofs of the underground passages, the surface of the ground in some limestone countries is so full of holes, and at the same time is so bare of soil, that it becomes a kind of barren and dry desert. All its rain, its springs and its rivers are withdrawn from the surface. In the course of time the ground has here and there subsided to such an extent as to form hollows in which the water collects into lakes. These however have no outlet at the surface. The water

issues from openings in the rocks to fill them and flows away by other openings of the same kind. The Zirknitz See in Carniola is a good instance of a lake of this kind. It is about five miles long and from one to two miles broad, but usually not more than from six to ten feet deep. Its bottom is said to be perforated with about 400 funnels or pipes through which the water ascends. In wet weather it rises to three times its ordinary height. But even at high water, it is so surrounded with high ground, that it cannot find any outlet at the surface and has to discharge its surplus waters down some of the innumerable caverns in the limestone (see Lesson XXVI., Art. 13).

48. From what has been said it appears that in some parts of the world (notably where limestone occurs largely) the surface of the ground is greatly changed by the chemical action of subterranean water. There is yet another way in which the circulation of water below ground affects the form of the surface. Where rain sinks through a porous sloping bed of soil or rock it sometimes forms a loose, watery layer underneath, which, by destroying the support of the overlying mass, allows the latter to slip down the slope and tumble into fragments below. This is called a landslip. Changes of this kind can of course only occur on the side of mountains, cliffs, ravines, or steep slopes generally, where movement by gravity from a higher to a lower level is possible. They are common along the sea-coast, many parts of the shore-line of the British islands being fringed with old landslips. When the slipped mass is large in extent and becomes covered with vegetation, it forms a strip of broken and picturesque ground in front of the higher cliff behind. Such is the undercliff of the Isle of Wight, and the long line of rough crags and grassy mounds flanking the steep cliffs of Antrim. In the latter case (Fig. 50) a great tableland of ancient hard lava beds (*b*)

rises from the coast in a line of noble cliffs resting upon layers of much softer and more porous rocks (*s*). Owing to the loosening of the support of the upper part of the cliff by the trickling of water between the beds in the



FIG. 50.—Section across the cliff and landslip of Antrim.

lower half, huge slices of the heavy solid lava-rocks have been launched down to the low ground. Many of these fallen fragments are themselves large enough to be called hills.



FIG. 51.—View of part of the cliffs and landslip of Antrim.

44. In hilly countries subject to heavy rains landslips are of frequent occurrence. The earth and upper layers of rock, saturated with water, are loosened and slide down

the slopes, carrying trees and fields to the valleys below, and piling up vast heaps of ruin there. In Sikkim, and other districts to the south of the Himalayan chain, the surface of the ground is being altered from this cause after every heavy fall of rain, vast spaces of mountain-slope many acres in extent being detached so as to sweep down, with their covering of forest, into the lower ground. Sometimes these fallen masses of earth and rock are thrown across a valley, so as to bar back the river and form a lake. But as the barrier consists only of loose rubbish, it is apt to give way to the pressure of the accumulating waters, which then pour down the valley with great force, sweeping everything before them, and desolating the district for many miles along their course.

45. When landslips take place in well-peopled and cultivated valleys they sometimes cause great destruction of life and property. Thus, in the valley of Goldau, Switzerland, in the year 1806, after a continuance of wet weather, a bed of rock, 100 feet thick, resting on saturated sandy layers, slipped down. The whole side of the mountain of the Rossberg seemed to be in motion. In a few minutes the descending mass had, with a terrible noise, rushed across the valley, burying five villages and about 500 people under a mass of ruined rocks 100 to 250 feet high.

LESSON XXV.—THE WATERS OF THE LAND.

Running water.—Brooks and Rivers.

1. Having in the last lesson followed the course of that portion of the rainfall which disappears into the ground, we have now to trace what becomes of the rest. Since

the surface of the land is higher than the sea, and slopes downward to the sea-level, the water which falls upon it from the sky cannot remain there but must in obedience to gravity seek the lowest level. This it can only do by moving downward till it flows into the sea. Were the land inclined from a central ridge with even slopes, like the roof of a house, we can suppose that the rain would run off in sheets of water. But instead of such uniformity it everywhere presents the most irregular surface. Even on what we might suppose to be a perfectly level and smooth piece of ground, there are innumerable little heights and hollows, which become at once apparent during a shower of rain, for then the hollows are marked by tiny runnels of water which course along them to a lower level.

2. Water-courses and River-basins.—Owing therefore to the unevenness of the surface of the land the surplus rain runs off into the hollows, down which it flows until it can descend no further. These hollows or channels which receive and conduct the drainage of the land are termed *water-courses* or *streams*. They vary in size from the merest rill or gutter up to the broad river which carries the drainage of half a continent to the sea.

3. Take the map of any large country or of a continent and notice the arrangement of the watercourses upon it. You observe that they are grouped somewhat like the branches and stem of a tree. In the low grounds, towards the sea, a river, in one single trunk, bends to and fro across the land. But further inland it divides into separate limbs, these again into smaller branches, and so on as far as the upper limit of the region drained by the river. This is the common plan on which the drainage of the land is arranged,—innumerable little runnels in the heart of a country, coalescing more and more as they descend until they all finally unite into a few main streams.

4. Let us in imagination trace the course of a large river from its beginning in the midst of a continent, to its end in the sea. Among the far mountains, where the sources of the river must be sought, the higher summits are covered, or at least streaked, with snow, while long tongues of snow and ice may often be seen creeping down the upper parts of the valleys. Perhaps the river issues from the melting end of a "glacier" (see Lesson XXVIII.). If so, it springs up at once as a tumultuous torrent of muddy water, and rushes down the valley receiving from either side innumerable minor torrents and runnels, which descend the rugged slopes, either from the melting edge of the snow, or from abundant clear bubbling springs. Or the river takes its origin in some single spring, not larger perhaps than many others in its neighbourhood, but which has been fixed upon from early times by the human population of the district as the true fountain of the river. Such a spring, either welling quietly from the ground or gushing out copiously, supplies the first little stream which dashes down its rocky channel, receiving from each side, as it descends, many tributary torrents, until after leaping from rock to rock in foaming cascades, and working its way through deep gullies, it reaches the more level part of the valley. In this first or torrent-part of its course the infant river is only one of many such streams by which the sides of the higher mountains are channelled.

5. But when it gains the valley it enters on a second and distinct portion of its journey. Its flow is less rapid, its channel less steep and uneven. It winds to and fro in many bends and loops across the flat parts of the valley, and rushes through the narrow gorges which occur at intervals. It still grows, by the addition of many smaller streams from either side, and becomes more and more like a true river as it rolls along. This valley part is by far

the longest and most important in its course. Here it receives most of its water, and puts on its distinguishing characters as a river. Its tributaries are no longer mere torrents or brooks, but rivers, sometimes as large as itself. But as these increase in size they become fewer in number, so that in the last stages of its journey, the main stream receives few or no affluents.

6. When at last, quitting the valley which has conducted it through the hills and the lower undulating country, the

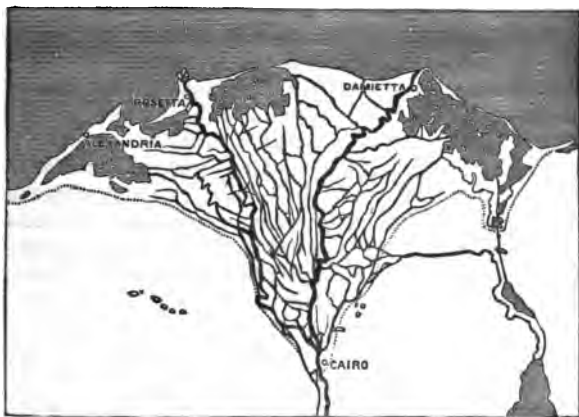


FIG. 52.—Delta of the Nile.

river reaches the low plains, towards the sea, it enters upon the last section of its course, that of the delta. Hitherto it has always been receiving but never giving off branches. But now, reaching the low level land of the delta, it begins to divide, sending off many arms which wind to and fro, and again divide among the swampy flats. So much may the river ramify in this part of its course, that it may enter the sea by many channels, or mouths,

often so nearly of the same size that it may be hard to say which of them should be called the chief. Moreover these intricate channels are constantly shifting their position as the river-water moves seaward. Hence what was once the chief mouth is now perhaps half-filled up, and the principal discharge takes place by another channel at some distance on the edge of the delta.

7. It will be noticed on a map, that except where flowing along some straight valley among mountains, most rivers do not run in a straight line for more than a very short distance. This is the case in every part of their course. The channel is continually winding from side to side. On a small scale you may observe the same arrangement of curving water-lines when the rain during a heavy shower is running down a sloping piece of roadway. At the upper part the runnels, though always rushing downwards, cannot, owing to the unevenness of the ground, descend in straight lines, but are deflected, now to one side, now to another, by little pebbles or bits of clay or other roughnesses on the road. Where straight ruts have been made for them by cart wheels they take advantage of these, and flow then in straight courses, as a river does in a longitudinal valley. But they escape before long, and resume their winding way as before, joining each other in the descent, and swelling the main runnel which sweeps down the road and eventually finds its way into some neighbouring ditch. What the pebbles, ruts, and other roughnesses on the road are to the rain, the uneven surface of the country is to the brooks and rivers. In either case the water seeks the readiest path of escape to lower levels. But this path is seldom the shortest. Every obstacle which the water cannot surmount or remove, serves to turn it aside, and, as such obstacles abound, the flow of the water is a continuous series of turnings and windings. Ridges and hollows, heights and valleys, turn

the streams, now to one side, now to another, until the waters find at last a rest in the great sea. Fig. 53 represents the serpentine curves of part of the course of the Mississippi.

8. Amidst these bendings it will sometimes happen



FIG. 53.—Windings of the Mississippi. The shaded part marks the alluvial plain.

that the river eventually cuts through the narrow part of the loop, and thus shortens and straightens its channel. The loop is gradually shut off from the river by the accumulation of sand and mud, and becomes a crescent-shaped pool, or lagoon of stagnant water. Instances of this kind

occur commonly along the courses of streams which flow through flat land, as in the case of the Mississippi (Fig. 53). In the delta of the Rhone they are called "Aigues Mortes."

9. One consequence of the frequent shifting of the direction of a river-course is that it becomes impossible to speak of the east or west side, or the north or south bank of a river. Obviously what is at one place the east side, may be successively the north, south, and west sides in the space of a few miles. Accordingly, it is usual to call one side the right bank, and the other the left bank, the observer being supposed to be looking down the river in the direction of its flow.

10. Another feature which a map makes clear is that each large river is the natural drain for a wide region. For instance, all the surplus rain and the discharge from melted snow and springs from by far the largest part of North America, find an outlet to the sea by a single river—the Mississippi. The space drained by this river is computed to be 1,244,000 square miles. This is termed the drainage-basin or catchment-basin of the river. The drainage-basin of the Ganges is estimated at 432,480 square miles; that of the Rhine at 75,000; that of the Thames at 5,162, of the Shannon at 4,590, and of the Tay at 2,090 square miles.

11. If you take a map of North America and trace a pencil line round the sources of all the streams which are tributary to the Mississippi, that line will represent what is called the water-shed¹ or water-parting. To the north of it you will find the basins of the Mackenzie,

¹ Sir John Herschel proposed to write this word *water-sched*, meaning "separation of the waters, not *water-shed*, the slope down which the waters run" (*Physical Geography*, p. 120). But the original meaning of "shed" was to divide, or part, and this use of the word still holds in Scotland, where a girl is said to *shed* her hair in the middle. The idea of disjunction is a secondary one, which has gradually come to be the common usage of the word.

and of the rivers which drain into Hudson's Bay, to the east the St. Lawrence and the smaller rivers of the Eastern States, while to the west the Fraser, Columbia, Sacramento, Colorado, and many lesser rivers carry the drainage of the Rocky Mountain slopes into the Pacific Ocean.

12. The water-shed of a country or continent is thus a line which divides the flow of the brooks and rivers on two opposite slopes. On many maps you will find it marked as if it were a ridge or mountain-chain. But in reality it does not necessarily coincide with the highest ground. Trace on the map of Europe, for example, a pencil line between the streams which drain to the Atlantic, Baltic, and North Sea on the one side, and those which drain to the Mediterranean, Black, and Caspian Seas on the other; you will thereby mark the general water-shed of the continent. You will find that the line, instead of running along and coinciding with any central mountain-chain, crosses all the great mountains, table-lands, and plains. Beginning at Gibraltar, it traverses the table-land of the Spanish Peninsula, crosses obliquely the chain of the Pyrenees, passes athwart the plateau of Central France on the right bank of the Rhone, runs through and across the ranges of the Alps, the Black Forest, and the Carpathians, and then descends into the vast plains of Russia, across which it winds in a general north-easterly course to the chain of the Urals. That the water-shed need not be a high ridge, may often be noticed even on low or gently undulating ground, where the same valley may have a stream at either end flowing in opposite directions, the water-shed being a quite imperceptible rise of the surface between them.

13. We may learn from a map some important lessons as to the form of a country by noting the line of its water-shed. That line seldom runs down the centre of a region

like the ridge on the roof of a house. Very commonly it lies much to one side, and winds in great curves as it traverses the land. Now the position of the water-shed, like the axis described in Lesson XIX., Art. 25, suffices to indicate the relative slopes of the two sides of a continent or country, especially where these two sides descend to the sea-level. Take South America by way of illustration. The water-shed of the whole continent lies near the western coast-line. The slope facing the Pacific must thus be very much steeper than that which looks towards the Atlantic. A heavy shower of rain falling on the water-shed will of course run off partly to the west and partly to the east. The westward portion, starting from a height of perhaps 10,000 feet, will reach the Pacific after a journey of not more than seventy miles in direct distance; while the other half, setting out from the same elevation, will have a journey of about 2,000 miles in a straight line before it can enter the Atlantic. The water-shed of Hindostan, south of the Gulf of Cambay, is another example of this one-sided position. On a smaller scale, Scandinavia, Great Britain, and Spain illustrate the same feature.

14. When a water-shed runs close to one edge of a continent, there is no room for large rivers on that side; these must flow on the opposite slope. America again furnishes an admirable instance of this obvious arrangement. In South America, for example, there is no river of large size flowing down the short slope into the Pacific, but on the east side the largest rivers of the world bear the drainage into the Atlantic.

15. *Sources of Rivers.*—Every shower of rain which falls, and every spring which rises, within the drainage basin of a river, may be regarded as one of the sources of the river. In many parts of the world indeed, such, for example, as the central and southern regions of India,

where there are dry and rainy seasons, and where the rivers do not take their rise in high snowy mountainous ground, the water which floods the streams during the wet months is mostly derived from the rain which runs off the saturated soil. In common language, however, the source of a river is understood to mean the point from which the head-waters of the main branch of the river take their rise. It is often hard to say which of the branches of a large river should be called the chief. One may be largest in volume of water, another greatest in length of course. One of the branches has usually been selected by the people of the country, and called by them the main stream.

16. Large rivers rise from various sources—springs, rains, melted snows or the ends of glaciers (see Lesson XXVIII.), and lakes. A great proportion of them may be traced up till their furthest little tributary brook is found gushing out as a spring from the side of some hill or mountain. In limestone countries, as we found in the foregoing Lesson, large rivers sometimes issue from the caverns by which the underground rocks are there perforated. Occasional or periodical rain directly supplies much of the water of many rivers. The Nile, for example, owes its annual rise to the heavy rainfall of the wet season among the mountains of Abyssinia. The snow-fields of the higher mountains furnish unfailling nourishment to many of the largest rivers of the globe. Thus in Europe, the Rhine and the Rhone take their rise from the melted snow and ice of the Alps. In Asia the rivers of Northern India descend from the melting snows and glaciers of the giant chain of the Himalaya. In North America the abundant patches of snow which mark the higher parts of the Rocky Mountains supply some of the water which drains from the west into the great valley of the Mississippi. Here and there among

the higher summits of the land little hollows arrest the first runnels of melted snow or of springs, and form little lakes, out of which the infant waters of important rivers flow. Or a lake, formed on lower ground by the confluence of several tributaries, fills a vast hollow of the land, whence the united drainage escapes by one large river, like the Rhine from the Lake of Constance.

17. Proportion of the Rain-fall carried to sea by Rivers.—Since the size of rivers depends upon the amount of rain-fall or snow-fall within their drainage-basins, we may naturally inquire how much of the total moisture discharged from the air upon the land is actually returned to the sea by the rivers. The proportion between rain-fall and river-discharge has never been very satisfactorily determined, but is said to vary from $\frac{1}{3}$ to $\frac{1}{4}$; that is to say, only about a third or a fourth part of the water which falls upon the land as rain or snow is carried off by the streams. The greater portion is returned to the air again by evaporation. From the moistened soil, from every surface of snow or water, from each spring, brook, lake, and river, vapour is continually passing into the air. The rivers, therefore, do not bear to the sea even all the water poured into them, for they are continually losing water by evaporation from their surface.

18. The diminution of rivers, the drying up of brooks, and the cessation of springs during seasons of drought, show how dependent is the flow of water over the land upon its circulation through the air. In countries such as Britain, where heavy rains may occur at any time of the year, the rivers are subject to irregular increase. In those regions, however, where a wet and dry season succeed each other at certain intervals, the rivers have their periodical rise and fall. The most familiar example of this regularity is that shown by the historical river—the Nile. Egypt, through which this river flows in

the lower part of its course, is a singularly dry country. Rain seldom falls there, and yet every year, and with such regularity that almost the very day of the change may be foretold, the river begins to rise, and continues to do so until the low plains on either side are overflowed. It then slowly subsides, leaving a film of fine mud over the ground, and resumes its former channel. This remarkable feature puzzled the ancients very much. They accounted for it by supposing that the Nile rose among snowy mountains far to the south, and that the inundations were caused by the melting of the snows. But in recent years the true cause has been ascertained. The high and rocky table-land of Abyssinia is visited by heavy rains during the months of March and April. The numerous gullies and gorges which intersect that rugged country, and which were previously quite dry, are then filled with torrents of water, which rush down and swell that branch of the great river known as the Blue Nile. It is these annual rains in the far highlands of Abyssinia, therefore, which cause the regular inundation of the rainless plains of Egypt.

19. In countries liable to heavy periodical rains, the relation between the flow of the rivers and the rain-fall is made strikingly clear. The Mahanadi River in Central India, for instance, has its basin within the area to which the south-west monsoon brings the copious rains alluded to in Lesson X. Falling upon high rocky ground, the rain at once rushes down by innumerable channels to swell the main river, which, unable to find in its channel room for all this water, overflows the surrounding country, producing great havoc in its course. But in the dry season, from February to May, the river is low, for then little or no rain falls, and it depends for its water upon the supplies which it receives from springs.

20. Another illustration of the relation between the

rain-fall and rivers is supplied by those singular dry ravines and gravel-tracks so common in Syria, Arabia, and the Valley of the Euphrates, called "wadys." Those within tracts which have a rainy season are turned into water-courses during the wet part of the year. But over a great part of the region where they occur, little or no rain falls, so that they remain constantly dry. That they were once the channels of brooks and rivers cannot be doubted. It is apparent that a change has taken place in the climate of these districts, partly caused or at least aggravated by the destruction of the ancient woods and forests, and the abandonment of the cultivation of the soil. The rains which once refreshed the land have ceased to fall, the river-beds which carried the surplus water to the sea are now dry, and the valleys are parched and barren.

21. In a country which is subject, like Britain, to heavy rains at different and uncertain times of the year, the rivers are liable to be swollen on any day; and as heavy rain often succeeds dry weather, the rivers may pass rapidly from a state of low water to full flood. Where, as in the case of the Nile, the rain comes regularly at the same season, the river swells and falls again with remarkable slowness and uniformity. But there is another cause of the regular, as well as irregular, flooding of rivers. It is evident that where a river draws its supplies of water in great part from the melting of snow among the mountains, it will have a larger volume in summer and autumn than in winter and spring. The Rhine and Rhone, for example, which take their rise among the snows and glaciers of the Alps, fill their channels with water during the dry hot weather of July and August, and shrink in size during the cold and often wet months of the year. Besides this annual increase and diminution, these rivers are liable to occasional and sometimes disastrous floods in summer, not caused by heavy rains, but by dry and warm

weather. When the warm south wind called *Föhn* blows upon the snowy slopes of the Alps, it causes a rapid thaw. Torrents of milky snow-water dart down the sides of the mountains and feed the brooks, these soon swell and rush forward into the rivers, which rapidly rise and sweep down their valleys, often carrying away bridges, inundating meadows and strewing them over with gravel, or even breaking down the houses in the lower parts of villages and towns built upon their banks.

22. Flow of Rivers.—Standing by the bank of a broad rushing river, we ask ourselves sometimes at what rate the water is moving, and how much passes before us in an hour or day. In most cases we shall probably guess the rate of motion to be faster than it is, for the broad stream, with its eddies and gurglings of water, impresses us with the idea of rapidity as well as of volume. The rate of flow is determined by the angle of slope and the volume of water. Measuring the average angle of slope down which rivers flow, we learn that it is considerably less than we might have expected it to be. The average declivity of the larger rivers of the continents probably does not exceed two feet in the mile. The Missouri has a mean fall of 28 inches in the mile; but the Volga does not slope more than 3 inches in the mile. To be easily navigable, a river should not have a mean declivity of more than 10 inches in the mile or 1 in 6336. Of course, in their mountain-track, streams have much higher angles of descent. That of the Arve at Châ-mounix is 1 in 616, and that of the Durance varies from 1 in 467 to 1 in 208. But such rapid declivities give rise to the motion of torrents rather than of rivers.

23. A moderate rate of the flow of a river is about $1\frac{1}{4}$ mile in the hour; that of a rapid torrent does not exceed 18 or 20 miles in the hour. The larger rivers of Britain have a velocity varying from about 1 mile to about

3 miles in the hour. Hence we may walk by the margin of a river and easily outstrip the rate of its current.

24. But the water of a river does not flow in every part of the channel at the same rate. Owing to friction against its bed, the river moves slower at the sides and bottom than in the middle. On a small stream you can easily prove this by throwing in pieces of wood and watching how much faster those travel which fall in the middle than those which have lighted near the edge. Evidently, therefore, the addition of more water to the same channel will increase the rate of motion of the stream. When two rivers join into one their united streams may occupy a channel no broader than one of them did before, but even without acquiring any increased slope in its bed, it runs faster, because the water has now the friction of only one channel instead of two. For the same reason, a river confined in a deep narrow gorge runs faster than where, with a like declivity, it spreads over a broad gravelly bed. When this cause is kept in mind we can understand the meaning of the curious fact that a river is sometimes not increased in breadth even by the influx of large tributaries. Thus the Mississippi becomes no wider even after it receives its largest affluents. Eighteen hundred miles above its mouth it is 5,000 feet, or nearly a mile, broad. But at New Orleans, where it enters the Gulf after having been successively fed by the vast streams of the Missouri, Ohio, Arkansas, and Red Rivers, it is only 2,470 feet broad, or less than half a mile, and yet in that comparatively narrow channel the drainage of nearly half a continent passes out to the sea.

25. Volume of Water discharged by Rivers.—Measurements and estimates have been made of the amount of water carried annually into the sea, or hourly past certain places by different rivers. The Mississippi, for example, has been found, after a careful survey of its operations, to

discharge annually into the Gulf of Mexico no less than 21,300,000,000 cubic feet of water, enough to make a lake as large as the whole of England and Wales and twelve feet deep. The Danube sends into the Black Sea 207,000 cubic feet of water every second. The Tay, though not the largest river in the British Islands, brings down more water into the sea than any other, its annual discharge being estimated at 144,020,000,000 cubic feet.

LESSON XXVI.—THE WATERS OF THE LAND.

Lakes and Inland Seas.

1. OWING to the inequalities on the surface of the land, the water which falls upon it cannot always flow off at once into the sea. Hollows occur which intercept the drainage until it fills them and flows over from the lowest parts of their margins, or otherwise escapes. Such water-filled basins are called lakes, but when they are of large size and are occupied by salt water, they receive the name of inland seas.

Lakes.

2. At first we might expect to meet with lakes indifferently on any part of the earth's surface. Any good general map of the world, however, shows that they are not distributed altogether at random, but are chiefly developed in certain regions. By considering this development we can better understand how the hollows came to exist at all, and why they should be where we now find them.

3. (1) First, then, in the northern hemisphere an extraordinary abundance of lakes may be observed to be scattered over the northern parts of Europe down to about the 52nd parallel of latitude, and in America to about the 42nd parallel. In some districts within these regions there

seems to be almost as much water as land. In Finland, for example, nearly a third of the country is covered with lakes and marshes, and as there are no mountain-ridges, and no dominant lines of valley, the undulating surface looks like that of a land which has been half submerged.

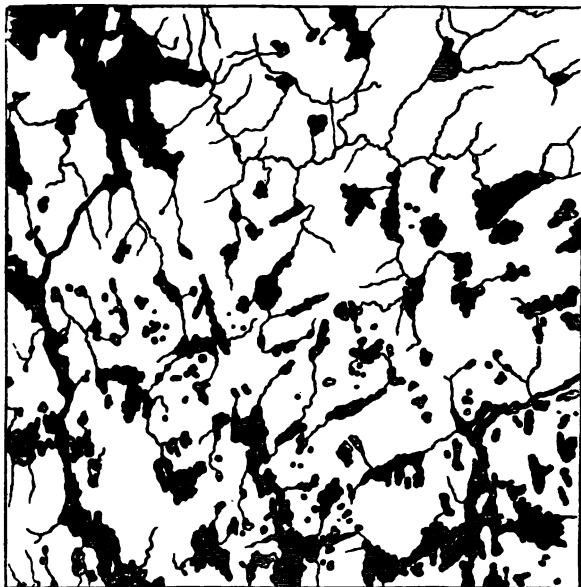


FIG. 54.—Part of the Island of Lewis, illustrating the abundant lakes of the north-west of Europe.

In the north-west of Scotland similar scenery occurs ; from some of the hill-tops there the undulating surface of the low country is seen to be plentifully strewn with lakes. In the accompanying figure (Fig. 54) a representation on the

scale of one inch to a mile is given of a portion of the surface of Lewis, one of the islands of the Outer Hebrides, from which some idea may be formed of the abundance and irregular distribution of these sheets of water. Again in North America, the British possessions and a large part of the north-eastern States are plentifully dotted over with lakes, varying in size from such vast basins as that of Lake Superior down to mere pools or tarns.

4. In these regions, the lakes not being confined to the valleys, but lying indiscriminately over the surface, may be found at any height from the sea-level to near the crest of the land. The ground is generally undulating rather than

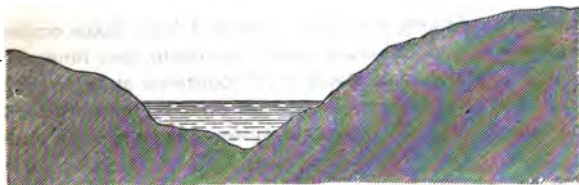


FIG. 55.—Section of a lake-basin excavated in solid rock.

hilly, and seldom rises more than a few hundred feet above the sea. The abundance of lakes in the northern portions of the northern hemisphere has been associated by many geologists with the fact that these regions have been covered with sheets of ice which, grinding down the surface of the land, have left it strewn with clay, gravel, and sand. How ice performs this kind of work is described in Lesson XXVIII. Many of the lakes lie in hollows excavated out of the solid rock (Fig. 55), which still retains abundantly the scratches and groovings made upon it by the movement of the ice. Other sheets of water are inclosed among the heaps of debris left on the land when the ice melted (Fig. 56).

5. (2) On further examination of the map it will be noticed that lakes occur more or less abundantly among mountains. Take Europe as an illustration. Even in the comparatively low mountain groups of Scotland, Cumberland, and Wales, lakes abound and form one of the great charms of the well-known scenery of these



FIG. 56.—Section of a lake-basin lying in a hollow of superficial detritus.

districts. Among the Alps a series of large lakes occurs on each side of the main axis of the chain, and innumerable minor sheets of water occur scattered at all heights among the central mountains, up even to the borders of the snow-line. All mountain systems have not the same



FIG. 57.—Section of a lake dammed up by a barrier of earth or gravel.

abundance of water-filled hollows ; in some, indeed, few or none occur. Lakes of this class may be hollows formed during the elevation of the mountains (see Lesson XXIX) ; or, like those referred to in Art. 4, may either have had their basins scooped out by glaciers or formed by the irregular piling up of ice-borne debris, as shown in Fig. 57.

In most volcanic districts lakes occur filling cavities which have been formerly blown open by explosions from below.

6. (3) A third series of lakes may be observed to occupy depressions on the table-lands. The most remarkable examples are furnished by the great lakes of Equatorial Africa. Of these the Victoria Nyanza lies at a height of about 3,300 feet above the sea, and is believed to cover an area of not less than 30,000 square miles. On the same continent a vast depression, with many small lakes, extends westward from the Nile valley, and stretches between the southern limits of Barbary and the country drained by the River Niger. Another smaller hollow lies in the southern part of the continent, and contains some small lakes, of which Lake Ngami (2,900 feet above the sea) is the largest. On the great table-land of Asia numerous lakes occur over Thibet, Turkestan, and Mongolia.

7. It will be noticed that many of these table-land lakes have no outlet. They lie in hollows below the general level of the surrounding country, and while they receive supplies of water from the streams which drain into them, no river escapes even from the lowest parts of their borders. Now almost all of them are filled with salt water. The reason of this is found in the fact, which has already been considered in Lesson XXIV., that water flowing over or through rocks dissolves out of them some of their soluble ingredients. Among the substances so removed, common salt, sulphate and carbonate of lime, and sulphate of magnesia are of common occurrence. These dissolved salts are carried down by rivers, and in most cases find their way ultimately into the sea. But in the depressions of the great table-lands the water flows into hollows from which it cannot escape except by evaporation. These hollows may be compared to great evaporating vats or

troughs, like those in which sea-water is boiled down in the manufacture of salt. The water passes off into invisible vapour, but leaves its various dissolved mineral substances behind. Hence year by year these lakes and pools become salter. When they dry up they leave a crust of salt upon the ground which they once covered. The soil is in such places so impregnated with salt, that plants will not grow upon it, and its arid sandy surface stretches for leagues as an inhospitable desert.

8. Where a fresh-water lake does occur in these regions it will be found to have some outlet by which its surplus water is removed, so as to prevent increase of saltness. Lake Chad in Central Africa lies in a hollow from which no river escapes to the sea, and it was believed in consequence that this fresh-water lake had no outflow, and was thus an exception to the general rule. More recently, however, a river has been found opening from its north-eastern margin and carrying the overflow along a wide valley, in which the water is finally dried up amid sandy wastes.

9. In the eastern countries bordering the Mediterranean many salt-lakes occur as well as ground incrustated or impregnated with salt. The naturally-formed salt has been used from time immemorial by the dwellers there. A little reflection on the mode of its formation and its composition explains the meaning of a curious passage in the Bible which is not in itself very intelligible. It occurs in the Sermon on the Mount. "If the salt have lost his savour, wherewith shall it be salted? it is thenceforth good for nothing, but to be cast out, and to be trodden under foot of men." (St. Matthew, v. 13.) Were the substance here spoken of our common salt (chloride of sodium), it would be difficult to explain how it could possibly lose its taste without ceasing to exist at all. Its taste is to us as essential a character

as its chemical constitution. But no doubt the substance referred to was the white incrustation obtained from the sides of the salt lakes and the bottoms of the dried-up saline pools. Now this incrustation, besides common salt, contains also magnesia, lime, and other ingredients. In short it resembles what may be obtained by the evaporation of sea-water. Of these various components common salt, as was shown in Lesson XIII. Art. 5, is the most soluble. It is the last to appear when the water is evaporated, and the first to disappear when moisture is supplied again. We can see therefore that the white natural crust used by the people of Syria for salt, being kept for a time and exposed to damp or rain, might lose all its salt. The more insoluble residue, consisting of gypsum, carbonate of lime, &c., though in appearance unchanged, yet having little or no taste, would be quite useless for the purposes for which the salt had been gathered. The question, therefore, might well be asked—"If the salt have lost his savour (or saltness), wherewith shall it be salted?"

10. (4) A fourth series of lakes occurs along many parts of the margin of the land where the ground is low and consists of soft sandy, clayey, or gravelly materials. These maritime sheets of water are known by the name of lagoons. In Europe they fringe all the Prussian shores of the Baltic, reappear on the west of Denmark, Holland, and Belgium, and are found at intervals along the northern coasts of the Mediterranean Sea, from the east of Spain to the western shores of Greece. In Asia they extend for hundreds of miles along the eastern and western sides of the peninsula of Hindostan. In America a long line of them skirts the Atlantic sea-board of the United States.

11. Lagoons along the sea-margin are for the most part shallow and narrow, running parallel with the coast, from which they are separated by a strip of low land formed

of sand, gravel, or other loose materials. When the sea flows into them the waters are salt or brackish. When they only communicate with the sea by a narrow outlet, or when they have no outflow, but soak through the porous bar which banks them out from the sea, they are fresh.

12. Most lakes derive their water from streams which flow into them. In Fig. 54, for example, this connection is well shown, though there the feeders of the smaller lakes are too minute to be shown upon the map. Many valleys contain chains of lakes along their course. A river flowing in one of these valleys appears alternately to contract and expand, having a comparatively rapid flow where it takes its own river-form, and losing itself in still water when it enters a lake. It would seem that this arrangement was formerly commoner than now, for the rivers have frequently filled up their lakes, and now wind to and fro across flat meadows where the lakes once stood.

13. But lakes likewise derive their supplies partly from springs which rise beneath them. In some cases, indeed, the whole supply comes from such underground sources. The Lake Zirknitz near Trieste (Lesson XXIV. Art. 42), affords an excellent illustration of this feature. It comes and goes with the seasons. After long drought it disappears, but when heavy rains fall on the surrounding mountains of Carniola, which are formed of remarkably honeycombed limestones, the water sinks out of sight, and after filling the underground passages, comes out with a roaring noise from the funnels and caverns which open upward into the hollow of Zirknitz. This hollow is then converted into a lake. No stream flows out of it, since both its supply and its overflow pass away by underground channels. This example shows how lakes, like rivers, depend ultimately upon the rain-fall for their water, and are apt to vary in level with the wetness or dryness of the seasons. In Northern Africa also, the

Sebka-el-Faroon, a hollow 100 miles in length, lying to the south of Tunis, at a level of several feet below the Mediterranean, is in winter covered with water to the depth of two or three feet. Having no outlet, and undergoing rapid evaporation during the parching summer of that hot climate, the lake disappears and leaves a salt-crusted floor.

14. Some of the largest fresh-water lakes in the world are those of North America, Lake Superior alone covering an area of 23,000 square miles, its surface being 627 feet above the level of the Atlantic, and its average depth nearly 1,000 feet. Another series of vast sheets of fresh-water lies in the eastern part of the table-land of equatorial Africa, and forms the source whence the Nile and the Congo take their rise. In the heart of Asia, Lake Baikal stands at a height of 1,363 feet above the sea, and covers a space 370 miles in length, by from 20 to 70 in breadth. On these vast sheets of water storms arise on a scale hardly inferior to that of the sea itself. Fresh-water, being lighter, is more easily stirred by the wind than salt-water. It is soon raised into ripples, and when deep and wide, and driven onward under the pressure of a continuous high gale, it rises into large waves, which roll across and burst in foam against the windward shore, heaping up gravel and sand, or cutting down the cliffs, as is elsewhere done by the waves of the sea (Lesson XVIII. Art. 11).

15. The depth of lakes varies almost within as wide and indefinite limits as their size. The form of the land round a lake generally affords an indication of the probable depth of the basin. Thus if the ground is low and slopes gently into the water, we may infer with some confidence that the lake must be shallow. If on the other hand the ground rises steeply out of the water, as when a lake fills a valley between two precipitous mountain-ridges, we may be prepared to find that the lake is deep. Several of the

Alpine lakes attain a great depth. Thus Lake Como is nearly 2,000 feet deep, and the Lago Maggiore 2,800 feet. In both of these lakes the bottom sinks below the level of the sea, that of the Lago Maggiore being 2,149, and that of Como 1,318 feet below the surface of the Mediterranean. But they all have outlets, and their waters are fresh. In Scotland, Loch Ness forms a remarkable feature in the great valley which cuts the Highlands in two, its surface being about 70 feet above the sea, and its greatest depth 810 feet. The bottom is therefore not only below the sea-level, but actually a good deal deeper than any part of the North Sea between Scotland and Denmark.

16. Little has yet been done in exploring the depths of these profound hollows on the land. Some of them, like Loch Ness, were at a comparatively recent period filled by the sea, and though probably all trace of the salt-water has long since been removed, perhaps some lingering forms of life may have survived the change, and may exist still in the undisturbed abysses, somewhat in the same way that some shore plants, such as the coast-plantain (*Plantago maritima*) are met with on the tops of the Scottish mountains.

17. It would be interesting also to know the distribution of temperature in the water of these deep lakes. According to observations made in Loch Lomond, in Scotland, the depth of which is about 600 feet and its surface 25 feet above the sea, it appears that a tolerably constant temperature of about 42° Fahr. characterises the lowest stratum of water for 100 feet or so above the bottom. The cold water of winter must sink down, and as the sun's rays can sensibly heat the water only for a short way beneath the surface, the temperature of the deeper parts of such lakes must be kept permanently low. In the Lake of Geneva in autumn, while the surface water showed a temperature of 78° Fahr., the bottom water, at a depth of

950 feet, marked $41^{\circ}7$. Similar observations on the other deep lakes of Switzerland and Northern Italy show that they all have a permanent mass of cold water at the bottom. Yet the mean annual temperature of the surface water of the Lake of Geneva as it issues into the Rhone is nearly 4° warmer than that of the air. Further south the Lago Sabatino, near Rome, was found to have a temperature of 77° at the surface, but one of 44° at a depth of 490 feet. Lakes, like the sea, exercise an important equalizing influence on temperature, preventing the air around them from being so much heated in summer and so much cooled in winter as would otherwise happen. (Lesson XXXI., Art. 20.)

18. One great and useful office of lakes is to regulate the flow of the rivers which issue from them. They receive the water discharged by heavy rains and rapidly-melted snows, spread it over a large surface, and allow it gradually to escape by the outflowing river. In this way they prevent the occurrence of those sudden and destructive floods, which, in the absence of such natural reservoirs, are apt to occur in all countries subject to copious rain, or where large masses of snow may be quickly thawed. Another part played by lakes is that of arresting the gravel, sand, and mud brought down by the streams which flow into them. These materials, which so often discolour the tributary brooks and rivers, fall to the bottom when the currents are checked by entering the still lake-water. The lakes in this way filter the rivers. The Rhone, for instance, is a muddy stream where it enters the Lake of Geneva, but at the lower end, where it quits the lake, its water is as pure and limpid as that of a spring. Its sediment has been dropped upon the bottom of the lake, which must consequently be slowly rising. The Lake of Geneva is being gradually filled up in the manner already referred to in Art. 12.

Inland Seas.

20. From what has been said in a previous part of this Lesson, it is not difficult to trace upon a map those areas of the earth's surface where the lakes are filled with salt-water. In every tract of land into which rivers flow without any outlet, and where the water must pass off by evaporation, we may expect to find brackish or salt lakes. A salt lake, therefore, need not necessarily have been once connected with the main ocean. So constantly are salts present in fresh-water, that any fresh-water lake where the only escape for the water is by evaporation, will eventually become salt. The salt lakes on the table-lands of Asia and Africa were, no doubt, fresh at first, and have gradually grown salt as the process of evaporation has been continued century after century. In North America, too, the Great Salt Lake in Utah, lying at a height of 4,200 feet above the sea, and covering a space of 75 miles in length, by from 15 to 40 in breadth, together with many other smaller salt lakes in the same region, show the same effect.

21. But in some parts of the world there exist sheets of salt-water which can be shown to have been once connected with the main body of the ocean, from which they have been separated by subterranean movements. Such Inland seas may have their surface below the level of the sea, or they may have been carried up together with the land around them, so as to lie at a higher level than the surface of the main ocean. By far the most remarkable of these is the chain of inland seas and smaller sheets of salt-water which extend from the Black Sea and Sea of Azov, eastward by the basin of the Caspian and Aral Sea, and thence northward through the vast steppes and mossy barren plains, or "tundras"—a low region

which has been elevated, at a comparatively recent time, out of the sea. The Arctic Ocean, as already stated (Lessons XII., Art. 21, and XXIII., Art. 17), is believed at one time to have run southward as a broad sea up to the mountains of Persia, a distance of more than 3,000 miles. Owing to some of those underground movements already described, that ancient mediterranean arm of the sea which separated Europe and Asia has been raised into land. But in the deeper cavities of the depression portions of the sea still remain, retaining even yet the marine shells, fishes, and seals which abounded in the water before the elevation of its bed. Of these relics the largest and most important is the Caspian Sea, which lies 84 feet below the level of the Black Sea, is from 2,000 to 3,000 feet deep in the central parts, and covers an area of about 180,000 square miles. It receives the drainage of the whole of the south-east of Russia in Europe by such important rivers as the Volga and the Ural, but it has no outlet. So large is the mass of fresh-water poured into the Caspian, that the saltness of the greater part of that sea is not more than about one-third that of ordinary sea-water (Lesson XIII., Art. 6). But along the shore numerous lagoons occur, where, in the dry and hot weather of summer, so much evaporation goes on that the water becomes intensely bitter and salt, and saline incrustations form at the bottom and on the shores. On the east side of the sea lies the wide but shallow Karaboghas Bay, which may be looked upon as a vast evaporating basin. A current is always passing in through the narrow opening, but there is said to be no compensating under-current outwards, so that, as the water-level does not rise, all this constant inflow must be supplying the loss by evaporation. The bay, therefore, grows every year salter. A solid layer of salt forms on the bottom of some of the shallower parts of its shores, and the water is there so saline that a

cord on being let down into it and pulled up again is immediately crusted over with salt. Seals, which used to flourish there, are said to have been driven away by the increasing saltiness of the water.

22. The Sea of Aral fills another of the hollows lying in the vast depression between the European and Asiatic high-grounds. It is a lake of brackish water 265 miles long and 145 broad. It is said to be at a height of only 33 feet above the level of the sea. On its southern side it receives the Oxus, which carries into it the drainage from the northern slopes of the great chain of the Hindu Kush mountains. It likewise obtains supplies of water and mud from the Jaxartes, which takes its rise among the lofty Thian-Shan mountains. Yet the Sea of Aral, like the Caspian, has no river flowing out of it. It loses by evaporation as much water as it receives. Indeed the loss from this cause would seem at present to be greater than the supply of water, for the sea is said to be sensibly decreasing in size.

23. Still further north smaller salt-lakes are scattered over the plain. That these are likewise parts of the vanished sea which once filled all that broad and long tract, seems to be put beyond doubt by the fact that they still contain shells and other kinds of marine life. Traces of dead shells are likewise met with all over the plains, and it has been observed that among these the characteristically Arctic species become more numerous towards the north.

24. The valley of the Dead Sea is remarkable as being the most depressed on any part of the land of the globe. The surface of that sheet of water is 1,298 feet below the level of the Mediterranean Sea. The water, so intensely salt as to be a kind of brine, contains in every 100 parts rather more than 24 parts by weight of salts, or about eight times the proportion in ordinary sea-water.

LESSON XXVII.—THE WATERS OF THE LAND.

The Work of Running Water.

1. In the two previous Lessons we have followed the circulation of running water over the surface of the land. That surface we have found to be traversed by an elaborate network of branching water-courses, which, stretching from the slopes of the central mountains down to the sea-shore, carry back to the ocean the surplus drainage of the land. We have seen how in the hollows of the land the water gathers into lakes, yet that it does not accumulate indefinitely there, since it either overflows and again takes the form of streams, or passes off by evaporation. Delivered to the sea, and mingling once more with that great body of water, it is anew raised by the sun's heat into invisible vapour, and carried by winds across the land to begin again the same circulation.

2. So vast a body of water ceaselessly moving over the land cannot but modify the forms of the mountains, hills, valleys, and plains on which it falls and flows. We shall now consider what the nature of such change may be, and how it is effected, beginning with the tiny rain-drop, and tracing the operations of running water down to the mouths of the great rivers as they enter the sea.

3. **Rain.**—The action of rain in washing the air was described in Lesson X., Art. 35. Its further influence in decomposing rocks underneath the surface of the ground was traced in Lesson XXIV. The same kind of change takes place upon the surface of the land. Rain-water, by means of the acids which it takes out of the air, or absorbs from decomposing vegetable matter when it

reaches the land, attacks rocks exposed to the air, dissolving and removing the more soluble parts of them, thereby loosening their cohesion and causing them gradually to crumble down. Calcareous rocks like marble, suffer much from this kind of waste ; but even hard rocks like granite, do not escape. In so far as it depends on the solvent action of the acids upon some of the ingredients of the stone, the process is distinctly a chemical one. But when the outer layer or crust of the rock has in this manner been loosened, heavy rain may wash off the disintegrated particles, and thereby expose a new surface to further decay. The action now becomes mechanical. These two combined kinds of change powerfully influence the scenery of the land (Lesson XXIX.).

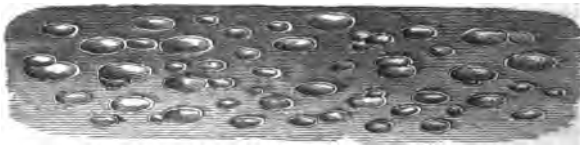


FIG. 58.—Prints made in soft mud or moist sand by rain-drops.

4. The little prints which the rain-drops leave upon a surface of moist clay or sand offer the simplest instance of the mechanical action of rain. From such apparently trivial effects many stages may be traced, until we reach huge pillars, like those shown in Fig. 59, which have been carved out by the blows of innumerable rain-drops. The material of these columns is a stiff earth or clay, stuck full of stones and large blocks of rock, and readily crumbling down under the influence of the weather. The large blocks, remaining of course unwasted, serve each to protect the portion of the earth lying underneath it, while the surrounding clay is washed down. Thus the block

becomes, as it were, the capital of a pillar which seems to rise slowly out of the rest of the earthy mass. Each pillar stands as a monument of the waste which is going

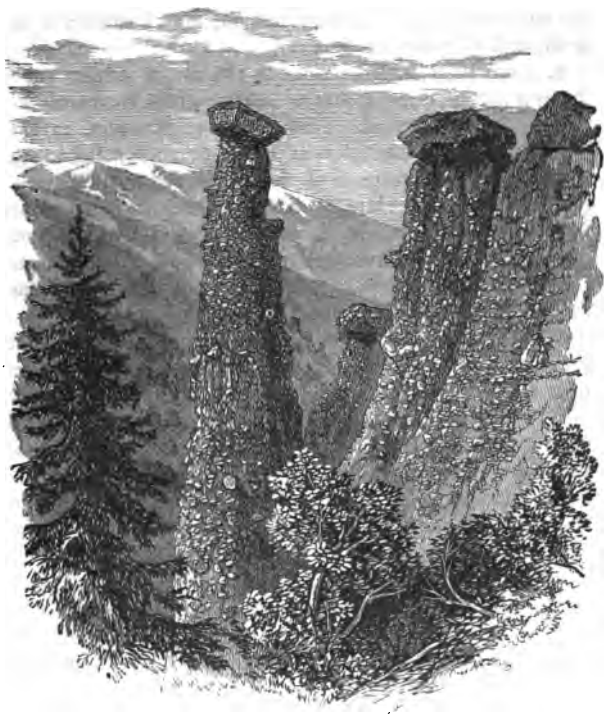


FIG. 59.—Earth-pillars of the Tyrol (from a photograph).

on, somewhat in the same way as the columns of rock or clay left by the workmen in a railway-cutting or a quarry

show the extent of material which has been removed. In such cases, following the rain after it has beaten upon the earth and coursed down the declivities, we find that on reaching the more level ground it has spread out some of the particles which it had washed off, and formed a layer or deposit of them there.

5. Here then may be seen in the action of the rain-drops a sort of type of that of all the great rivers of the globe. It is threefold. First we have Erosion, in the loosening of the particles of earth or rock. Secondly, Transport, in the removal of these loosened particles, and the consequent exposure of fresh surfaces to waste; and thirdly, Deposit, in the laying down of a new stratum formed out of the removed materials. Let us now see how these three kinds of action are manifested over the surface of the land by rivers.

6. **Erosion.**—Every rannel, brook, and river, in short every current of water, no matter how small, which moves over the land is busy removing part of the soil or rock over which it flows. This work, like that of the rain-drops (Art. 3), is twofold. In the first place, the rain and the water of the streams dissolve certain parts of the solid substance of the land and carry them away in chemical solution. In countries where the rocks consist of limestone, or contain a large proportion of the more readily soluble substances, a very considerable amount of waste is carried on by this means, although the water is not visibly affected in colour or transparency. Some idea of the extent of the loss thus sustained by the surface of the land may be formed from the amount of dissolved mineral matter which is found in the water of rivers (Art. 14).

7. In the second place, the solvent action of rain and the disintegrating effects of frost (Lesson XXVII.) cause the surface of exposed rocks to crumble down. Loose clay

and sand are thereby produced, and the hardest rocks are split into fragments. This debris of the land, washed away by the rain and brooks, becomes the instrument of still further destruction. As it is hurried along, its particles, ground against each other, are reduced still further in size, and at the same time are worn smooth and round. They thus acquire that familiar water-worn character which we recognise as the most obvious and distinguishing feature of the detritus in the channel of a stream. So constant is this character, that when, at any point in the course of a river, sharp-edged fragments appear on the banks, we naturally conclude that they cannot have lain very long in the current, nor have travelled very far. The fragments, the longer they are rolled about, and the further they are carried, grow smaller in size, until at the far end of the river they may be found as mere fine sand and mud. If you set out to trace this fine sediment to its source, you would have to seek it among the rocks of the far distant mountains, in the higher part of the river-basin. And you might be able to trace it step by step until you found the huge blocks loosened from the parent cliffs, which in the course of ages have supplied a constant and abundant tribute of detritus to the river. In some respects a river may be compared to an enormous grinding-mill, where large pieces of stone go in at one end, and only fine sand and mud are seen to emerge at the other.

8. But the loose materials swept away by the streams not only wear each other down, they likewise erode the sides and bottom of the water-courses. The water-worn character is thus not confined to the loose sand, gravel, and stones, but is as marked upon the solid rocks over which these materials are driven. Even the hardest kinds of stone cannot resist constant friction. They become smooth and polished, though where out of the reach of the water they may have a rough surface and sharp edges.

Sometimes the upper limit of the grinding action of the flooded stream is in this way well defined along the sides of a rocky ravine.

9. In the course of time the stream grinds out a channel for itself through even the hardest rocks. This channel, however, is seldom a mere deep straight trench. The rocks offering many varying degrees of resistance to erosion, they are worn down unequally, being scooped



FIG. 60.—View of ravines cut by streams out of a table-land.

out where more easily worn away, and projecting where more durable. The stream, thrown from side to side, dashes along, sweeping onward the sand and mud which it drives over its rocky bed, and excavating those winding picturesque ravines which are familiar features in the water-courses of all countries.

10. Along the walls of these ravines when the water is low, curious round cauldron-shaped cavities with smooth

sides may often be observed. These are known as *pot-holes* (Fig. 61), and are formed, as may be observed when the water is in flood, by the grinding action of loose stones, which, caught in eddies of the water, are kept in



FIG. 61.—Cascade and pot-holes of a water-course.

rapid rotation, so as gradually to excavate the holes. Of course the stones are themselves reduced to sand or fine gravel in this process, but other stones are swept in to

supply their place. On the sides of many a narrow gorge through which a stream forces its way, traces of numerous old pot-holes may be seen high above the present water-line. These mark former levels of the stream, and show how in course of time the rocky bed has been gradually dug out.

11. Another way in which a stream erodes its channel is by means of waterfalls (Fig. 62). It is not always



FIG. 62.—Section of a waterfall and ravine.

possible to tell from what cause a particular waterfall may have been formed at first: some original cliff or steep bank, when the stream began to flow, may have determined it. But since that time the waterfall can be shown to have gradually crept up the stream, carving out a ravine between its original point of commencement and the point which it has now reached. This excavation is done by the recoil of the water and spray, whereby the rocks behind the bottom of the fall are loosened. So long

as these crumble down more easily than those at the top of the fall, the cliff over which the water dashes will remain precipitous. Slice after slice being cut off its face, it will shift its place further and further up the stream. The fall will still continue to pour over its edge, and, though slowly moving up the gorge, will preserve from year to year the same general appearance. If, however, from any difference in the nature or position of the rocks, the rate of waste should become more rapid at the top of the cliff than at the bottom, the cliff, instead of overhanging, will then begin to retire at its upper part. As this goes on the waterfall will gradually grow less marked, until it passes into the condition of rapids, that is, a shoot of water over a steep and rough part of the bed. Finally, these rapids will themselves be worn down, and all trace of the original fall will disappear, except the gorge which it excavated during its recession.

12. Almost every large river flowing through a hilly or mountainous region illustrates these features of the erosive action of running water. Perhaps the most stupendous example is that of the Niagara River. The famous Falls of Niagara consist of two vast cascades separated by a small island, and having a united breadth of 950 yards and a height of 140 to 160 feet. It has been computed that 670,000 tons of water are poured every minute over these falls into the foaming torrent below, from which vast clouds of spray rise up into the air. Originally the falls stood at Queenstown, where the limestone forms a cliff above a great plain. Since that period, however, the cataract has slowly receded for about seven miles to its present position. That it is still moving upwards is shown by the large slices of rock which from time to time fall from the cliff over which the water rushes. The rate of retreat has been computed at about one foot in the year. Probably this is an exaggerated estimate; if it be taken

as an average for the past work of the river, somewhere about 35,000 years must have been required to excavate the ravine between the present Niagara Falls and Queenstown.

13. Every stream, then, which drives along sand and gravel on its bottom is busy with the work of erosion. Even in ordinary weather this action may be perceived; how much more stupendous must it be in floods, when ever little runnel is swollen, when earth, sand, and stones are swept by rain off the ground, and when the rivers, rising high above their ordinary level, and acquiring from the increase of their volume augmented velocity of flow, rush over the land and bear their vast burden of detritus down to the sea.

14. **Transport.**—The loose materials acquired in the process of erosion must be removed and disposed of by the streams. As long as the water has velocity enough it keeps the sediment moving, and conveys it sometimes to vast distances. Any check to the velocity causes some of the sediment to fall to the bottom. In considering the nature and amount of work done by rivers in the transport of mineral materials on the land we must bear in mind that these materials exist not only in visible form, such as sand and mud, but invisibly dissolved in the water. So that we should carry away an incomplete notion of the work unless we took this latter part of it into account. We have traced how every spring is busily employed in bringing up to the surface the mineral substances which the water has dissolved out of the underground rocks, and how rain and the water of streams are similarly engaged above ground. Now this dissolved material is conveyed by runnels, brooks, and rivers into the sea. It is not difficult to make a probable estimate of the amount of invisible mineral substance thus carried by a river. First the amount of water discharged by the

river must be ascertained, then the average proportion of mineral ingredients contained in a gallon or other given quantity of the water. The one sum multiplied by the other will of course give us the required result. The celebrated chemist Bischoff calculated that the Rhine carries past Bonn every year enough carbonate of lime chemically dissolved in its water to form three hundred and thirty-two thousand millions of oysters of the usual size, and that if all these oysters could be put together they would form a cube measuring 560 feet in the side. The river Rhone is estimated to carry past Avignon every year 8,290,464 tons of dissolved salts in its water.

15. But by far the largest amount of mineral matter borne by rivers from the land is in the form of mechanical sediment—gravel, sand, and mud. Every river is more or less muddy. After heavy rain even the clearest brook has its water discoloured by the earth it is carrying down. The mere discoloration therefore is a proof of the constant transport of sediment by running water. The amount of material thus transported depends partly, of course, upon the carrying power of the river, which is regulated by its volume and velocity, partly upon the nature of the soil and rocks of its drainage-basin, whether they happen to be earthy and easily worn away or the reverse, partly upon the distribution of the rain-fall, whether it is spread over all the seasons of the year, or crowded into a few weeks or months, so as to produce swollen and muddy rivers while it lasts, and partly, where the river takes its rise from a glacier, upon the quantity of mud which escapes from the melting end of the ice. (See Lesson XXVIII.)

16. A stream having a current of about half a mile in the hour, which is a comparatively feeble flow, can carry along ordinary sandy soil suspended in the water. With a velocity of twelve inches in a second, which is about

two-thirds of a mile in the hour, it can roll along fine gravel, while, when the rate rises to three feet in a second, or a little more than two miles in the hour, it can sweep away slippery angular stones as large as an egg. We can readily understand that in torrents, where the velocity rises so high, as was stated in Lesson XXV., Art. 22, the power of transport must be enormous. Huge masses of rock as large as a house have been known to be moved during heavy floods.

17. It is evident, therefore, that in a rapid river or brook, the mud which discolours its water represents only a part of the sediment which it is carrying down. A great deal of sand and gravel, or even coarse shingle, is at the same time being pushed along the bottom. We may even hear the stones rattling against each other as they are rolled onward by the current.

18. Measurements and estimates have been made of the proportion of sediment in the water of different rivers. This proportion varies of course in different seasons, being greatest during floods and least when the rivers are low. The Ganges during its four months of flood is stated to contain one part of sediment in every 428 parts by weight of water, while the mean average for the year is one part in 510. In the water of the Irrawaddy the proportion was found to be one part in 1,700 by weight of water during floods, and one part in 5,725 during the dry season. In the water of the Mississippi the average proportion was determined to be one part in 1,500 by weight, or one part in every 2,900 by volume of water. The Danube has been found to contain a mean proportion of $\frac{1}{1100}$ by weight of suspended matter, and during extraordinary floods discharges into the Black Sea as much as 2,500,000 tons of silt in twenty-four hours.

19. But in any estimate of the total discharge, the coarse, heavy sediment which is pushed along the bottom

must also be allowed for. In the case of the Mississippi this moving layer has been estimated to deliver into the Gulf of Mexico the vast amount of 750,000,000 cubic feet of earth, sand, and gravel every year.

20. Having ascertained the average quantity of mineral matter suspended in the water or pushed along the bottom, and having estimated the average amount of water carried by a river into the sea, we easily obtain by mere multiplication the total quantity of sediment removed from the land by that river in a year. Thus the Rhone is estimated to carry into the Mediterranean every year rather more than six hundred millions of cubic feet of sediment. The discharge of the Danube into the Black Sea has been determined to be 67,760,000 tons of silt annually. The mean yearly amount of solid matter carried in suspension by the Mississippi into the Gulf of Mexico is estimated about 362,723,000 tons ; and this, including the coarse sand and gravel which are pushed along the bottom, would make a column one mile square and 268 feet high. We may form some notion of this amount of material by supposing that 1,100 merchantmen, each laden with 1,000 tons of it, were to arrive every day for a whole year at the mouth of the Mississippi and discharge their cargoes into the sea. They would only equal as carriers the work of this single river.

21. But many rivers greatly exceed the Mississippi in the proportion of solid matter which they transport. During the rainy season in India the streams become torrents of mud. Dr. Livingstone in his African travels came upon sand-rivers—currents of sand moving along with a comparatively small amount of water. In trying to ford them he felt thousands of particles of sand and pebbles of gravel striking against his legs, even in dry weather, and he saw that after the rains the quantity of detritus removed by these streams must be enormous.

22. The amount of sediment carried down by a river to the sea in a year represents the extent of loss which the region drained by the river has sustained within that time. Knowing the quantity of sediment and the area of country from which it has been derived, we can determine the amount by which the general surface of the river-basin has been lowered. Thus at its present rate of work the Mississippi reduces the general level of its drainage area $\frac{1}{6000}$ of a foot annually, or one foot in 6,000 years. Could this rate of denudation be continually kept up over the surface of North America, which is computed to have an average height of 748 feet, the whole of this half of the continent would be reduced to the sea level in about 4,500,000 years. Such calculations are of importance in showing that the present surface of the land must be continually changing, and cannot, therefore, be comparatively of high antiquity.

23. **Deposit.**—All rivers, then, are constantly busy grinding down and transporting gravel, sand, and mud over the surface of the land. To ascertain what becomes of all this material, let us follow the course of a river from the mountains to the sea, and watch how the sediment is disposed of by the way.

24. When running water has its velocity checked it loses some of its power to transport sediment, which then partly sinks to the bottom. This may happen when a stream enters upon a gentle slope or plain, where it must move more slowly, or when it joins a larger and more gently-flowing current, or when it falls into still water, like that of a lake, or into the sea. Hence, during its course as well as at its termination, a river necessarily encounters many obstacles to its progress, by which it is compelled to slacken its pace and to drop some of its sediment.

25. Beginning among the mountains, we meet with abundant examples of this arrest of the detritus which

the torrents have been sweeping down the declivities. A steep slope may be deeply trenched with gullies which have been cut out of the soil or rock by the torrents. But wherever these ravines reach a strip of more level ground they each show a pile of rubbish, which the headlong brooks, checked in their flow by the change in their angle of descent, have been forced to throw down. In a long, narrow valley among lofty mountains, so numerous may be the torrents, and so vast the heaps of gravel and stones which they tear out of the hill-sides and heap upon the bottom of the slopes, that it may be difficult to maintain a roadway, which is liable at one point to be buried under huge masses of debris, and at another to be swept away by a flooded stream.

26. But only the larger and coarser kind of sediment is usually arrested at the foot of these mountain slopes. The finer parts and even some of the rougher gravel are carried further into the valley, where the various brooks unite into one main stream. At every point where this stream is checked, an accompanying deposit of sediment occurs. Among its many serpentine windings the current, though rushing briskly round the sharp side of a bend, is thrown into an eddy or slack water at the concave side, and there you find a bank of sand or shingle.

27. When the stream is flooded it not only fills its ordinary channel, but rises, and overflows the flat meadows on either side. These tracts of level ground diminish the velocity of the overflowing water, which must therefore allow some of its sand and mud to fall upon them. Should they be covered with herbage or brushwood, the leaves and branches act as filters to retain the sediment. So that when the stream has subsided the inundated ground is found to have received a coating of fine silt, or even, it may be, of coarse gravel.

28. If, then, each flood adds to the height of the level

ground on either side of the stream in a valley, the time must come when, even at the highest flood, the water will not be able to overspread the plain. This would arise not only from the heightening of the plain by such frequent deposits of sediment, but from the gradual deepening of the stream-channel by the scour of the current. As the channel is deepened the current will continue as now to eat into its banks, curving from side to side, and forming a new plain at a lower level. This process has been in progress for a long time in most river-valleys. A succession of terraces marking former levels of the river-floods may be seen rising even up to heights of several hundred feet above the present river. The accompanying figure (Fig. 63) represents a section drawn through one of

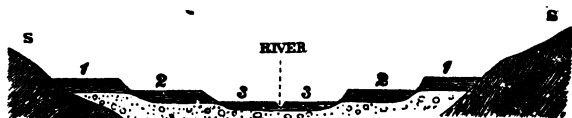


FIG. 63.—Terraces of gravel, sand, and mud, left by a river.

these river-valleys, and shows the relation which the low-level terrace (3) at present in course of formation bears to the higher and of course older ones. In some of the latter, remains of primitive man, such as chipped stone spear-heads and other implements, have been met with in different parts of Europe, showing that when the rivers flowed at the level of these higher terraces, a rude human population already existed in countries where certainly the river-valleys have undergone hardly any appreciable change within historic times.

29. While still in the mountainous or hilly part of its course, a river may have to traverse one or more lakes. Each of these sheets of still water arrests the current and

compels it to drop its burden of sediment. As we saw in Lesson XXVI., lakes filter the river-water, which, leaving on their floors its sand and mud, issues at the lower end quite clear. It is where the river enters the lake that the chief deposit takes place. By degrees that portion of the lake is filled up and converted into a plain which, after being gradually heightened by the river-floods, at last comes to be above the flood-limit. In this way the upper end of the Lake of Geneva has been so diminished by the deposits of the Rhone that a Roman port, still called Port Valais, is now nearly two miles from the edge of the lake, the intervening space consisting of meadows and marshes.

30. If no other agency should intervene to keep them open, lakes must evidently in course of time be filled up by the earth and sand washed into them. This has already been the fate of many. The once united lakes of Thun and Brienz in Switzerland have been separated by the tract of land formed by the streams at Interlaken. In Great Britain, as well as generally throughout Northern Europe, every stage in the disappearance of lakes is abundantly shown, from the mere tongue of sand encroaching upon the edge of a deep mountain tarn up to the flat moss-or fertile meadow which marks where a former lake has been silted up.

31. When the river enters the lowlands the diminished slope of its course causes it to flow with a weaker current, and therefore to drop some of its burden of mud and sand as it moves along. Winding to and fro, it cuts down its banks at one part and heaps up sediment at another, so that in course of time the whole of a wide plain comes piece by piece to be levelled by the shifting stream. The plain is formed indeed out of the sediment which the water has carried down from the higher ground. A well or pit sunk anywhere over its surface shows that beneath the soil

there lie layers of water-worn silt, sand, or gravel, like the materials now being transported and deposited by the river.

32. When the plain is long and the seaward inclination slight, the flow of the river may be so lazy that instead of scouring out its channel it may be unable to prevent the sediment from sinking to the bottom and actually heightening the bed of the channel. During floods, too, the chief deposit of silt takes place on the banks on either side of the river; consequently, as the ground there is raised in level more than the plain beyond, which receives less sediment, the river gradually comes to flow at a higher level between broad embankments of its own building. From time to time it breaks through the lower or weaker parts of those embankments and inundates the plain, perhaps even scouring out here and there a new bed. In cultivated regions, like those watered by the courses of the Po and Adige, in the plain of Lombardy, and by the lower portion of the Mississippi, much care is needed to strengthen the banks, with the view of averting floods. The rivers have sometimes so heightened their channels that the surface of the current flows during floods at a higher level than the streets of towns on their banks.

33. As a familiar example of such broad plains formed by the accumulation of sediment brought from the interior of the land and laid down by the long-continued flow and overflow of rivers, we may take the valley of Egypt, which is a river-formed plain laid down between two ranges of higher ground. In Lower Egypt at midsummer every year the Nile begins to rise and cover the flat land on either side of its course. The inundation is at its height in about three months, and then the waters, after remaining stationary for nearly a fortnight, retire to their former level. This annual rise corresponds (as mentioned in Lesson XXV., Art. 18) to the rainy season in

the mountainous table-land of Abyssinia. When the monsoon blows from the Indian Ocean it brings torrents of rain, which rush down the rugged slopes of that country and sweep along a vast amount of mud. Thousands of swollen and muddy streams are at last united in the Blue Nile, which carries this vast body of discoloured water down into Egypt. After the inundation has ceased the ground is found to be covered with a thin coating of rich fertile mud left by the water. It has been estimated that the thickness of this annual deposit does not exceed that of a thin sheet of pasteboard, so that a depth of only two or three feet of the soil of Lower Egypt represents the continuous deposits of a thousand years. The increase of the Egyptian plain evidently takes place at the expense of the high grounds of Abyssinia. It is the finer particles, worn away from the rocks of these uplands, which form the mud spread over Egypt. Here we see how the erosion, transport, and deposit of materials by running water combine to build up a plain and renew its fertility.

34. The great plains of India furnish likewise admirable illustrations of the way in which the debris of the mountains is spread out by rivers on the low grounds. The valleys of the Indus, Ganges, and Brahmaputra have been filled up by the accumulations carried down from the Himalaya chain by these great rivers. The Tigris and Euphrates have likewise combined to fill up the upper half of the valley of which the Persian Gulf is the still remaining lower half. On the American continent this process is exhibited on the most stupendous scale. Most of the eastern coast-line of the United States is a plain formed by the deposit of sediment washed off the land. The valley of the Amazon, with its vast forests or *silvas*, forms so long and so level a plain that ships can sail up the river to the very foot of the Andes, a distance of 2,000 miles inland from the sea.

35. We have now, lastly, to see how the sediment borne down by rivers is disposed of when it reaches the sea. Many rivers have at their mouths what is called a bar, that is a ridge of gravel or sand stretching across the channel, and always under water. From what has been said already in this Lesson, the origin of this bar will now be understood. It is evidently due mainly to the arrest of the sediment where the current of the river is checked by coming in contact with the sea. The coarser gravel and sand which have been pushed along the bottom of its channel by the river now encounter the opposition of the salt-water, over which the lighter river flows onward. The sea too piles up fresh materials upon the bar from the outside, while, on the other hand, the river when flooded drives its bar further seaward. Hence this barrier to navigation at the mouth of a river is continually shifting its position and altering its shape and size, according as the action of the river or that of the sea predominates.

36. Some of the larger rivers of the globe carry out at their mouths on a vast scale the same kind of action which was described as going on where streams enter lakes. The sediment deposited in the sea has gradually increased so much as to fill up the bay or gulf into which the river originally flowed, and to convert it into a territory of flat land. This advancing tract of river-formed ground is usually triangular in form, the apex pointing up the river. It was this resemblance to the Greek letter Δ that suggested the name of delta, as applied to these accumulations at the mouths of rivers. Down to the head of its delta a river usually is augmented by tributaries from either side, and does not branch, except for a limited space, as where it encircles an island in its course. Its current thus becomes more and more ample. But when it reaches the delta it begins to subdivide, and continues to branch out until in some cases the flat plains and marshes

are traversed by innumerable tortuous channels of water. By this ramification over the flat ground the deposit of sediment is facilitated, consequently the channels are being continually filled up, while new ones are cut through the soft earth and silt. Two or more main branches of the river carry out the chief part of the water and sediment to sea, and at their mouths the increase of the delta is most apparent.

37. The accompanying figure (Fig. 64) shows the shape of the lower part of the delta of the Mississippi, and the way in which the tongues of low muddy swamps are pushed outwards into the Gulf of Mexico. The average rate of advance of the delta has been estimated to amount to 86 yards in the year. The Tiber throws forward its delta at the annual rate of about 12 to 13 feet. The delta of the Po has increased at such a rate that the port of Adria, which stands on it and was so important in Greek and Roman times as to give its name to the Adriatic sea, is now 14 miles inland.

38. Since this action must have been in progress for a vast number of centuries, we may be prepared to find that some deltas are of enormous size. That of the Mississippi embraces an area of about 40,000 square miles. That of the Ganges and Brahmaputra is as large as the whole of England and Wales. Their vast superficial extent indicates of course the high antiquity of these deltas. But we must remember also that as they have been formed by the gradual filling up of gulfs of the sea, where these gulfs were deep a prolonged period of time would be occupied in bringing up the level of the bottom to that of the surface before any advance of the plain of the delta could take place. Where the tides and currents of the sea interfered to sweep away some of the sediment, the time required would be still further increased. The delta of the Ganges has been bored into at Calcutta, and its

deposits, consisting of sand, gravel, silt, and layers of vegetation, were found to be more than 400 feet thick.

39. In cases where the form of the coast is unfavourable to the formation of deposits, or where the rivers flow with sufficient velocity to carry out their mud to sea, or where marine currents sweep past the river-mouths and



FIG. 64.—Delta of the Mississippi.

carry away the sediment, deltas do not occur. The Amazon and La Plata, for example, form no deltas. But they nevertheless pour a vast quantity of fine silt into the Atlantic. Even for a distance of 300 miles from the mouth of the Amazon the sea is said to be perceptibly discoloured by the mud of that river.

LESSON XXVIII.—THE FROZEN WATERS OF THE LAND.

Frost, Snow-fields, Glaciers.

1. In Lesson X. we traced how the moisture of the air, when the temperature sinks to the freezing-point, passes into ice. We have now to see what becomes of this ice when it falls from the air to the land, and how the waters of the land are affected when they too are cooled down below the freezing-point. Changed into solid ice, they cannot but acquire new and different powers from those which have been the subject of the last four Lessons. What part then does the frozen water of the land play in the general plan of the globe? We shall find some answer to this question by considering three aspects of the ice of the land—Frost, Snow-fields, and Glaciers.

2. **Frost.**—Most substances suffer contraction from cold, and consequently increase in density. A cubic foot of pure water, for example, at a temperature of 40° Fahr. weighs more than the same quantity at 60° : or, in other words, a vessel exactly filled with water at 60° would be found to be not quite full were the water cooled down to 40° .

3. But a remarkable change takes place during the further cooling of the water. When the temperature reaches 39.1° Fahr., contraction ceases. This is, therefore, the temperature at which pure water is heaviest, and is called the *point of maximum density* of fresh-water. Below this point the water expands, and instead of sinking, remains on the surface of the warmer layers, where it is at last converted into ice when the temperature has fallen to 32° . Hence it is that ice is formed

on the surface of lakes and rivers, where, being lighter than water, it floats. As the cold continues the first thin crust of ice is increased by additions to its under surface, until after a severe winter it may be a foot or two in thickness. In this way, after a long and hard frost, rivers and lakes become so firmly frozen over that carts and heavy waggons may be drawn over them, though the water with its living inhabitants remains still liquid underneath. Thus canals and rivers which for most part of the year are highways for boats and barges, become in these cold winters thoroughfares for carriages and sledges, as well as passengers on skates or on foot.

4. The parts of the earth's surface where this curious transformation may be ordinarily seen in winter are indicated on Plates V. and VI. by the position of the isotherm of 32° . In the northern hemisphere, over all the ground lying to the north of that line, and in the southern hemisphere, on all the ground lying to the south of it, the waters are frozen during winter. But even in much more temperate latitudes the cold is occasionally severe enough to make the rivers and canals passable on foot, and what is more remarkable, actually to encrust the sea with ice. Thus in the year 401, and again in 642, the winter was so intensely cold in Southern Europe that the Black Sea was entirely frozen over; in 850 the Adriatic was covered with ice; while in 1233 and in 1314 the rivers in northern Italy were likewise frozen. In the year 1205 a severe frost, lasting for three months, bound up the soil of England so that it could not be ploughed, and froze the canals, rivers, and ponds. The last time that the Thames at London was passable on ice was in the year 1814. Thousands of passengers crossed the river on foot between London and Blackfriars bridges, and numerous wooden booths and tents were erected on the ice, together with merry-

go-rounds and other amusements, so that the busy concourse of people was called "Frost Fair."

5. The freezing of sheets of water on the land may take place so equably, and the ice may disappear again so quietly, that after the frost has passed away no trace may be left of its having occurred at all. But should a frozen lake break up under a storm of wind, large masses of ice will be driven ashore and may push up the sand, gravel, and stones lying at the water's edge. Such heaps of ice often take some weeks to thaw, and after they have melted, the stones and sand which they had driven ashore are found scattered over the ground. This can be seen on a great scale on the shores of the large Canadian lakes in spring, where also another result of the severe frosts may be observed. Blocks of stone lying in shallow water are frozen in the ice. When a thaw comes and the ice breaks up, they are actually lifted from their places by the fringes of ice which have formed round them, and are floated away to some other part of the shore, or may even be driven out into the deeper water, so as at last to be dropped there. This transport of stones may be watched along some of the rivers as well as the lakes of Canada. It has been particularly observed on the shores of the St. Lawrence.

6. But the effects of the fall of temperature in winter are not confined to large sheets of water on the land, and indeed it is not there that the most remarkable results of frost are to be observed. We have seen how the rain which falls upon the land soaks through the soil and rocks. Both soil and rocks therefore contain abundant water in their pores and cavities. When severe frost sets in, this water freezes. The soil becomes consequently hard, so that even muddy places where we should have sunk deep in mire can now be safely and easily traversed on foot.

7. Let us consider the consequences of the freezing of the soil. When a jug of water is exposed to severe frost it is apt to be split. And often in insufficiently warmed dwelling-houses water-pipes in the same way burst during frost. The reason of these mishaps is to be sought in the expansion of water in the act of freezing. When passing from the liquid to the solid state, water undergoes a sudden and remarkable expansion, amounting to about one-tenth of its volume. At the moment of change it exerts great force upon the sides of the vessel or cavity containing it, and if it cannot readily escape it will do its best to burst these sides, that it may find room for its increase in bulk. Now what takes place in a water-jug or pipe goes on also among the little particles of water inclosed among the grains of the soil and the rocks. The frost expands them, and they push these grains aside. On a winter morning after a night of sharp frost we may notice that the expansion has been great enough to force the little pebbles and stones of a roadway out of their places. So also, in countries like Canada, where the winters are extremely cold, wooden fences are in the course of a year or two twisted out of the ground.

8. It is not until after thaw that we are made fully aware of what the frost has done. So long as the cold continues the separated particles of soil are kept together by the ice, which binds them into a hard solid mass. But when this ice melts, the grains of sand and earth become loosened from each other. Walking on a road or over a ploughed field after frost, we find that this loosening has been carried so far that the ground has become coated with mud. Indeed the millions of little ice crystals, which the frost wedges in among the grains of the soil, have much the same effect as if the earth were ground down in a grinding-mill or mortar. They pulverise

it, that is, reduce it to powder, and thus lay it more open to the branching roots of plants which obtain so much of their sustenance from the soil. Farmers are in the habit of ploughing their land before the cold weather sets in, and leaving the upturned loam exposed to the beneficial effects of the succeeding frosts.


9. The powerful mechanical effects of frost are so well seen in soil owing to the abundant moisture retained among the particles of which the soil is composed. But any porous rock which contains sufficient water, and is exposed to a great enough cold, may show the same kind of disintegration. Hence in countries where the winters are severe, ordinary building-stones and mortar are found to peel off in successive crusts, or to crumble down into powder after the frosts have given way to milder weather. Even in the comparatively mild winters of Britain this may be constantly seen; in the severer climate of North America it is a serious and costly evil, since it prevents the use of many kinds of building-stone, which in the absence of frost would be very valuable.

10. Passing now from the water retained and frozen among the pores of the soil and of rocks, let us consider further the result of the formation of ice in the larger cracks and crevices of cliffs and crags. To realise properly how abundant these natural lines of division are in all rocks, look at any exposed face of rock such as a sea-cliff, the ravines of a river, or the precipitous sides of a mountain. No matter what may be the kind of rock, you will find that it is traversed by many parallel or intersecting divisional planes or "joints." We saw in Lesson XXIV. that these lines are made use of as channels by which the surface water descends underground, and by which it re-ascends to form springs. When a frost comes, severe enough to penetrate beyond the mere outer layers of rock, the water contained in the external parts of these

joints freezes. It often happens that this takes place in cavities where there is little room for expansion, and where therefore the ice exerts its great force in pushing the walls of the joints asunder. Winter after winter the process is repeated, until at last the portion of rock on the outside gets so far wedged off from the rest that it loses its balance and falls with a crash to the bottom of the cliff. In all countries subject to intense frosts the bases of cliffs and crags may be seen to be strewn with large rough blocks, which have in this manner been loosened from their places above. Among the valleys of mountains which rise high up above the snow-line the operations of frost are powerful in splintering the crags and pinnacles, and giving them the sharp spiry forms they so often assume. (See Lesson XXIX.)

11. Snowfields.—Wherever the land rises above the snow-line it is buried under a permanent sheet of snow, from which only the higher and steeper mountain peaks project. In some regions, as in the table-land of Norway, the broad and tolerably level surface of the ground allows the snowy covering to spread in a vast undulating sheet three or four thousand feet above the level of the valleys. Standing upon one of the heights at the edge of such an expanse of snow, you see what looks like a white frozen plain, with no limit save the line where it meets the sky. In other parts of the globe, where rugged groups of mountains tower above the snow-line, and there are consequently no such level tracts, the snow accumulates in the hollows and on the higher slopes. The semicircular ranges of mountain cliffs and crests often inclose vast basin-shaped depressions, each of which becomes a gathering place for the snow. To all the permanent sheets of snow, whether occurring on table-lands or on mountains, the general term of snow-fields is given.

12. In these regions, since the moisture falls from the



air as snow rather than rain, and the heat of summer is insufficient to melt it all, the quantity of snow would increase indefinitely were there no provision whereby the surplus could be removed. The thickness of snow on some snow-fields must amount to many hundred feet. Greenland, for example, is almost wholly buried under one vast snow-field, so deep as to cover over the inequalities of the surface of the land almost as completely as the furrows of a ploughed field are lost beneath a heavy snow-wreath.

13. There are two ways, besides melting and evaporation, in which the snow-fields get rid of their excess of snow ; these are avalanches and glaciers. Where the edges of a snow-field overhang steep slopes, portions of the more or less consolidated snow from time to time break off and rush with a noise like thunder and a terrific force into the valleys, tearing up the soil, sweeping down loose blocks of stone, uprooting or breaking trees, and carrying destruction as far as they reach. These snow-falls are called avalanches. In Alpine countries forests which lie in the pathway of such descending masses of snow are carefully preserved as barriers to protect the meadows and villages from ruin. Roads which pass along the base of snowy mountains require in some places to be covered over with a strong archway of masonry to protect them from the disasters caused by frequent snow-falls, which would not only sweep away every traveller and carriage in their course, but would even tear up the road itself, or bury it under heaps of earth and stones.

14. **Glaciers**, however, are the chief means whereby the superabundant snow above the snow-line is removed. While the snow forming the surface of a snow-field is loose and open in texture, like that which covers the ground in winter and disappears in spring, the under portions become more and more close and firm under

the increasing pressure of the overlying mass. This compacted snow (*névé* or *firn*, as it is called in Switzerland) passes by degrees into clear blue ice, as the imprisoned air is more and more completely squeezed out of it. If the snow-fields lay on perfectly level table-lands, there would be no general movement of the snow except along the edges of the plateaux, where the gathering snow-sheet would break and slide off in fragments into the valleys. But since the ground usually has a marked inclination



FIG. 65.— Snow-field and glaciers of Holands Fjord, Arctic Norway.

away from the axis or water-shed, the snow, by virtue of the action of gravity, is forced to slide downward even upon a very gentle slope. It is during this movement that the air is chiefly pressed out, whereby the loose white, opaque snow is converted into solid, blue, transparent ice. Having acquired this slow-sliding motion, the mass of snow must needs seek the lowest levels. It therefore moves downward into the heads of the valleys, which ascend into the snow-fields. Each of these hollows

becomes a receptacle in which the snow, pressing downwards from each side as well as from behind, accumulates to a great depth, and is so jammed up between the mountain slopes as to take the form of solid ice from top to bottom. Driven onwards by the pressure of the advancing mass behind and by its own gravity, this ice fills up each valley sometimes to a depth of several hundred feet, and for a distance of many miles. Such a tongue of ice, proceeding from a snow-field above and protruding below the snow-line down a valley, is termed a glacier. In the accompanying drawing (Fig. 65) two glaciers are seen descending from one of the great snow-fields of Arctic Norway. In one case the ice almost reaches the sea; in the other the glacier is smaller, and does not proceed so far from its parent snow.

15. A glacier, then, represents the escaping drainage of the snow-fall above the limit of perpetual snow, as a river does the surplus drainage from the rain-fall on lower grounds. Its size must depend upon that of the valley and the extent and declivity of the snowy gathering ground from which it issues. In the Alps, for example, the Great Aletsch Glacier extends about fifteen miles down its valley. Sometimes a little glacier fills up a high recess on the flank of a mountain, and never reaches even to the nearest valley. In other cases the glacier descends far below the snow-line, even into the region of meadows, corn-fields, and gardens. The lower glacier of Grindelwald, in the Bernese Oberland, reaches to a point about 3,500 feet below the snow-line on the north side of the Alps.

16. To realise as clearly as possible the general appearance presented by a glacier, let us suppose ourselves placed at the lower end of one of those among the Alps. On either side of the valley the slopes are clothed with pine. Patches of green pasture catch the sunlight in the

hollows and on the lower projecting hills ; while around us lie scattered cottages and bright meadows. In front stands the abrupt end of the glacier—a steep, but broken slope of ice, from the base of which issues a river of pale muddy water. Numerous large blocks of rock are scattered about on the valley-bottom below the ice. The ground there, indeed, is mainly composed of coarse shingle, like



FIG. 66.—View of a glacier, with its lines of rubbish (moraines) and the river which escapes from its end. Ice-worn hummock of rock and transported stones are shown in the foreground.

that which forms the bed of the present river. Even on the ice itself we may see heaps of stones, some, perhaps, poised just on the verge of the last steep slope of the glacier, whence they must soon roll down to join the crowd of others which have preceded them. Looking into one of the deeper rents in the ice, we see it to

be of wonderful purity, and of the most exquisite blue colour. And yet most of its outer surface is so obscured by earth and stones, that at first, perhaps, we can hardly be persuaded that this same clear, transparent ice really lies below.

17. We ascend to the surface of the glacier either by mounting among the broken cliffs of ice forming its abrupt front, or by choosing a safer and easier path up the slope on one side of the valley. The ice is now seen to lie as a great sheet filling the bottom of the valley from side to side, and stretching far up into the heart of the mountains. Its surface has at first a gentle slope, and is comparatively smooth, though many glaciers even at the lower end present a marvellously rugged aspect, like that of a tempestuous sea suddenly frozen. The ice is much obscured in places by earth, gravel, and stones which cover it. As we ascend the valley we notice that these surface accumulations are especially abundant along the sides of the glacier, and likewise in one or more ridges along its centre. During the day, when the sun shines out warmly, the surface of the ice is thawed, and consequently abundant little runnels of water flow over it. At night, when these are frozen, the glacier becomes once more silent.

18. In several respects a glacier resembles a river; and this resemblance is closer than might at first appear. If the position of any prominent object on the surface of the ice be observed with relation to some fixed point on the bank, or if, as in the original survey of the Mer-de-Glace by Professor James D. Forbes, a transverse series of stakes be driven into the ice across the breadth of the glacier, and their position observed at intervals from the bank, the ice is found to be slowly moving down the valley, and faster in the centre than at the sides, because the friction of the sides impedes the flow. The average

rate of movement on the Mer-de-Glace was thus ascertained to be, during summer and autumn, from 20 to 27 inches in the twenty-four hours at the centre, and from 13 to 19½ inches at the side.

19. Bending from side to side, and travelling over an uneven floor, the glacier is split across by rents called *crevasses*, which often open and form wide and deep chasms, extending down to the very bottom of the glacier. Stones and earth lying on the surface of the ice frequently fall into these fissures, and thus reach the floor of the valley below the glacier, or are imprisoned in the ice when the sides of the yawning crevasses are pressed together again as the glacier moves onward.

20. When a river reaches a steep and rocky part of its bed, it forms a rapid ; when it comes to a cliff, it takes the shape of a water-fall. Ice not having the same mobility as running water, cannot so easily adapt itself to the irregularities of its channel. But it shows these irregularities in a very marked way. In the course of the glacier which we have supposed ourselves to be visiting, after, perhaps, some miles of slow ascent, we come to a precipitous slope, where the ice, completely shattered by innumerable fissures, rises into pinnacles and sharp crests of many fantastic forms. Could we watch that tumultuous descent of broken ice for a long enough time, we should find it to be all in slow motion downwards. It is an ice-fall, and answers in the mechanism of a glacier to the water-fall in that of a river. Underneath it the bottom of the valley is steep or precipitous, and the ice, unable to descend the declivity in one unbroken sheet, is cracked and splintered in this wonderful way. But just as a river, no matter how much it may have been tossed into foam by its descent in a water-fall, speedily takes its usual shape and rolls onward as if no sudden plunge had so recently disturbed its current, so a glacier, though it

may have been reduced, as it were, to fragments at one of these ice-falls, soon reunites at the bottom, and again pursues its course as a solid and continuous sheet of ice. If the glacier is a large one it may receive tributary glaciers from valleys on either side. The Mer-de-Glace of Chamouni, for example, is formed by the united mass of several glaciers, as shown in Fig. 67. Each of these, as well as the main stream, may be traced upward until it is found insensibly merging into the snows which fill up the hollows in the higher part of the mountains.

21. One cannot trace the course of a glacier and realise on the ground the size and character of those vast tongues of ice which carry off the drainage of the snow-fields, without wishing to know more of the work done by them. The rivers which bear away the surplus water of the land are busy in a stupendous work,—wearing down the mountains and valleys, and strewing their debris over the plains, or sweeping it out to sea (Lesson XXVII.). Glaciers, too, are engaged in a similar task. They transport the waste of the mountains down to lower levels, and they evade the sides and bottoms of the valleys in which they move.

22. Transport.—We have already noticed that a glacier is largely covered with earth and stones. Whence come these materials which so darken and obscure the surface of the ice, and which, far below the snow-line, where the glacier melts almost at the edge of the meadows and gardens, lie piled in heaps on every side? From a good point of observation above the glacier you may notice that the heaps are not scattered wholly at random across the surface of the ice; but that there are long lines of stones which keep apart from, but parallel to, each other, and run along the length of the glacier. You see them winding to and fro with the varying curves of the glacier in its course until they are lost in the distance.

In the drawing of a glacier in Fig. 66, for example, some of these lines of stones are seen both in the centre and



FIG. 67.—Plan of the Mer-de-Glace of Chamouni and its tributary glaciers, showing the way in which lateral moraines became medial.

towards either side. Follow the central line on a glacier (sometimes there are several lines down the middle), and

you will find at last that at some higher part of the valley it brings you to a point where two branches of the glacier join, and where the line of stones either continues up the side of one of the branches, or divides into two, one portion keeping to the right-hand side of one of the tributary glaciers, the other remaining on the left-hand side of the other branch (Fig. 67). In every case you will observe, if you trace the glacier far enough up its valley, that a line of stones in the middle of the ice comes really from the side, and is due to the confluence of two branches of the glacier. Owing to the irregularities in the slope and breadth of the glacier-channel, and the manner in which the ice is consequently driven to adapt itself to these, as well as to the melting of the surface of the glacier (Art. 25) the lines of stones are apt to lose their distinctness as they advance down the valley, until, after some great ice-falls and abundant crevasses, the rubbish comes to be spread more or less over the whole of the glacier's surface, as it is towards the lower end of the Mer-de-Glace (Fig. 67).

23. These heaps of stones, earth, and gravel lying on the ice are known by the name of moraines. When on the side of the glacier, they are termed *lateral* moraines; on the centre, they are called *medial*; at the end, where the glacier throws down its burden as the ice melts, they are spoken of as *terminal*.

24. In all cases, then, these moraine-heaps can be followed up the glacier to the base of some cliff or craggy mountain-slope, whence the blocks of rock have been derived. In such positions you see how the fragments have been loosened by the severe and prolonged frosts of these high grounds, until, wedged off from the face of the cliffs, they have rolled down to find a resting-place upon the glacier below. Once on the ice, they are slowly borne down the valley and dropped among the heaps of

rubbish at the far end of the glacier. Much of the rubbish, however, falls down the numerous rents or crevasses in the ice, and reaches the bottom of the glacier, there to aid the ice in accomplishing another important part of its work (Art. 28).

25. It is curious to watch the progress of one of the large blocks as it travels with the ice. We thereby learn, what might not otherwise be so evident, that the surface of the glacier is continually being lowered by melting and evaporation. Take the case of a large flat stone, which, bounding from some high crag, has found a lodgment upon the ice. The portion of the ice lying below



FIG. 68.—Glacier table—a pillar of ice supporting a block of stone.

the stone is screened from loss by thawing and evaporation, but the surrounding parts of the glacier, not so protected, are insensibly wasted away. Consequently the stone begins, as it were, to rise out of the glacier. Its pedestal of ice continues to increase in height, but being exposed on the sides to sun and air, is lessened in diameter, until it becomes too slim to support the heavy burden of stone, which then tumbles down upon the surface of the glacier (Fig. 68). But as the general waste of the surface of the ice continues, the new position of the stone

is soon marked by the rise of a new pillar of ice as before. The same block of rock may thus in the course of its journey down the glacier become the capital of several successive ice-columns. The same kind of testimony to the remarkable lowering of the surface of the glacier is shown by the long parallel moraine mounds. Looking at one of these ridges, and even climbing and standing on it, you would suppose it to consist of fragments of stone throughout. But pull down some of the loose blocks, and



FIG. 69.—The Pierre-à-Bot, near Neufchâtel (J. D. Forbes).

you will discover that the solid ice lies immediately below. It is in fact a ridge of ice with a coating of debris which has protected it from the general waste, as the detached stones screen the ice-pedestals beneath them.

26. In most valleys with glaciers in them large blocks of rock may be observed above the present limits of the ice, but in places to which at one time the ice evidently reached. They occur poised sometimes in the most

precarious positions, as if a man's strength would be sufficient to dislodge and send them down the slope. These perched blocks, as they are called, furnish good proof of the former extent of the glaciers. By their means, for instance, combined with other evidence, it can be proved that the glaciers of the Alps at one period filled up the Swiss valleys, and even spread over the broad plain of Switzerland between the Lake of Geneva and the Jura. On these vast ice-rivers blocks of granite from the Mont Blanc group of mountains were transported across what is now the valley of the Rhone and the Lake of Geneva, and were stranded high on the sides of the Jura range. The accompanying figure (Fig. 69) by the late Principal Forbes, representing one of these travelled blocks, erratics, or, as the Swiss call them, foundlings, shows the great size which some of them attain.

27. We conclude then that one important part of the work of glaciers is the transport of the materials of the mountains from higher to lower levels. In the case of such a mountain chain as the Alps or the Himalaya the glaciers melt long before they can reach the outskirts of the high grounds. They, therefore, do not bear their burden beyond the mountain region, though, as we have just seen, they once carried it much farther than they do now. But in Arctic and Antarctic regions the glaciers actually reach the sea-level, and even pushing their way out to sea, break off into icebergs, as already described (Lesson XVI.). The accompanying drawing (Fig. 70) represents the little glacier at the head of the Jokuls Fjord in the north of Norway, which descends into the sea. You observe a few transverse crevasses in it near the base. From time to time the outer portions break away and float slowly on the current of fresh-water which is moving down the inlet. Pieces of stone may now and then be seen upon these little bergs. In Green-

land vast glaciers, like the Humboldt glacier, which measures 60 miles in width, descend to the sea-level, and extend for some distance from the shore till their seaward ends split off and become icebergs. (See Fig. 16.) Occasional blocks of stone have been noticed on these bergs. The debris of the mountains is thus actually borne by the ice away out to sea, and may travel for



FIG. 70.—Glacier descending to the sea. Head of Jokuls Fjord, Arctic Norway.

many hundreds of miles before it sinks at last to the bottom of the ocean.

28. Erosion.—In the second place, a little further observation will suffice to show that the carrying of materials down its valley is not all the work which the glacier performs. Let us consider the river of muddy water which escapes from the end of the ice. Muddy at every season of the year, it is most copious and

discoloured during the warm dry weather of summer and autumn. The mud cannot come from the melting of the glacier itself, for the ice is clear and pure. It is not derived from the bright little brooks and torrents which come down from the melting snows and the springs on either side of the valley. Yet it undoubtedly comes out from beneath the glacier. Mud consists merely of the finer particles worn from rocks. There must, therefore, be a great deal of waste somewhere to account for this constant and plentiful supply of mud.

29. Following this inquiry further, we observe that the rocks which rise on either side of the valley from under the glacier are remarkably smoothed. Underneath the ice, as we may sometimes detect, the floor of the valley is similarly smoothed and polished. The contrast between the rounded and smoothed outlines of the knobs and hummocks of rock near the ice, and the sharp, rugged forms of the crags above, is often singularly well-defined. Besides this smoothing and polishing, the surface of the rocks may be observed to be covered with many parallel or intersecting scratches and groovings, varying from such lines as might have been graven by a hard grain of sand to such deep ruts as would require the forcible pressure of some sharp edge or blunt corner of stone. These markings are seen to run in a general direction down the valley. They have been evidently produced by some agent which has descended with sufficient force and steadiness to grind the rocks along the bottom and the lower slopes of the valley.

30. It is now and then possible to creep in under the ice at the end of the glacier, and to see where it rests upon its rocky bed. At such times you may, as it were, catch the glacier in the very act of grinding down and striating the rocks below it. Pieces of stone and grains of sand are jammed between the ice and the rock over which it

moves. Held there and pressed against the rock, they groove and scratch it. As this goes on year after year, the surface of the rock necessarily undergoes continual waste, and acquires that smoothed, polished, and striated appearance so characteristic of the bottom and sides of glacier-valleys.

31. Here, then, is the cause of the unfailing muddiness of the water which issues from the end of a glacier. The glacier is busily engaged in wearing down its channel with the same kind of grinding powder which a river uses—the sand and stones that fall into it from the



FIG. 71.—Stone polished and striated under glacier-ice.

sides of its valley—but employing these materials in its own way, and forming with them a peculiar smoothed and grooved surface, which no other agent in nature can produce. And here we see what becomes of the earth, sand, and stones which fall from the moraine heaps through the crevasses. They form the hard materials by which the ice erodes its channel. While they are used to wear down the solid rocks, they themselves are undergoing constant wear and tear. They become smoothed, polished, and striated like the solid rocks. One of these ice-smoothed and striated stones is represented in Fig. 71.

32. We conclude, then, that another great task in which every glacier is ceaselessly engaged is the erosion of the sides and bottom of its valley. That this is an important work in regard to the general scenery of a country may be shown by the great height and the long distances to which the peculiar forms of ice-worn rocks may be traced. The former greater thickness and wider extent of the Alpine glaciers are not more decisively shown by the range of the erratic blocks than by that of the polished and striated rock-surfaces. By this test it can be proved that a great part of Northern Europe and America has been under moving sheets of land-ice, which, passing over the land from mountain-ridge to sea-shore, have left behind them their memorial in the almost indelible markings engraved upon the rocks.

33. Again, no one can see the tumultuous body of discoloured water escape from a glacier without appreciating that in time a sensible deepening of the valley must take place in consequence of this ceaseless erosion and removal of materials. The deepening cannot be uniformly spread over the whole of the valley. There are places where the ice exerts more grinding power than at others, as a river at its falls and rapids effects more destruction than among meadows and plains. The rocks, too, of the glacier's bed vary much in hardness and power of resistance. Some parts must be more readily scooped out than others. So that, should the glacier retreat up the valley, these more deeply excavated portions, unless concealed under moraine rubbish, would become basins filled with water. Lake-basins of this kind in the midst of ice-worn rocks are a marked feature of glacier districts and of all those regions of Northern Europe and Northern America which have just been referred to as having been at one time buried under ice (see Lesson XXVI. Art. 4).

LESSON XXIX.—THE SCULPTURE OF THE LAND.

1. In Lesson XX. we traced the character of the leading features of the general external form of the land—its mountains, table-lands, valleys, and plains. At the end of that Lesson the question naturally arose whether any explanation could be given of the origin and history of these various features, but the answer to this question was postponed until after some consideration had been given to the nature of the materials and the internal constitution of the globe, and to the operations of water upon the surface of the land. We are now therefore in a position to return to the subject and to apply to the investigation of it the facts and deductions about the interior and exterior of the land which have come before us in the last eight Lessons.

2. At the outset of this interesting inquiry it should be borne in mind that the existing land consists of the higher parts (projecting above sea-level) of those ridges into which the exterior of our planet has been wrinkled during its gradual consolidation and contraction from an original fluid condition (Lesson XXII. Art. 20). We must not hastily conclude however that the land, such as we now see it, is part of the original surface of the globe. That it cannot be so must be evident from two considerations to which reference has already been made. 1st. All the land as far as we know has been under the sea, and even up to its mountain tops consists in great part of hardened and altered sand, mud, and other materials, which were originally laid down upon the floor of the sea (Lesson XXI., Art. 13). 2nd. The whole surface of the land is subject to a constant and enormous though unequally

distributed decay and removal. A comparatively short period would suffice for the entire destruction of the continents were their surfaces to be continuously wasted even at the rate of the Mississippi's operations—and other rivers work considerably faster. At any probable rate of degradation the original land must have been long since destroyed, and we can hardly hope to find any trace of it even buried under the later accumulations of which the continents consist (Lesson XXVII., Art. 22).

3. But though no portion of the present land can be looked upon as part of the original or earliest solid surface of the planet, there can be no doubt that the existing continents must be very old. Not improbably they occupy the sites of the first ridges which appeared upon the cooling and shrinking mass of the globe. In the course of ages these primeval ridges would be worn down by the action of water and air. But from time to time, if renewed uprisings took place along the same original lines, the land would be formed and destroyed, and then formed and destroyed again. That this is not mere theory, but rests on a strong basis of probability, may be shown by a consideration of the way in which the materials forming the land have actually been put together. If the existing ridges can be shown to have been upraised again and again during past ages they may at least be plausibly conjectured to mark generally the primeval lines of elevation on the surface of the globe.

4. Let us then return to the composition of the earth described in Lesson XXI. All over the globe it is found that by far the largest mass of the land is built up of materials which have been accumulated slowly as sediment on the floor of the sea. These materials are arranged in layers or strata which have been laid down upon each other until a depth of many thousands of feet has been formed. It is evident that the original position of these

strata must have been nearly or quite horizontal, seeing that they were piled one upon another as sand and mud are laid down on the level or gently sloping bed of the sea at the present time. The subterranean movements by which they have been raised above the sea into dry land have taken place over such wide regions that this original level or gently inclined position may remain little or not altered. The flat bedding of the rocks shown in Fig. 26 is of common occurrence, and most people will recognise the familiar look of such sections as that of the quarry

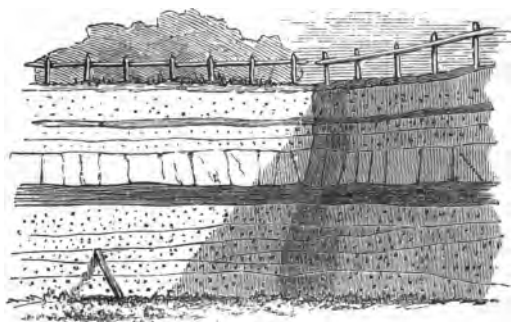


FIG. 72.—Quarry in flat stratified rocks.

shown in Fig. 72. In these cases we may conceive a large tract of the sea-floor to have been raised up into land so uniformly and equably, that the sheets of hardened sand and mud remained horizontal. In Central and Northern Russia, in the United States, and in Canada, this gentle and equable upheaval has taken place over regions many thousands of square miles in extent.

5. While the earth has been contracting as a result of its cooling, the effects of this contraction have not been



FIG. 73.—Section across a mountain-chain to show how the level rocks of the plains are bent and inverted along the flanks of the mountain, while the lowest and oldest rocks are made to form the central and highest point of the chain.

uniformly distributed over the surface. The vast basins of the oceans no doubt mark the regions where the subsidence has been greatest and most continuous. Probably they have been depressions from the beginning, though portions of them, particularly along their margins, have from time to time oscillated up and down. Every tract which sinks requires of course to accommodate itself to a diminished superficial area, and tends to exert a strong lateral thrust upon the adjoining more stable parts. Under the influence of this force long ridges have been raised into land between the ocean basins. Every successive subsidence may thus have carried with it a corresponding upheaval. So that although subsidence was the rule, and although the land was being continually wasted by air, rain, rivers, and the sea, nevertheless these periodic uplifts have compensated for the loss, and seem to have maintained, on the whole, the balance of dry land.

6. But it could hardly happen in the midst of these movements that the upheaval should be always so gentle and uniform as not to disturb the original level, or nearly level, position of the strata. On the contrary, over wide tracts of land, and more particularly along vast extended lines, the rocks have not only been upraised, but have been crumpled up and broken. Instead of remaining in horizontal or slightly inclined sheets, they may be found tilted

up in all directions, and often placed on end like books on a library-shelf. Every great mountain-chain furnishes examples of these more complex arrangements. On the plains and lowlands the rocks may stretch for hundreds of miles as level as before they rose out of the sea. But towards the interior they begin to bend in wave-like undulations, which increase in magnitude until, along the flanks of the mountains, the rocks are sometimes found actually so thrown over that the lowest lie uppermost. This structure is explained in the accompanying diagram (Fig. 73).

7. In this section, which represents an extremely simple form of mountain-structure, there is evidence of only one



FIG. 74.—Section across a mountain-chain, showing two successive periods of uplift.

general upheaval. The date of the movement must of course be subsequent to the formation of all the rocks of the region, for they have all participated in the change of position. A rather more complex case is illustrated in Fig. 74. Here there is proof of two upheavals. First, the older series of rocks, A, was contorted and raised; then, against its sides and upon its broken and worn edges, the series B was formed, and thereafter raised into land. A still longer succession of movements is represented in Fig. 75. We there see that after the two movements which disturbed the rocks A and B, a third uplift took place, whereby a still newer series, C, which had been deposited upon the side of B, came to be raised into land.

9. Two facts about mountains are presented to the mind by such sections as these. First, that a great axis of elevation on the earth's surface may have again and again served as a line of relief from the strain of terrestrial contraction, and may have consequently been successively pushed up between the subsiding areas on either side. And secondly, that between each period of uplift there has been great waste of the upraised land—air, rain, frosts, brooks, rivers, glaciers, and the sea, all wearing down the surface and producing the materials out of which the next series of rocks was formed.

10. It is evident also that by means of such sections we may compare the relative dates of different mountain-chains. If the rocks represented in Fig. 73 be the same as



FIG. 75.—Section across a mountain-chain showing three successive periods of uplift.

those marked C in Fig. 75, then the mountain shown in the former section must be much younger than that shown in the latter. We see indeed that the mountain in Fig. 75 is not only older, but older by two earlier upheavals. It is the province of the science of Geology to investigate these questions regarding the structure and age of mountain-chains. Geologists have discovered that among the stratified rocks of the earth's crust, a determinate order can be made out, and they can thus assign to each great elevation on the surface of the globe its relative date in the history of the mountains.

10. But while the existence of dry land and of the

mountain-chains which traverse it is due to the movements of the solid crust of the earth, it is clear that the present aspect of the land must have been largely determined by the action of those various agents which wear down its surface. The vast amount of mud annually transported into the sea by rivers, proves how much material is continually removed from the land, and therefore, how greatly, though it may be insensibly, the height and appearance of the mountains and valleys must in course of time be changed. This process of disintegration and removal is going on all over the globe, here more rapidly, there more slowly, but always advancing, and always bringing with it changes in the aspect of the land. In the lapse of the long periods of time during which it has been in progress, how vast must have been the mutations on the surface of the globe, how many successive mountain ranges may have been upraised and worn away!

11. We have followed the working of the chief agents by which the surface of the land is eroded—the air by its gases and vapours, its winds and changes of temperature; frost by its oft-renewed wedges of ice; rain, brooks, and rivers by their movement down the land, and their power to sweep before them the loosened debris; avalanches and glaciers by the fragments of stone with which they grind and polish the rocks of the valleys; the sea by the waves thrown incessantly against the shores which bound it. We may compare the general results of the co-operation of all these forces to the work of a sculptor. They are, so to speak, the different tools with which the framework of the land is carved. But the sculpture they achieve is never completed. It goes on continuously so long as the land remains above the sea.

12. At first it may seem almost incredible that the whole surface of the land, even the loftiest and stateliest

mountains, should thus be crumbling down. But the more we search for proofs the more clear and abundant do they become. We learn that whatever may have been the aspect of the land when first pushed out of the sea, it has been, and is now being, chiselled from its highest peaks down to below the tide-marks. Its cliffs and pinnacles are split up and grow more shattered and sharp every year. Its ravines are widened and deepened. Its hilly surfaces become more roughened and more deeply seamed by the lines which running water traces over them. Its valleys and plains are levelled and strewn with the debris washed down from the higher grounds.

13. In travelling from place to place you will do well to note as you go the proofs of this universal decay. While so doing you will not fail to observe that though the wearing down of the land may be traced more or less clearly everywhere, its rate and the changes of scenery which it brings with it depend very much upon the nature of the rocks of each region. Here again we may have recourse to the simile of the sculptor's work. The character of a statue depends not only on the design and manipulation of the artist, but also on the material employed. Out of a piece of granite or of pudding-stone, no matter what amount of genius and skill were bestowed on it, the same effect could never be produced as from a block of pure white marble. So we find that the hills, valleys, and mountains differ from each other in great measure according to the nature of the rocks of which they consist. In your journeys, whether in your own district or in other regions, you will find it not uninteresting or unimportant to take note of the changes in the aspect of the scenery through which you pass, and to connect them when you can with variations in the character of the rocks.

14. Apart, however, from the varying nature of the materials, nothing contributes more to the character of

scenery than those lines of division or "joints" which have already been referred to as traversing all rocks. They serve as channels for the descent and reascent of subterranean water (Lesson XXIII. Art. 17). They are made

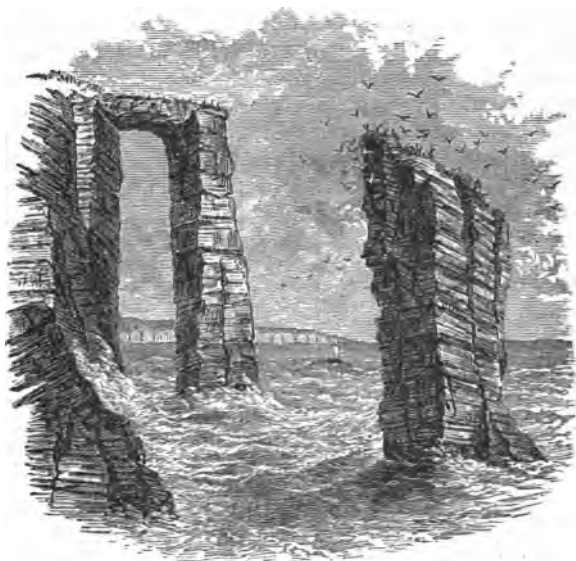


FIG. 76.—Scene on the Coast of Caithness.—Influence of joints among stratified rocks in the formation of vertical cliffs and outstanding square blocks of rock.

use of by frost as the lines along which the wedges of ice are driven most effectively into exposed faces of rock (Lesson XXVIII. Art. 10). Every mountain-peak and cliff, indeed every large naked mass of rock which projects into the air, owes more or less of its characteristic outlines to

the way in which its joints have been split open. Among the stratified rocks the joints allow vertical cliffs to be formed and large square buttress-like masses to project from the cliff or even to be isolated from them. The accompanying drawing (Fig. 76), for example, shows how the forms of coast-cliffs are determined by the position of the intersecting lines of joint, each vertical face of rock corresponding with the direction of one of these lines. Where such rocks form lofty mountainous ground they are often carved into the most picturesque forms of pinnacle and buttress. Again, among the unstratified rocks, such as granite and basalt, the influence of the joints is no less marked, as, for instance, where it defines the ledges and rifts in a precipice, or where it has allowed the most solid rock to be so completely shattered as to look like a huge mass of ruin. In Fig. 77 a representation is given of part of the face of a basalt cliff, where the abundant joints traverse the rock in such a way as to divide it into rude prisms, which are gradually wedged off from each other by frost, until, detached at last, they fall to the base of the precipice. The ruined masses below are further broken up and carried away piecemeal by frost, rain, and in the general process of "weathering," or in other cases, by the waves along the sea-shore or by the flow of a brook or river. Their removal permits the continuous decay of the cliff and preserves the steep face of the precipice, which thus slowly recedes as slice after slice is cut away from its front. But where the detached blocks gather at the base they form in the end a protecting bulwark there, and either retard or prevent the further recession of the line of cliff.

15. It is among the higher parts of the mountains that this kind of rude chiselling of the rocks can be most conspicuously seen. Not only are steep crags and lofty precipices formed, but the very mountain ridges are cut

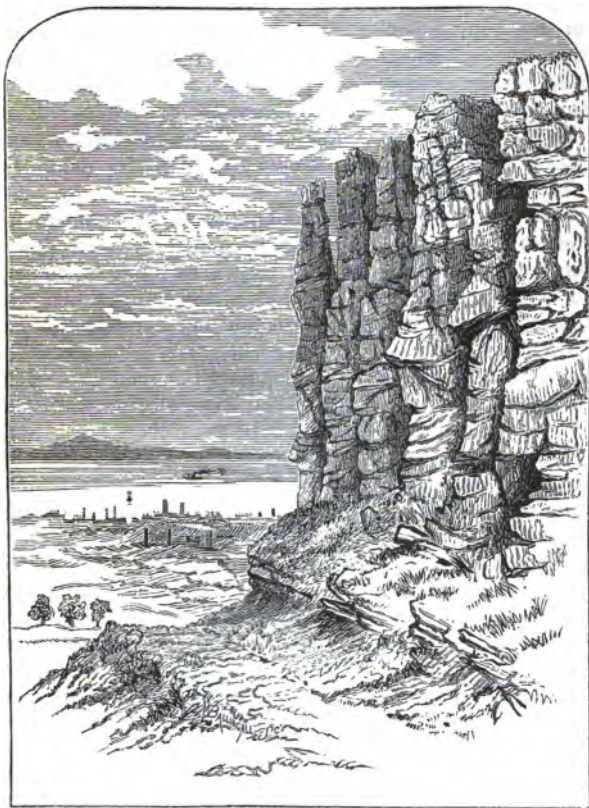


FIG. 77.—Portion of the west front of Salisbury Crag, Edinburgh, showing the influence of joints in promoting the splitting up of the rock (basalt), and the preservation of a vertical face to the cliff.

away into sharp crests. These, again, still further split and splintered by the severe frosts and the furious storms of the mountain climate, are cut into slender pinnacles and spires, sometimes at a distance seeming so needle-like in their slimness and sharpness, as to have received among the Alps the name of *aiguilles*, or needles. It is the blocks loosened from these high crags and crests which furnish abundant stones for the glacier moraines (Lesson XXVIII. Art. 22).

16. From the top of a high hill or of a mountain beneath the snow-line one may sometimes look down upon a wide region and mark, as in a vast map or model, how the little gullies on the sides of the slopes widen out into larger channels, how these run together into valleys, and how the whole landscape seems thus to be trenched with water-courses. When you have the good fortune to see such a scene as this, if you have already learnt to appreciate how ceaselessly and potently every brook and river is cutting out its channel (Lesson XXVI.), you will realise more vividly than from any map or description how the valleys are carved out by the power of running water. Every little gully and ravine down the steeper declivities around you is a sample of the way in which the forms of the solid land are changed ; you may trace every gradation of size and shape from the trench that was opened by some storm last winter up to the deep and wide gorge through which the foaming river rushes, or the broad ample valley down which the collected waters from a whole range of mountains sweep onward to the sea. You may not be able to tell how far the line of any particular valley may have been originally determined by the shape which the ground had when it first rose out of the sea. But as a heavy shower of rain produces runnels which soon cut out a miniature drainage system on a road-way, so in the course of time the flow of

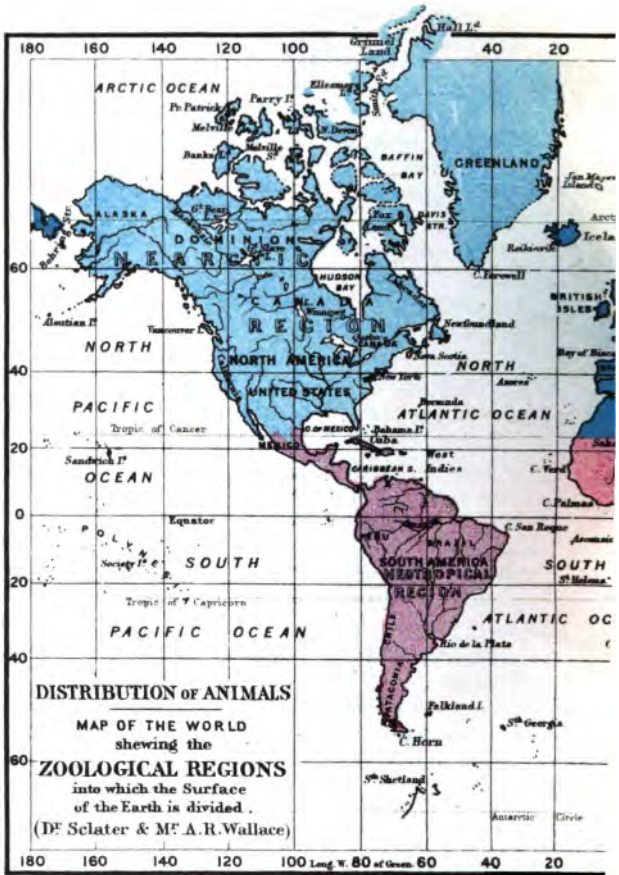
brooks and rivers over the surface of the land must necessarily erode systems of valleys. You will feel assured therefore, that whatever might have been the original shape of a country, rain and frost, brooks and rivers, snow-fields and glaciers could not have been at work upon it for even a comparatively short time without carving out valleys for themselves, and sculpturing the mountains into such sharp and rugged forms as they now wear. Everywhere you see proofs of excavation; hill-sides are furrowed with gulleys, mountain slopes are trenched with ravines, table-lands are cut down until they become only chains of ridges separating the valleys which have been carved out of them (Lesson XX. Art. 19).

17. While the effects of disintegration in roughening the surface of the land are most marked among the high grounds, the results of this process show themselves among the lower regions in the strewing of the crumbled fragments of the hills over the valleys and plains. Every tract of fertile meadow and level field bears witness to the way in which the lowlands are smoothed and heightened by the sand and earth spread over them by the streams which come from the hills. And yet this increase of the height of the plains does not really compensate for the waste of the high-grounds. A little consideration of the matter shows, in the first place, that though the plains do obtain considerable additions to their surface from the materials swept down upon them by rivers, they receive only a part of these materials, the remainder being carried out to sea by the rivers; and in the second place, that even the plains themselves are wasted; floods tear up their soil and sweep away their river-banks.

18. It appears, then, that the tendency of the process of sculpture, which gives to landscapes their characteristic details of outline, is in the end to reduce the dry land to the level of the sea. But this is not all. While the

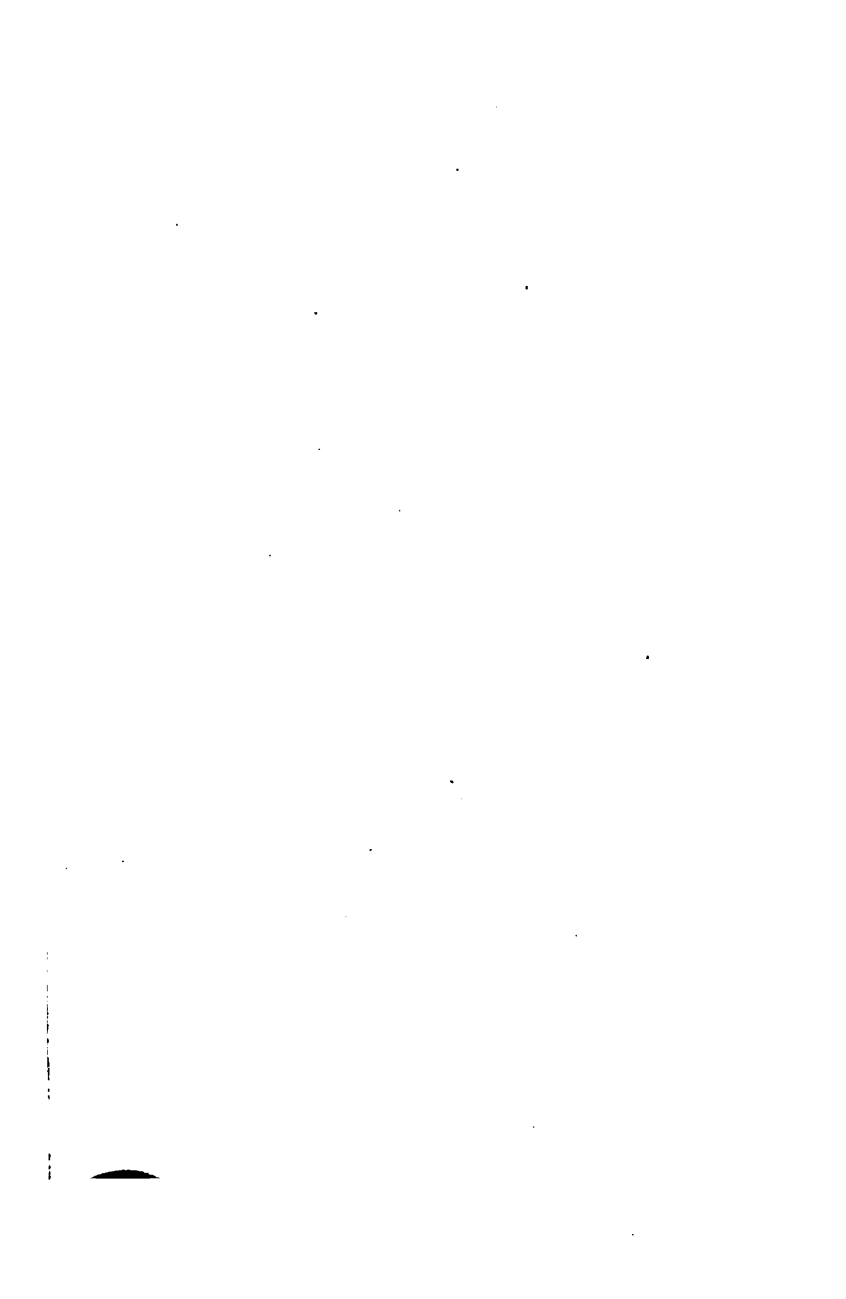
general surface of the land undergoes attacks from the atmospheric influences, its margin is continually suffering from the assaults of the waves. It is indeed only the parts of the earth's surface lying under a considerable depth of ocean which are protected from decay (Lesson XVIII. Art. 19). Elsewhere the waves are gnawing away the coast line, or are only kept from doing so by the bar of detritus which has been thrown up against them.

19. Were no other operation to come into play, the natural and inevitable result of this ceaseless destruction would be the final disappearance of dry land. But here we see the meaning and importance of the underground movements already referred to as the results of terrestrial contraction ; the great ocean basins have from time to time sunk down, and in so doing have pushed up ridges of land between them. These ridges on each successive uplift would consist mainly of the more or less consolidated debris worn away from their predecessors. The same materials have thus served over and over again to form the framework of the upraised land, and thus while looking at the subject from one side we see only ceaseless destruction, mountain and valley continually crumbling down before us, yet, when we take a wider view, we perceive that the decay of the surface is needed to furnish soil for the support of the living plants and animals which people the earth, and that the materials so removed from our eyes are not lost, but are carefully stored away to be in some future time raised into new land, and thereafter to go again through the same cycle of change.





Edinburgh and London.



CHAPTER V.

LIFE.

LESSON XXX.--THE GEOGRAPHICAL DISTRIBUTION OF PLANTS AND ANIMALS.

1. IN the foregoing Lessons we have considered the parts of the earth, their relation to each other, and the constant changes and reactions between them which constitute the Life of the globe. But above and apart from all these movements within and upon the surface of the earth, another kind of activity and progress now claims our attention, where the forces concerned are not air, sea, and land, but the living energy of plants and animals. Our planet is not merely a theatre for the evolution of physical phenomena. It has been appointed as the dwelling-place of a vast and varied series of living things, which move through the air and people both land and water. The study of these living organisms is comprised under the general name of Biology, or the science which deals with vegetable and animal life.

2. So vast a study, opening up wide fields of inquiry far beyond those over which we have been travelling in these Lessons, must needs be subdivided into different departments. Thus one branch inquires into the structure and growth of plants, another deals with the way in which

plants are distributed over the globe, a third treats of the structure of the various tribes of animals, a fourth follows the action of the different parts of an animal's body and the part which each of these plays in the life of that body, a fifth arranges the enormous numbers of animal forms in due order, to show their grade in the scale of being, and to allow the general assemblage of living forms in one region to be compared with that in another. These and the other branches of biology deal chiefly with the plants and animals as they are in themselves, or as they stand in relation to each other.

3. But it is evident that just as we ourselves are encompassed by external conditions of geography, climate, and vegetation which, it may be unconsciously, govern our everyday life, so each plant and animal on the globe comes under the control of the same surrounding influences. Apart therefore from the structure, or functions, or classification of vegetable and animal life, we may study it with reference to its relation to these external and dominant conditions. From this point of view physical geography and biology are seen to be closely linked to each other. It is impossible to gain any intelligent conception of the present distribution of plants and animals over the globe without entering upon inquiries which form part of the scope of physical geography.

4. We all know how greatly the plants and animals of different quarters of the globe differ from each other. The equatorial regions nourish a rank and luxuriant vegetation, including palms, bananas, tall tree-grasses with rope-like lianas twisting round their stems, and bright-hued, strangely-shaped orchids hanging from their branches. The animals are equally characteristic, for they include lions, tigers, elephants, rhinoceroses, camels, giraffes, crocodiles, large serpents, with crowds of gorgeously-plumed birds and brilliant butterflies.

5. The temperate zone is distinguished by a very different assemblage of plants and animals. The forests and woods show such trees as the oak, ash, elm, sycamore, beech, poplar, birch, hazel, and pine. The dells are bright in spring with snowdrops, anemones, and primroses; and in summer with speedwells, geraniums, and wild roses. But both the plants and the animals are more sober in colouring than in the hotter parts of the earth. The birds include the thrushes, larks, and other songsters. Among the wild animals of the low-ground we find mice, rats, weasels, hedgehogs, badgers, otters, foxes; and in hilly districts wild-cats, wolves, and bears.

6. Within the Arctic and Polar regions life becomes much less abundant and varied. As we advance trees disappear, though stunted forms of birch, fir, and willow extend a long way northwards. By degrees these too die out, and the scanty vegetation consists mainly of mosses, lichens, with saxifrages, gentians, and a few more flowering plants. These snow-covered lands are wandered over by polar bears, white foxes, reindeer, musk-oxen, ermines; the seas are frequented by seals, walruses, and whales; while the coasts are sought by large flocks of northern sea-fowl and by snowy falcons, buntings, ptarmigans, owls, and other white-feathered birds. At the extreme northern limits reached by explorers, life of any kind is hardly to be met with among the deep snow-fields and piled-up heaps of ice which cover land and sea.

7. We see then that from the extreme exuberance and variety of plant and animal life in the equatorial and tropical lands there is a gradual diminution, until in the far polar regions we reach almost the zero of vegetation and of animal existence. There can be no doubt therefore that one grand influence in the distribution of plants and animals is temperature, warmth being favourable, cold unfavourable, to the growth of living things.

●. But if temperature were the sole cause that determined the character of the plants and animals of any country, then every zone of latitude should be marked by the same kind of vegetation and by the same groups of animals. It would then be enough to know the geographical position of any place to be able to tell what must be its *flora*, that is, its assemblage of plants, and its *fauna*, or population of animals. But a very little inquiry suffices to show that no such strict coincidence between latitude and botanical or zoological distribution really exists. The Old and the New Worlds are traversed by similar isothermal bands, and have the same kinds of climate, soil, and exposure. Yet even in regions closely resembling each other as to conditions of physical geography the plants and animals, though sometimes presenting such general resemblance as to show that they are mutually representative of each other, are sometimes widely different in the two hemispheres. The lion of the Old World gives way to the puma in the New; the tiger is replaced by the jaguar; the elephant, rhinoceros, and hippopotamus, by tapirs and peccaries; camels by vicunas; apes and baboons by flat-nosed monkeys and marmosets. The birds are not less distinct; the Old-World vulture being represented by the New-World condor; sun-birds by humming-birds.

●. The present arrangement of plants and animals over the earth's surface is one of the most difficult questions with which science has to deal. The first step towards its investigation must necessarily be a careful inquiry into the actual facts regarding the distribution of the various kinds of plant and animal life. This inquiry has been carried on by naturalists in all parts of the world with the result of enabling them to parcel out the earth's surface into distinct regions and sub-regions, each characterised by a peculiar flora and fauna, though sometimes containing a number of the same plants and animals in common.

These regions, to which the following names and limits have been assigned, are shown in Plate X.¹

10. (i.) **Palaearctic Region**, comprising all Europe with the temperate parts of Asia, and the tracts of Africa lying to the north of the Sahara desert ; or the northern parts of the Old World, from Iceland to Behring Strait, and from the Azores to Japan. In the Arctic portions of this vast region the vegetation is comparatively meagre, showing an abundance of mosses and lichens, which in the tundras of Siberia cover thousands of square miles of barren waste. In favourable places flowering plants, such as saxifrages and gentians, peep out from beneath the snows during the short, but warm summer ; stunted forms of willow, azalea, and rhododendron form here and there a kind of scrub upon the slopes, while southwards of the mean annual isotherm of 32° pine-trees make their appearance and increase in number, till they form wide ranges of dark forest, as in Norway, and on the higher mountain groups further south. Between the isotherms of 40° and 60° the sombre pine-forests, retaining their leaves throughout the year, give place to a more varied and luxuriant *deciduous* vegetation, which sheds its leaves in autumn and renews them again in spring. The trees include many noble kinds, birch, alder, beech, ash, oak, elm, sycamore, walnut, chestnut, and maple. Northern fruits, like the cranberry, cloudberry, bilberry, strawberry, currant, and raspberry, are succeeded further south by luscious pears and apples, almonds, olives, figs, grapes, and oranges. The cereals—wheat, barley, oats, &c.—are abundantly cultivated throughout most of the region. In the more southern countries, with a mean annual temperature, between 60° and 70°, the trees do not lose their leaves in winter. Here we find the evergreen oak, myrtle, laurel, with some plants,

¹ The arrangement here followed is that of Mr. Sclater, as modified by Mr. Wallace in his recent admirable work on "The Geographical Distribution of Animals."

such as palms, which more properly belong to the tracts lying nearer the equator.

11. The animals of the region show a development similarly connected with the distribution of temperature. In the extreme north white bears and foxes, reindeer, whales, walruses and seals, many peculiar sea-birds, together with white owls and ptarmigan, form a distinct assemblage. As most of these species disappear southwards, their places are taken by brown bears, badgers, otters, horses, buffaloes, fallow and roe-deer, chamois, wild goats, wild sheep, hares, rabbits, moles, hedgehogs, and dormice; golden eagles, hawks, grouse, pheasants, house-sparrows, magpies, jays, and thrushes. Further south the camel is the most characteristic animal.

12. (ii.) **The Ethiopian Region** embraces Central and Southern Africa, with the tropical part of Arabia, Madagascar, and the neighbouring islands. To the south of the great desert of Sahara the western portion of the African continent is largely covered with dense forests, where the oil-palm, the huge baobab, euphorbias, bignonias, tamarinds, and many other tropical plants form a dense luxuriant vegetation in the hot moist air of that climate. Further to the east lie vast elevated lands, covered with tall grasses and sedges, and dotted with patches of forest. The flora of the southern part of Africa is distinguished by its great variety of heaths, its fig-marigolds, carrion-scented stapelias, aloes, and pelargoniums.

13. The fauna of the Ethiopian region is marked by the absence of such wide-spread animals as camels, deer, goats, sheep, wild oxen, wild boars, and bears, and by the presence of many remarkable forms of life, including the gorilla, chimpanzee, baboon, lemurs, aye-aye, lion, leopard, civet, hyæna, zebra, rhinoceros, hippopotamus, giraffe, antelope, elephant, ostrich, ibis, flamingo, chameleon, and crocodile.

14. (iii.) **The Oriental Region** includes Southern Asia from the mouth of the Indus along the southward slopes of the Himalaya mountains and the Chinese uplands to Ningpo, with Formosa, the Philippines, Borneo, and the Malay Islands as far as the south-east end of Java. Much of the surface of this region is covered with dense forests of tropical vegetation. Among the better-known plants occur the ginger, arrow-root, banana, cocoa-nut, screw-pine, yam, bamboo, rice, gourd, custard-apple, mango, coffee-tree, mangrove, ebony-tree, bignonia, hemp, and sandalwood.

15. The fauna contains some characteristic animals, the ourang-utan, long-armed monkeys, flying lemur, many civets, the tiger, hyæna, jackal, wild cattle, elephant, rhinoceros; many bright-feathered birds, as trogons, hornbills, goat-suckers, sunbirds, long-tailed parrots, and peacocks; numerous reptiles, including ground and tree-snakes, cobras, and crocodiles; and a vast assemblage of insects, among which the size and brilliancy of the butterflies and many of the beetles are remarkable.

16. (iv.) **The Australian Region**, embracing Australia, New Zealand, and the numerous islands to the east of Java, Borneo, and the Philippine group, consists wholly of islands, which, lying apart from all the great continental masses of land, show a peculiar assemblage of plants and animals. The vast insular expanse of Australia, situated partly within and partly without the tropics, and exposing a wide desert interior to the hot rays of the sun, while its coast-line is washed by the open sea, presents contrasts of climate not met with in the smaller islands of the region. From its size, also, and its proximity to the south-eastern limits of the Oriental region, it contains a greater diversity in its flora and fauna. Over the dry and warm tracts of Australia the general heath-like vegetation is marked by a pervading dead blue-green colour, with

dull leaves so arranged upon the plants as to afford but little shade. The eucalyptus, or gum-tree, and other trees and shrubs bearing bright honeyed flowers, together with thickets of acacia and scattered marsh-oaks, give a peculiar character to the forest-lands. Vast regions are covered with grasses, and furnish good pasture. Along the northern limits, where this region borders the Oriental islands, the flora contains some of the forms of vegetation found more plentifully towards the north and north-west, such as pandanus, cabbage-palm, fig, nutmeg, and sandalwood. On its southern margin, where the climate assumes a moister and more temperate character, ferns, cycads, and pines abound, while the heath-like epacris, and numerous proteas enliven the surface with their bright blossoms. New Zealand is distinguished by the verdure of its flora, which consists largely of ferns, often growing as trees, and many kinds of pines. The scattered islands of the Pacific have their cocoa-nut palm, bread-fruit, tacca, grasses, and sedges.

17. The fauna of the Australian region is the most peculiar on the face of the globe, both for the types which it contains, and for the almost universally diffused forms which it has not. No apes or monkeys chatter in its woods, no wild horses, cattle, or sheep browse over its pastures, no wolves, foxes, tigers, or other similar beasts of prey prowl across its hills and valleys. The place of these various and wide-spread species is taken in Australia by a totally distinct and less highly-organized class of animals called marsupials, of which the kangaroo may serve as the type. Of these there are many varieties, some living on fruits, others on roots, others on smaller animals of their own kind, others on insects; some keeping to the ground, others taking to trees. The birds likewise are peculiar, for they include the bird-of-paradise, lyre-bird, cassowary, paroquets, and honey-suckers.

18. (v.) **Neotropical Region.**—Under this term are included the whole of South America, the islands of the Antilles group, and the tropical part of North America. Ranging across the whole zone of the tropics, and as far south as the 56th parallel of south latitude, and rising up to the snow-line in the Andes, this region presents many varieties of climate, which are well shown by differences of vegetation. The lower grounds within the tropics present the most luxuriant flora in the world. It abounds in mangroves, palms (cabbage-palm, ivory-palm, and other kinds), bananas, tree-ferns, and mimosas, growing in dense jungles, and having their stems and branches clustered round with many smaller plants, such as lianas and ferns, or gorgeously-blossomed cactuses, orchids, and passion-flowers. The vast plains or *llanos* of the Orinoco, with their tall grasses and occasional clumps of pines and mimosas, represent the pasture lands of the Old World, but with their bright lilies show a brilliancy of colour peculiar to themselves. Further south the basin of the La Plata presents similar plains or *pampas*, which, getting less and less luxuriant in their vegetation, are at length succeeded by the barren moors of Patagonia and Terra del Fuego. On the lower mountain slopes, characteristic trees are cinchonas, from the bark of which quinine is prepared. Mahogany, rosewood, the indiarubber tree, with many plants yielding spices, balsams, and perfumes, give a distinctive character to the South American flora. On the more elevated tracts, calceolarias, gentians, and low-growing plants occur, that remind the traveller of some features of mountain vegetation near the snow-line in the Old World.

19. The fauna of this region is regarded as more varied than that of any other. It contains the peculiar jaguar, flat-nosed monkeys, and marmosets, blood-sucking bats, chinchillas, sloths, armadilloes, ant-eaters, racoons, opos-

sums, deer, llamas, alpacas, vicunas, tapirs, and peccaries ; but no native sheep or oxen. Among the birds occur condors, curassows, rheas, or American ostriches, toucans, jacamars, mot-mots, macaws, and numerous forms of humming-bird. The reptiles include the boa-constrictor and many other serpents, the alligator, crocodiles, tortoises, and turtles. The insect life is immensely abundant and varied.

20. (vi.) **The Nearctic Region** embraces all North America lying to the north of the tropic of Cancer. Its greatest breadth being towards the cold northern regions, it rapidly narrows southward, so as to be connected with the Neotropical region merely by a narrow strip of land. This isolation is accompanied by a somewhat less varied flora and fauna than in the corresponding region of the Old World. The plants and animals, taken as a whole, present much less contrast to those of the Palæarctic region than those of the Neotropical and Ethiopian regions do to each other. *The northern tracts of North America extend far within the Arctic circle, into the snow-covered lands where vegetation reaches its lowest point of development. The southern limits of the province, on the other hand, lie towards the tropical zone, where the sugarcane, yucca, cotton, maize, and tobacco are characteristic plants. In California and Oregon many large and distinct kinds of pine occur in the forests, such as the gigantic Wellingtonia and the Douglas pine. Eastward of the Rocky Mountains vast undulating pasture lands or prairies stretch over the basin of the Mississippi. The British Possessions lie mostly upon forests, which, as the country becomes peopled, are gradually giving way to pasture and cultivation.

21. The fauna varies with the latitude. In the north are found musk-sheep, moose-elks, reindeer, gluttons, skunks, racoons, beavers, lemmings, jumping-mice, and tree-

porcupines. Further south vast herds of bisons roam over the prairies. Other typical animals are the grizzly bear black bear, puma, lynx, prong-horned antelope, prairie-dog, flying-squirrel, pouched-rat, opossum, humming-bird, blue-crow, and rattle-snake.

LESSON XXXI.—THE DIFFUSION OF PLANTS AND ANIMALS.

Climate.—Migration and Transport.—Changes of Land and Sea.

1. IT is not enough to know how the various tribes of plants and animals are distributed over the surface of the globe. We are irresistibly led to ask ourselves why and how their distribution has come to be as we have traced it in the preceding lesson. Not very long ago men were content with supposing that the present arrangement had always existed ever since the different continents and islands rose out of the sea and received their earliest inhabitants.

2. But in the gravels and clays beneath the soil, or in the limestones, sandstones, and other rocks lying below, traces have been found all over the world of older and different plants and animals, which occupied the land before the present denizens had appeared. Our modern horses and cattle, for example, were preceded by other kinds which are no longer living. The bears, wolves, and hyænas of to-day are not the same as those of which the teeth and bones are found in caves and ancient alluvial deposits.

3. A thoughtful study of this subject suggests three main influences which may have guided the distribution of the present faunas and floras of the earth's surface.

1st. Climate. 2nd. Migration and transport. 3rd. Changes in the form and height of the land and in the depth and extent of the seas. A knowledge of the nature and effect of these influences helps us to understand much that would otherwise be inexplicable, but there still remain, and perhaps must ever remain, many difficulties which no amount of research may be able to remove.

I.—Climate.

4. This term includes the general temperature, moisture, winds, and other atmospheric conditions which prevail in any district, and which directly affect the growth and vigour of plants and animals. From the statements made in Lesson IX. it appeared that the climates of different portions of the globe greatly differ, and some of the causes of such differences were there traced. But a knowledge of the annual distribution of heat at any place, though it gives us one main element in determining the climate of that place, requires to be enlarged by further knowledge respecting the rainfall, the direction of the prevalent winds, the shape, height, and position of the ground, the character of the soil, vegetation, and other more local details. Since Lesson IX. you have been led over most of these subjects, so that you may now return to the question of climate in reference to the arrangement of plants and animals.

5. On consideration we may recognise five distinct influences by which the climate of any place is determined. 1st. Distance from the equator. 2nd. Distance from the sea. 3rd. Height above the sea. 4th. Prevailing winds ; and 5th. Local influences, such as soil, vegetation, and proximity to lakes or to mountains. Each of these causes directly, and often powerfully, controls the spread of vegetable and animal life.

6. (i.) Distance from the Equator.—Climate, having temperature for its main element, must follow the same general distribution over the earth's surface. The warmest climates are necessarily those of the inter-tropical regions, where the sun's rays are vertical, or not much inclined from vertical. In proportion as we recede from the equator the rays fall more and more obliquely, and the same amount of heat-rays is therefore spread over an increasing breadth of surface, while, moreover, they have to pierce a greater mass of air. Round the poles the least amount of heat is received, and the climates are coldest.

7. Were the earth's surface either wholly land or water the climates would be arranged in parallel and regular bands from the equator to the poles. But it was shown in Lesson IX. that, owing to the way in which land and water are grouped, such an arrangement has been prevented. Two places on the same latitude may have very different average temperatures, and therefore very dissimilar climates. Nevertheless the great predominating influence of position with regard to the equator is on the whole maintained in the arrangement of climates. It influences in a marked way vegetable growth, as is shown by the time of flowering or of ripening among widely-distributed plants. On the continents this time becomes later in proportion as the country is distant from the equator. Thus the elm comes into leaf at Naples about the beginning of February, at Paris not until late in March, and in the centre of England not until the middle of April. Ripe cherries may be gathered in the south of Italy about the beginning of May; they are ready in Northern France and Central Germany at the end of June; but not generally in England for three or four weeks later. Nothing could show more strikingly the difference of climates between the different latitudes of a continent.

8. But these differences are not merely marked by the variations in the growth of the same plants. As shown in the last lesson, when we pass from one climate to another we encounter different plants and different animals. One by one characteristic forms of life drop away and their places are taken by others. So constant and marked are these changes that such expressions as "an arctic vegetation," "a temperate flora," "a tropical fauna," have passed into general use, and convey a distinct picture to the mind.

9. We have found, however, that if such a distribution of plants and animals were due to differences of climate alone, wherever the same climate recurs it should be accompanied by the same kind of vegetation and of animal life, but that no general coincidence of this kind exists, when regions remote from each other are compared. The climate of the centre of Europe closely resembles that of parts of the United States. Yet the wild animals and birds are strikingly different; mice, hedgehogs, buffaloes, chamois, and jays in the Old World are replaced by jumping-mice, racoons, opossums, bisons, llamas, and humming-birds in the New. In central South America the forests are tenanted by jaguars, sloths, armadilloes, tapirs, curassows, and toucans. On corresponding latitudes in equatorial Africa these animals are represented by lions, leopards, hyænas, hippopotamuses, elephants, guinea-fowl, and touracoes. In Australia these forms are again replaced by a strange and peculiar assemblage of animals, including kangaroos, wombats, flying-opossums, emus, lyre-birds, and crested pigeons. While therefore difference of latitude usually means difference of climate and of plant and animal life, identity of latitude with similarity of climate does not necessarily imply agreement in the character of the flora and fauna.

10. (ii.) **Distance from the Sea.**—The influence of the

sea upon the distribution of temperature and moisture has been already described (Lessons IX., X., and XVIII.). As water is more slowly heated and cooled than land, the climates of the sea and of the coasts of the land are much more moist and equable than those of the interior of the land. In proportion therefore as places recede from the sea their climates become more extreme. An insular or oceanic climate is one where the difference between summer and winter temperature is reduced to a minimum, and where there is a copious supply of moisture from the large water-surface. A continental climate is one where the summer is hot, the winter cold, and where the rainfall is comparatively slight.

11. These variations cannot but make themselves visible in the distribution of plant and animal life. They are well shown by contrasting the times of flowering and ripening of the same plants along the Atlantic border and in the central countries of Europe. It will be remembered that owing to the influence of the warm Atlantic water the temperature of the whole of the north-west of that continent is raised considerably higher than it would otherwise be. Consequently vegetation is much earlier in the south of Sweden than in the same latitudes to the east. The lilac and elm begin to show their leaves sooner at Upsala than at Paris, and while winter still reigns to the east of the Baltic, spring blossoms have already spread far up into Scandinavia.

12. The difference between an insular or oceanic and a continental climate is likewise well brought out by the fact that such evergreens as the Portugal laurel, aucuba, and laurustinus grow luxuriantly even in the north of Scotland, while they cannot withstand the severe cold of the winter at Lyons.

13. (iii.) **Height above the Sea.**—In Lesson IX. Art. 20 reference was made to the gradual diminution of temper-

ature with increase of elevation above the sea. This cause of variation in climate is of a more local character than the two already illustrated. But its effects upon the spread of plants and animals is singularly well marked. If the fall of the thermometer be taken at 1° Fahr. for every 300 feet of ascent we can readily perceive that the times of flowering and ripening of the same plants must become later in proportion to the height of their place of growth, until at last the ground is too high and bleak to let them ripen at all before the winter sets in. They have thus an upper limit beyond which they cannot extend. But while they disappear, other plants, better able to withstand the rigorous climate of the uplands, take their place. As we climb to higher elevations the familiar vegetation of the plains is gradually succeeded by a vegetation peculiar to the mountains, until at last we reach the edge of the snow line. This influence of height on vegetation is illustrated in a graphical form by Fig. 78.

On the animal world, too, the influence of elevation may be distinctly seen. In Europe rabbits, moles, hedgehogs, otters, foxes, larks, thrushes, lapwings, and many other common forms occur among the lower grounds, while in the mountains such animals as the marmot, goat, ibex, chamois, brown bear, and eagle find their congenial home.

14. (iv.) Prevailing Winds.—Air lying upon the surface of any part of the globe tends to acquire the temperature of that surface. Consequently winds which come from a cold region are cold, those from a warm region are warm. Winds from the sea are usually moist, those from the land are generally dry. Sea-breezes are not liable to the same extremes of temperature as those from the land. The vapour which they carry with them cools the heat of summer and lessens the cold of winter. On the other hand, winds blowing from the interior of a continent are

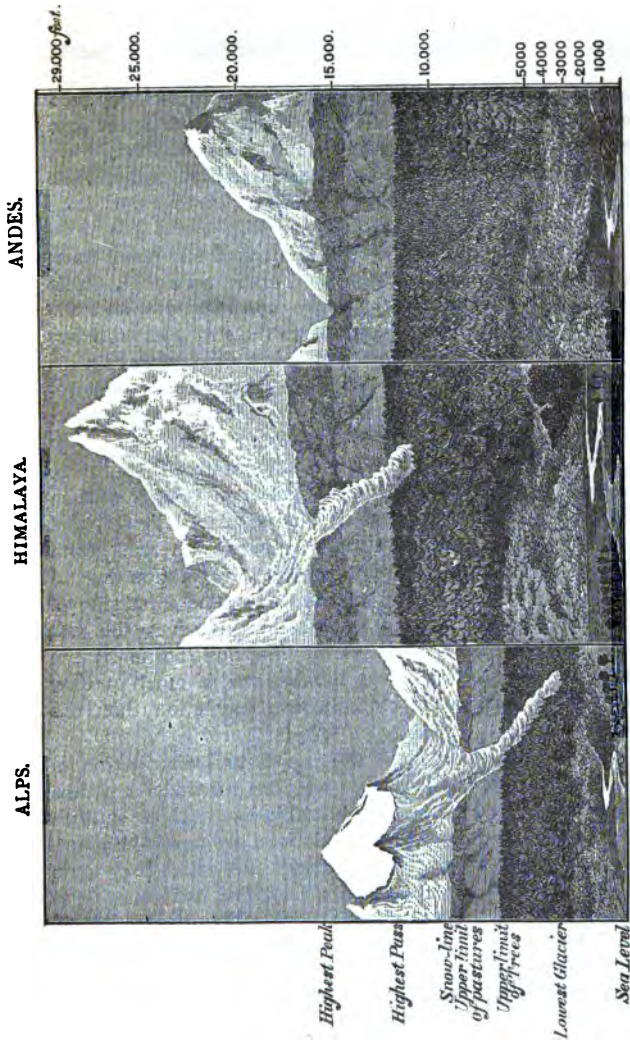


Fig. 78.—Vertical distribution of climate on mountains.

apt to be hot and suffocating in summer, piercingly cold and dry in winter. Winds which come from lower into higher latitudes, or from warmer to cooler climates, have their moisture condensed, and are therefore rainy, while those which blow from higher to lower latitudes, or from cold to warm regions, are dry.

15. Much, therefore, in the climate of any place must be due to the prevailing winds. This is more particularly noticeable on the coasts of the continents, where the winds blow alternately from and to the sea. The striking contrasts between the extremely rainy and almost rainless districts in certain parts of India have already been referred to as showing the great influence of the winds in determining the moisture of a climate.

16. It is evident that, as regards plant-growth, moisture is hardly less important an element than temperature in the climate. Those tracts are most plentifully covered with vegetation which are most copiously watered. Both the character and the abundance of the vegetation depend greatly upon the amount of rain-fall. The west side of the British Islands, for example, which receives the first and largest precipitation of the moisture borne by the south-westerly breezes from the Atlantic, is much greener and more luxuriantly clothed with vegetation than the east side.

17. Whatever regulates the growth and distribution of plants must tell effectually upon the spread of animals. The herbivorous species naturally haunt those regions where their supplies of vegetable food are most abundant. In their train come the predatory kinds which prey upon them. Any change of climate, therefore, unfavourable to the vitality of the pasture will drive away or even locally exterminate the herds of plant-eating animals, and when they disappear the beasts of prey must vanish also.

18. (v.) **Local Influences.**—Various minor causes of a

more local kind help to modify the climates of different places, and thereby to affect the flora and fauna. The nature of the soil is one of the most important of these. Wet, marshy ground lowers the mean temperature, seeing that its water absorbs and conveys downward the heat which would otherwise warm the soil. Consequently the effect of drainage is to raise the mean annual temperature. In Britain this increase amounts sometimes to as much as $1\frac{1}{2}$ °—3° Fahr., which is as great a change as if the ground had actually been transported 100 or 150 miles further south. A waste of sand presents the greatest extremes of climate, for while the dry surface readily absorbs the sun's heat so as to rise even to 200° Fahr. during the day, it cools rapidly by radiation, and during a clear night may grow ice-cold.

19. A surface of vegetation prevents the soil from being as much warmed and cooled as it would be if bare, and as leaves never become so hot as soil, they equalise the temperature. A large mass of forest thus exercises a well-marked influence on the climate of the region on which it lies, tempering alike the heat of the day and the cold of night.

20. Similar effects are produced by lakes and other large inland surfaces of water. The surface water chilled by the cold of winter descends to the bottom, leaving a warmer layer at the top, which, cooled in its turn, sinks down and allows another warmer portion to lie at the surface. By this means the temperature of the air overlying the water is kept above that of the air overlying the adjoining land, while the colder air from all sides flows down to the lake and is there warmed. Many deep lakes do not freeze in winter, and then serve as reservoirs of warmth to keep the temperature of the surrounding ground higher than that of places only a short distance away. On the other hand, during summer the water cools the air

lying upon it, and thereby lessens the heat of the locality.

21. One other local cause affecting climate may be referred to, viz., proximity to hills or mountains. The influence of high ground shows itself in augmenting rain-fall (Lesson X. Art 29), in producing currents of air, which move alternately up and down the valleys (Lesson XI. Art. 8), giving rise to gusts and blasts of cold wind, which rush down to the plains.

22. It was formerly imagined that each climate had its own characteristic forms of life, and that the boundaries between the different botanical and zoological regions were as ancient and as well defined as between the various climates. But while similarity of climate does not always bring similarity of vegetation and of animals, the want of resemblance between the plants and animals of two distant countries having similar climates does not arise from any unfitness in the one country for the organisms of the other. Cattle and horses introduced by the Spaniards into South America now run wild there in vast herds. The rat, originally not a native of America, may now be found in all parts of the continent. Hogs, goats, cats, and dogs, first brought into the New World by Columbus and his successors, are to-day found running wild in great numbers. In Australia, too, the domestic animals introduced by the colonists are rapidly supplanting the kangaroos and other aboriginal forms. A fresh-water plant accidentally imported from America has spread rapidly over England, and is choking up canals and the channels of rivers.

23. We must conclude that similar climates may have remarkably dissimilar assemblages of plants and animals if they are sufficiently isolated from each other, that such differences are not referable to the influence of climate, for the plants and animals have been found to increase

rapidly when transported to a distant but similar climate, and that while climate has evidently an important influence in the distribution of life over the globe, it is not sufficient to account for all that we see.

II.—Migration and Transport.

24. It might be supposed that the present plants and animals first appeared in one region or continent, whence they gradually spread over the whole of the globe. No doubt many species are endowed with remarkable powers of diffusing themselves and of living even vigorously under the greatest extremes of climate. But further consideration suffices to convince us that this explanation is wholly incapable of accounting for the existing arrangement of the faunas and floras of the earth.

25. Plants have many facilities for spreading themselves. Their seeds are often swept up into the air by whirlwinds, and may be carried along for hundreds of miles before being dropt again to the ground. Should the soil, climate, and other conditions be favourable, these transported seeds may take root and spread over their new abode. In other cases, seeds may be borne for long distances over the sea, either floating by themselves or inclosed among earth and leaves in masses of drift-wood. Cast up at last on some remote shore, they sometimes find a fitting home and take root there. To the feathers of birds and the fur of animals seeds must often adhere, and may thus be carried far away from their original source. Seeds which have been for many hours in the crops of birds have been found to be still alive. It may now and then happen, therefore, that when, after flying across hundreds of miles of sea the birds fall a prey to other predatory members of the feathered race, seeds falling from their torn crops find a lodgment in the earth, where eventually they

spring into leaf. In these and other ways many kinds of plants may have spread far beyond their original bounds. Yet at the best these are but limited means of transport. Differences of climate and soil, lofty mountain-chains, and intervening seas have placed barriers in the way of such diffusion which comparatively few species of plants can ever surmount.

26. Animals enjoy greater facilities for dispersal, since their movements are voluntary as well as involuntary. On some of the large tropical rivers rafts of drift-wood are now and then to be seen with monkeys and other wild animals upon them, all sailing down the current on their way to the ocean. In the great majority of cases rafts of this kind are broken up at sea, and their unfortunate denizens are drowned. But cases have been known where the animals have actually found their way to land. We may suppose therefore that islands in mid-ocean may sometimes have had both plants and animals introduced into them by these means. Again, in the Arctic seas, polar bears have been noticed upon icebergs at a great distance from land; so that by drifting ice as well as by floating vegetation, animals may be diffused from one country to another. But here again the means of transport are so scanty and the chances of the animals being able to live and multiply in their new home are so small, that we may be sure it is not in this way that the continents have been peopled.

27. While most animals live within tolerably well-defined limits, marked out by the climate and the kind of vegetation which the climate supports, some species have great powers of migration, and when impelled by their migratory instinct, whether from stress of hunger, or from change of season, will travel for hundreds or thousands of miles. In North America many remarkable instances are on record of the hordes of bisons, beavers, and squirrels

which from time to time have quitted their previous haunts in search of a new home. Birds show this instinct strongly. Many of the most familiar birds of the temperate region, both in the Old and the New World, are migratory. They go north in summer to breed, and after spending some months in a cooler climate, and seeing their young brood able to fly, they again take wing and return to their winter quarters in the south. In Europe the swift, swallow, and cuckoo wing their way in summer even far up within the Arctic Circle, but before winter has set in they have crossed the Mediterranean to the milder air of Northern Africa.

28. But while a limited number of animals are fitted to spread over wide regions and to endure great diversities of climate, the vast majority are confined within their own district, beyond which they cannot stray, not because they are in all cases unfitted for distant journeys, but because of the insuperable barriers to their advance. Of these obstacles the most potent is no doubt climate, which acts not only on the animals themselves, but on the vegetation, which directly or indirectly furnishes their food. Some species of animals can live only in woods, others cannot stray far from the marshes and jungles, where alone they find their proper support ; some are adapted for life solely in the moist hot air and luxuriant vegetation of the tropics ; others find their congenial home among arctic snows.

29. Even, however, where an animal is endowed, like the tiger, with extraordinary powers of accommodating itself to wide extremes of temperature and great variety of food, it is surrounded by many obstacles to its diffusion. A lofty snow-covered mountain-chain may effectually prevent it from crossing into districts where, if it could once reach them, it would find abundance of food and shelter. A strip of barren arid desert is another efficacious barrier,

preventing the animals of one province from passing over into another. An arm of the sea or strait only a few miles in breadth suffices to keep the plants and animals of the opposite shores distinct ; while of course a wide and deep ocean is an insuperable barrier.

30. Let us suppose, however, that through some exceptionally favourable circumstances, a few animals of one or more species have succeeded in crossing one of these natural barriers, what are the chances that they will be able to establish themselves on the further side? The climate must be one in which they can live and increase. They must be able to find enough of their proper food. If herbivorous they must needs find vegetation fitted to support them ; and if carnivorous they will require to meet with animals less powerful than themselves in numbers sufficient to yield them subsistence. It would seldom happen that the invaders would have all these chances in their favour. But if they were able to maintain themselves at first, they would soon find their advent opposed by some rival species already long established and numerous. In the struggle that would follow the new comers could seldom make good their hold in the country, and unless able to return to their original haunts would in most cases perish.

31. The vast majority of animals are thus hemmed in by barriers—climate, food, mountains, deserts, wide rivers, seas, or rival species—barriers which, whether seen or unseen, effectually restrain them from spreading beyond the limits of their own district or region. We are therefore forced to conclude that most of the present species of animals (and the same holds true for the great majority of plants) cannot have spread from any one common centre, but from their very nature and requirements must always have been restricted in their distribution, generally to the same tracts in which they now live. While certain forms

of plant and animal have been able to diffuse themselves over almost the whole globe, the flora and fauna of each of the great biological regions remain distinctly marked out by more or less definite boundaries. That these regions have in every case had a long history, and that their existing species of plants and animals have been preceded by other and different species, is shown by the rocks which form the land, and the traces both of vegetable and animal life found in these rocks. In trying to discover how and whence the continents have received their mantle of vegetation and their hosts of animals, science needs to grope backwards among the records contained in the rocks, and which form the subject of Geology. To some aspects of this question, which show how closely Physical Geography is linked with Biology, how the plants and the animals of a continent may be made to tell a part of the ancient history of the land on which they live, the concluding pages of these Lessons may fittingly be given.

III.—Changes of Land and Sea and of Climate.

32. From the facts recorded in the preceding Chapters it appeared that the present heights and hollows of the land have not always existed,—that the continents have been uplifted at different times from the bed of the sea, that each mountain-chain has been ridged up and altered at various periods, and that the valleys have been slowly deepened and widened by the rivers flowing in them. We cannot suppose that such important changes could take place without affecting more or less potently the various forms of plants and animals. Investigation shows that the present distribution of life sometimes bears striking and independent witness to these changes. How this is

made evident will be clear from one of the simpler illustrations.

33. The plains of Central Europe up to the shores of the Baltic are clothed with a vegetation which has one common character throughout. Many of the plants are of course local, but a vast number range far and wide over the region. Crossing from the continent into Britain we meet with the same general assemblage of plants, and as the greater number of these could not have been drifted across the intervening sea, but must have travelled by land, they show that originally Britain formed a part of the European continent, and that its separation into islands has taken place since the present species of plants spread over its surface.

34. Any one who ascends the higher hills in Britain, or a part of the mountain-chain of the Alps or the Pyrenees, finds that as he reaches higher elevations he loses the common and characteristic plants of the plains. The vegetation as it becomes less luxuriant assumes a more and more distinct type, not only by the disappearance of lowland species, but by the occurrence of others, such as peculiar gentians and saxifrages never seen below, but plentiful on the higher hill-tops and mountain-slopes. So general is this change that every hill or mountain in Central and Northern Europe, rising high enough to reach the fitting climate, may be expected to contain more or less fully this alpine flora. And this is found to be the case even when the groups of mountains are separated by wide intervals of low country. The Scottish Highlands furnish on their higher slopes an abundant growth of alpine forms of vegetation. Crossing the Lowlands, where none of these plants occur, we again meet with them on some of the more elevated summits in the Cheviot Hills. After another interval they reappear on the hill-tops in the Cumberland

lake-district, and again on the higher mountains of Wales. Across the whole breadth of England they are absent from the low grounds. They are not to be found on the opposite shores of the Continent. But far to the south they reappear abundantly on the tops of the Pyrenees and below the snow line along the whole chain of the Alps.

35. We must not suppose these plants to be merely species peculiar to lofty elevations and always found there. They do not occur on mountains lying to the south of the Palæarctic region. On the Peak of Teneriffe, for instance, they are absent, though the climate and soil would have been well fitted for them. On comparing the heights at which they are met with we perceive that they approach nearer the sea-level the further north we trace them. In the Alps they grow in the zone between the upper limit of trees and the snow-line, or at a height of from 6,000 to 10,000 feet above the sea. In the Scottish Highlands they descend to 2,000 feet, or even lower. In Scandinavia they come down to the sea-level, and grow in such vigour and abundance there as to show that they are really northern or arctic plants.

36. How then did an arctic flora overspread the mountains so far south as the Alps? It could not do so at present. The intervening low grounds are clothed with an abundant vegetation of another type, across which the northern plants could not force their way. To enable them to advance southwards this lowland vegetation would need to be removed. If the climate of Central Europe were to grow as severe as that of Norway or of the higher Alps, the effect of the change would be to kill the vegetation of the plains, or as much of it as could not endure the greater cold. At the same time the arctic plants, finding their congenial temperature extending across

the European plain, would gradually spread southwards, ascend the hills, and finally become the dominant vegetation over the whole region where the arctic climate prevailed. This seems to be the only way in which the presence of hundreds of northern plants upon the Alps and Pyrenees can be satisfactorily explained.

37. Is there then any corroborative evidence that such a wonderful change in the aspect of Europe did really take place? Undoubtedly there is. Below the soil in different parts of France and England, as well as in the deposits covering the floors of caves, bones of reindeer have been found in considerable abundance. This we know is an arctic animal. Remains also of the musk-sheep, glutton, arctic fox, lemming and others, which are all northern forms, have been exhumed from similar situations. So that there can be no doubt that at one time characteristic animals of the arctic regions roamed over Europe as far at least as the south of France. We cannot doubt that this could have happened only through some change of climate which, driving out the usual denizens of the plains and forests, allowed the northern animals to occupy their place.

38. But further and abundant proof of an extremely rigorous climate having overspread Europe is supplied by the polished rocks and heaps of earth, which were described in Lesson XXVIII. as part of the work of glacier-ice. Traces of former glaciers are found throughout the more hilly districts of Britain. Similar traces occur both in Norway and among the Alps and Pyrenees, far beyond the limits of the present glaciers, showing that the ice formerly extended in vast sheets into the plains.

39. It was during this cold period that the arctic plants and animals overspread Europe. Since that time the climate has gradually ameliorated. Step by step the

northern plants have been driven out of the plains and up into the mountains, where among the congenial frosts and snows they are able to maintain themselves in scattered colonies. The animals have long since been pushed back into the icy north.

40. In North America a similar record has been preserved. As far south as the White Mountains of New Hampshire in the United States (lat. 45°) the summits are peopled with Labrador plants, which once no doubt extended over the low grounds up into the northern lands, where the same species are now found abundantly down to the sea-level.

41. This illustration—by showing how far the present distribution of plants and animals may be from that which once existed, and also how distinctly groups of plants or of animals may sometimes tell of former changes in physical geography—may serve to indicate why the problem of accounting for the existence and differences of the biological regions should be so difficult. So many causes have to be considered and so much knowledge is needed regarding the events which preceded the present state of things. Much has been done by naturalists in this department of research during recent years, but they have as yet only entered upon the beginning of the inquiry. The story of the peopling of each of the great regions with plants and animals may never be fully told. But that it will be made far fuller and clearer than it is, and that it will help to illustrate the history of the continents themselves, cannot be doubted.

42. One grand object of science is to link the present with the past, to show how the condition of the globe today is the result of former changes, to trace the progress of the continents back through long ages to their earliest beginnings, to connect the abundant life now teeming in

air or land and in the sea with earlier forms long since extinct, but which all bore their part in the grand onward march of life now headed by man ; and thus, learning ever more and more of that marvellous plan after which this vast world has been framed, to gain a deeper insight into the harmony and beauty of creation, with a yet profounder reverence for Him who made and who upholds it all.

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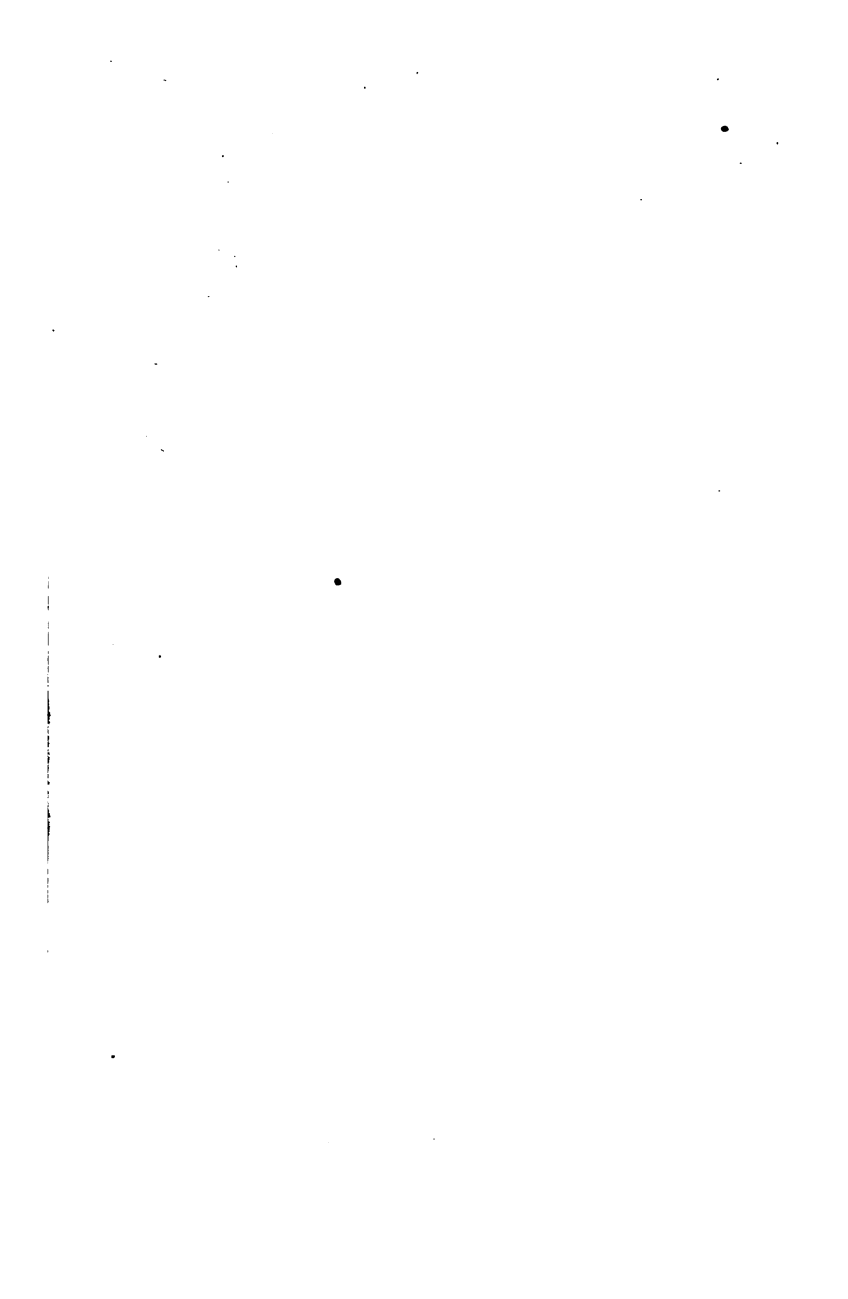
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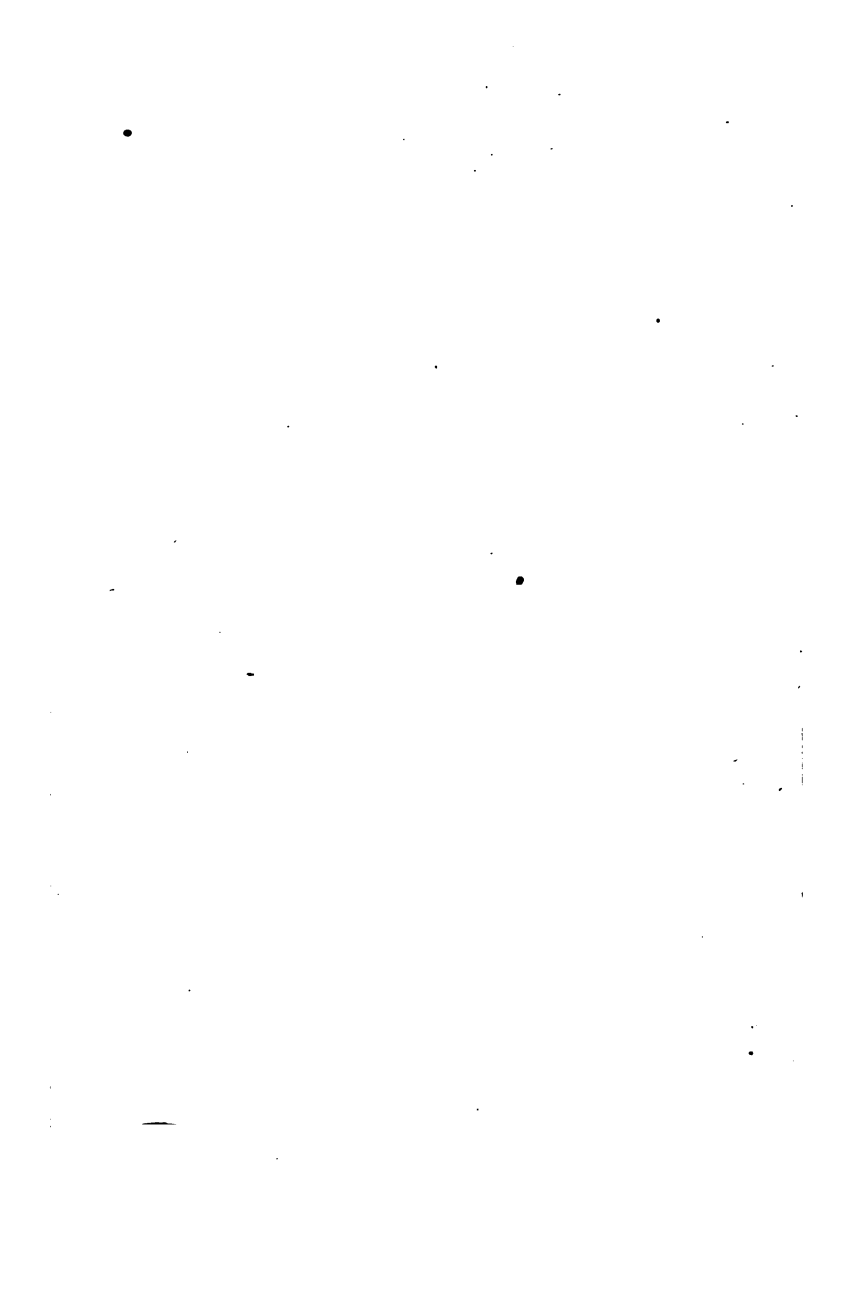
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